

Global change pressures on soils from land use and management

Article

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Global Change Pressures on Soils from Land Use and Management

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57

58

59 **Abstract**

60

61 Soils are subject to varying degrees of direct or indirect human disturbance, constituting a
62 major global change driver. Factoring out natural from direct and indirect human influence is
63 not always straightforward, but some human activities have clear impacts. These include land
64 use change, land management, and land degradation (erosion, compaction, sealing and
65 salinization). The intensity of land use also exerts a great impact on soils, and soils are also
66 subject to indirect impacts arising from human activity, such as acid deposition (sulphur and
67 nitrogen) and heavy metal pollution. In this critical review, we report the state-of-the-art
68 understanding of these global change pressures on soils, identify knowledge gaps and
69 research challenges, and highlight actions and policies to minimise adverse environmental
70 impacts arising from these global change drivers.

71

72 Soils are central to considerations of what constitutes sustainable intensification. Therefore,
73 ensuring that vulnerable and high environmental value soils are considered when protecting
74 important habitats and ecosystems, will help to reduce the pressure on land from global
75 change drivers. To ensure that soils are protected as part of wider environmental efforts, a
76 global soil resilience programme should be considered, to monitor, recover or sustain soil
77 fertility and function, and to enhance the ecosystem services provided by soils. Soils cannot,
78 and should not, be considered in isolation of the ecosystems that they underpin and vice
79 versa. The role of soils in supporting ecosystems and natural capital needs greater
80 recognition. The lasting legacy of the International Year of Soils in 2015 should be to put
81 soils at the centre of policy supporting environmental protection and sustainable
82 development.

83 **1. Introduction**

84

85 2015 is the International Year of Soil. This represents an ideal time to take stock of scientific
86 knowledge about the changing global pressures that humans are exerting on soils. 2015 is
87 also the year when policy makers will adopt a new legally-binding climate agreement under
88 the United Nations Framework Convention on Climate Change (UNFCCC), with individual
89 countries and businesses making policies and targets on greenhouse gas emissions and
90 removals. Soils storage and cycling of carbon and nitrogen are part of emissions and
91 removals from the land sector. Furthermore, 2015 is the year when countries will shape and
92 adopt a new development agenda that will build on the Millennium Development Goals
93 (MDGs). With increasing population, issues such as food security, water security, energy
94 security (including bioenergy production) and sustainable integrated land and resource
95 management are central to many development research and policy agendas. Soils underpin
96 the provision of many ecosystem services related to development.

97

98 Soils provide multiple ecosystem services, allowing sustained food and fibre production, and
99 delivering climate regulation, flood regulation, improved air and water quality, reducing soil
100 erosion, and provide a reservoir for biodiversity (Smith et al. 2015). All soils are subject to
101 some degree of human disturbance, either directly through land-use and land management, or
102 indirectly through responses to human-induced global change such as pollution and climate
103 change. Distinguishing natural from direct and indirect human influence is not always
104 straightforward (Smith, 2005), but some human activities and their consequences have clear
105 impacts, and despite large heterogeneity in soil properties and responses, robust scientific
106 knowledge exists.

107

108 Human impacts on soils largely emerge from the need to meet the food, fibre, and fuel
109 demands of a growing population including an increase in meat consumption as developing
110 nations become wealthier, the production of biofuels, and increasing areas of urbanization.
111 This has led to conversion of natural land to managed land (extensification) and
112 intensification of agricultural and other management practices on existing land such as
113 increasing nutrient and water inputs and increasing harvest frequency to increase yields per
114 hectare.

115

116 Land cover or land use change (e.g. from forest or natural grassland to pasture or cropland),
117 removes biomass, changes vegetation and disturbs soils, leading to loss of soil carbon and
118 other nutrients, changes in soil properties, and changes to above- and below-ground
119 biodiversity. Some land cover conversions e.g. reforestation after abandonment of cropland,
120 can increase both above- and below-ground carbon and nutrients. Land use or land
121 management that does not result in a change of cover (e.g. forest harvest and regrowth,
122 increased grazing intensity and intensification of crop production), can potentially result in
123 degradation of soil properties, depending on the characteristics of the management practices.

124

125 Land use change has been accelerated by population increases and migration as food, shelter,
126 and materials are sought and acquired. It is estimated that humans have directly modified at
127 least 70 Mkm², or >50 percent of Earth's ice-free land area (Hooke et al. 2012). The new
128 Global Land Cover Share-database (Latham et al., 2014) represents the major land cover
129 classes defined by the FAO. Croplands and grasslands (including both natural grasslands and
130 managed grazing lands) each covered 13.0 %. "Tree-covered areas" (i.e. both natural and
131 managed forests) covered 28%, shrub-covered areas 9.5 %. Artificial surfaces (including
132 urbanised areas) occupy 1 %. Land degradation can be found in all land cover types.

133 Degraded land covers approximately 24% of the global land area (35 Mkm²). 23% of
134 degrading land is under broadleaved forest, 19% under needle-leaved forests and 20-25% on
135 rangeland (Bai *et al.*, 2008).

136

137 In this review we report the state-of-the-art understanding, and knowledge gaps concerning
138 impacts of changes in anthropogenic land use and land management on soils, including
139 interactions with other anthropogenic global change pressures. We also review actions and
140 policies that limit the adverse impacts arising from these global change drivers. We make the
141 case to put soils at the centre of research strategy and policy actions as a legacy of the
142 International Year of Soils.

143

144 **2. Land use/land cover change**

145

146 Land cover change has been dominated by deforestation, but also conversion of grasslands to
147 cropland and grazing land. Deforestation has had the greatest impact on historical soil carbon
148 change, causing on average around 25% of soil carbon to be lost (Guo & Gifford, 2002;
149 Murty *et al.*, 2002). Soil carbon losses largely stem from oxidation of the organic matter as

150 well as soil erosion.

151

152 Deforestation affected an estimated 13 million hectares per year between 2000 and 2010; net
153 forest loss was 5.2 million hectares per year (FAO, 2010). Most of this recent deforestation
154 has taken place in tropical countries (FAO, 2010; Hansen et al., 2013). Over 50% of tropical
155 forest loss occurred in Brazil and Indonesia, largely driven by a few commodities: timber,
156 soy, beef, and oil palm (West *et al.*, 2014). There has been a reduced rate of deforestation in
157 some regions over the last decade, most notably Brazil (INPE, 2014), largely because of land
158 use conservation policies (Soares-Filho *et al.*, 2014; Nolte *et al.*, 2013) as well as economics.
159 Most developed countries with temperate and boreal forest ecosystems – and more recently,
160 countries in the Near East and Asia – are experiencing stable or increasing forest areas in
161 contrast to the large scale historic deforestation in these regions, with afforestation reported in
162 Europe, USA, China, Vietnam and India (FAO 2013).

163

164 Changes in soil properties can vary markedly with type of land cover change, climate, and
165 method, extent of vegetation removal (e.g. land clearing, fires, mechanical harvest) and
166 management post-harvest. For example, West *et al.* (2010) estimated that clearing land in the
167 tropics generally emits three times the amount of carbon per ton of annual crop production
168 compared to clearing land in temperate areas. Emissions are particularly high when organic
169 peatland/wetland soils are drained to enable agriculture as the initial soil carbon is higher, and
170 drainage results in large losses of carbon as previously anaerobic soils become aerobic,
171 allowing the organic matter to oxidise. For example, clearing forest on organic soils for palm
172 oil production in Kalimantan emits nine-times more carbon than clearing on neighbouring
173 mineral soils (Carlson & Curran, 2013). Impacts of deforestation can be reduced by avoiding
174 deforestation on organic soils, and on steep slopes prone to erosion.

175

176 There is large heterogeneity in soil measurements of carbon, nitrogen, microbes etc., and still
177 many areas of the world with poor data coverage. Models can be used to fill gaps in spatial
178 coverage and look at past and future time periods, but these too give very variable results.
179 Nevertheless there are some clear signals that can be obtained from meta-analyses of field
180 data and global model results of land use/land cover change with respect to soil carbon.

181

182 ***2.1. Observations of impacts of land cover change***

183

184 Table 1 presents the results of different meta-analysis studies across different climatic zones
185 that compared the impacts of land use changes on SOC (Guo & Gifford 2002; Don *et al.*
186 2011; Poeplau *et al.* 2011; Bárcena *et al.* 2014; Murty *et al.* 2002; Wei *et al.* 2014a). Changes
187 in SOC after the conversion of forests to croplands ranged from -24 to -52% without marked
188 differences between climatic regions. The conversion of pastures to other uses (tree
189 plantations and particularly, croplands) also induced decreases in SOC (-10% and -59%,
190 respectively). On the other hand, the substitution of croplands by other land uses (forest
191 regrowth, tree plantation, grassland, pasture) resulted in an increase of SOC (+18 to +53%).
192 In the case of afforestation, soil C increase with time after afforestation, and C sequestration
193 depends on prior land use, climate and the tree species planted.

194

195 Fewer meta-analysis studies are available for changes in soil N with changes in land uses. A
196 compilation with predominance of data from tropical sites indicated that average loss of 15%
197 of soil N after conversion of forests to croplands (Murty *et al.* 2002). In Australia, N losses
198 after conversion of native vegetation to perennial pasture and cropland were more than 20%
199 and 38%, respectively (Dalal *et al.* 2013) while in China N loss (0-10 cm depth) was 21%
200 and 31% after 4 and 50 years after conversion of forests to cropland (Wei *et al.* 2014b).
201 Similarly to what was described for SOC, afforestation in subtropical zone results in a
202 significant increase of N stocks 50 years after conversion (Li *et al.* 2012).

203

204 [Table 1 here]

205

206 ***2.2. Modelled impacts of land cover change***

207

208 Dynamic Global Vegetation Models (DGVMs) are used to look at the combined effects of
209 land use change, climate, CO₂, and in some cases N deposition, on vegetation and soil
210 properties over time. A few global models include some aspects of forest, grassland or
211 cropland management (Bondeau *et al.* 2007; Lindeskog *et al.* 2013; Drewniak *et al.* 2013;
212 Jain *et al.* 2005). Most DGVMs do not currently model peatland soils. In Tables 1 and 2, and
213 Figures 1 and 2, we show impacts of past land cover and management change on soil carbon
214 and nitrogen as calculated by three DGVMs: ISAM (Jain *et al.* 2013; El-Masri *et al.* 2013;
215 Barman *et al.* 2014 a,b); LPJ-GUESS (Smith *et al.* 2001; Lindeskog *et al.* 2013); and LPJmL
216 (Bondeau *et al.* 2007). The ISAM and LPJ-GUESS models were run with the HYDE
217 historical land use change data set (History Database of the Global Environment; Klein

218 Goldewijk *et al.* 2011). ISAM included wood harvest following (Hurtt *et al.* 2011). The
219 LPJmL group combined 3 land use change data sets with the geographic distribution of
220 global agricultural lands in the year 2000. All models were run with historical climate and
221 CO₂, and additionally N deposition in the ISAM model only as it includes a nitrogen cycle.
222 The effects of land cover change were isolated by comparing model runs with and without
223 land use/management (Le Quéré *et al.* 2014). Table 2 and Figure 1 show the loss of soil
224 carbon due to historical land use change from 1860 to 2010 (note there was land use change
225 causing soil carbon loss prior to 1860 particularly in Europe and central Asia, but results are
226 not shown as they were not available for all three models). As with the observed data (Table
227 1) high carbon losses are associated with the conversion of forests to croplands. Figure 2
228 shows the mineral soil C and N concentration of different land cover types in different
229 geographic ranges.

230

231 [Figure 1 & 2; Table 2 here]

232

233 Differences between the models are large for some systems and regions due to different land
234 use change data, different land cover definitions, and different processes included in the
235 models. For example, soil carbon losses are higher in the LPJmL model (Table 2, Figure 1) in
236 part due to greater land cover change in their land cover reconstructions, while their boreal
237 grassland soil carbon is high due to the inclusion of permafrost slowing soil carbon
238 decomposition (Figure 2). Treatment of management processes turns out to be an important
239 differentiator. ISAM shows strong decreases of soil carbon in some regions e.g. the southern
240 Boreal zone (Figure 1) where the inclusion of wood harvest removes carbon and nutrients
241 from the soil, while increases in soil carbon in parts of the mid.-latitudes are due to regrowth
242 of forest following abandonment of agricultural land.

243 In semi-arid to arid regions, LPJ-GUESS and LPJmL show opposite signs of soil carbon
244 change after conversion of natural land to pastures (Figure 1), primarily because LPJ-GUESS
245 simulates a greater fraction of woody vegetation than LPJmL in these regions under potential
246 natural vegetation. Conversion of woody vegetation to pasture slightly increases soil carbon
247 (see the meta analysis of Guo & Gifford 2002), partly because of boosted productivity and
248 higher turnover rates adding more C to the soil, while the change from potential natural
249 grassland to managed pasture (for which the literature is sparse) results in a soil carbon

250 decrease in LPJmL. Pasture management strategies can have a large influence on the soil
251 carbon storage (see Section 4.3), and may also be partly be responsible for differences.

252 Vegetation models are embedded in Earth System Models (ESMs) used to project future
253 climates under different human activity including different land management. Some
254 significant differences between future model climate projections stem from the differences in
255 modeling soil carbon, in particular, the strength of the relationship between increasing
256 temperatures and the increasing rate of soil carbon decomposition (Q_{10}) causing climate-
257 carbon feedbacks *via* CO₂ emissions (Friedlingstein *et al.* 2006). A recent intercomparison of
258 11 ESMs used in the IPCC 5th Assessment Report (Todd-Brown *et al.* 2013), found the
259 estimate of global soil carbon from ESMs ranged from 510 to 3040 PgC across 11 ESMs
260 compared to an estimate of 890-1600 PgC (95% confidence interval) from the Harmonized
261 World Soil Data Base (FAO/IIASA/ISRIC/ISSCAS/JRC, 2012), with all models having
262 difficulty representing the spatial variability of soil carbon at smaller (1 degree) scales
263 compared to empirical data. In all models NPP and temperature strongly influenced soil
264 carbon stocks, much more so than in the observational data, and differences between models
265 was found to be largely due to the representation of NPP and the parameterization of soil
266 decomposition sub-models. A similar, systematic analysis of DGVMs including
267 benchmarking with observational data, and careful testing of assumptions and process
268 representations in these models, making use of the very large number of observations that
269 have become available in the years since these algorithms were formulated (e.g. Medlyn *et al.*
270 2015), could significantly improve model performance. This, along with better representation
271 of critical biological and geochemical mechanisms would improve model capability (Todd-
272 Brown *et al.* 2013).

273

274 ***2.3 Drainage and conversion of peatlands/wetlands for agriculture***

275

276 The organic soils in peatlands/wetlands store vast quantities of carbon which decomposes
277 rapidly when they are drained for agriculture or commercial forestry, resulting in emissions
278 of CO₂ and N₂O to the atmosphere (Hooijer *et al.*, 2010). Other services, in particular water
279 storage and biodiversity, are negatively impacted. Drainage increases vulnerability to further
280 losses through fire. The majority of soil carbon is concentrated in peatlands in the boreal zone
281 and tropical peatland forests in Southeast Asia. These areas, along with wetlands along the
282 banks of rivers, lakes and estuaries have increasingly been developed for croplands/bioenergy

283 production over recent decades. The FAO emissions database estimates that globally there
284 are 250 000 km² of drained organic soils under cropland and grassland, with total GHG
285 emissions (N₂O plus CO₂) of 0.9 Pg CO₂eq yr⁻¹ in 2010, with the largest contributions from
286 Asia (0.44 Pg CO₂eq yr⁻¹) and Europe (0.18 Pg CO₂eq yr⁻¹; FAOSTAT, 2013; Tubiello *et al.*,
287 2015). Joosten (2010) estimated that there are >500 000 km² of drained peatlands in the
288 world, including under forests, with CO₂ emissions having increased from 1.06 Pg CO₂ yr⁻¹
289 in 1990 to 1.30 Pg CO₂ yr⁻¹ in 2008, despite a decreasing trend in developed countries, from
290 0.65 to 0.49 Pg CO₂ yr⁻¹, primarily due to natural and artificial rewetting of peatlands. In
291 Southeast Asia, CO₂ emissions from drained peatlands in 2006 were 0.61 ± 0.25 Pg CO₂ yr⁻¹
292 (Hooijer *et al.*, 2010). Conversion of peatlands in Southeast Asia is increasing, particularly
293 for oil palm expansion, where cleared peatlands typically emit ~9 times more carbon than
294 neighbouring mineral soils (Carlson & Curran 2013). In China, between 1950 and 2000, 13
295 000 km² of wetland soils were shifted to cultivated arable lands, which led to a SOC loss of
296 5.5 Pg CO₂, mostly from peatlands in Northeast China and Tibet (Zhang *et al.*, 2008).

297

298 Soil drainage also affects mineral soils. Meersmans *et al.* (2009) showed that initially poorly
299 drained valley soils in Belgium have lost significant amount of topsoil SOC (i.e. between ~70
300 and 150 t CO₂ ha⁻¹ over the 1960 – 2006 period), most probably as a consequence of
301 intensified soil drainage practices for cultivation purposes.

302

303 **3. Agricultural management**

304

305 To meet projected increases in food demand, crop production will need to increase by 70-
306 110% by 2050 (World Bank, 2008; Royal Society of London, 2009; Tilman *et al.*, 2011).
307 This can be achieved either through further expansion of agricultural land (extensification),
308 or through intensification of production on existing land. Intensification is widely promoted
309 as the more sustainable option because of the negative environmental consequences of land
310 expansion through deforestation and wetland cultivation (Foley *et al.*, 2011). For example,
311 Burney *et al.* (2010) estimate that intensification of production on croplands between 1961
312 and 2010 avoided the release of 590 PgCO₂eq. Increased productivity per unit land area can
313 be achieved through a variety of management practices, such as fertilization, irrigation and
314 increased livestock density, but these can lead to adverse consequences for the soil and wider
315 environment (Tilman *et al.*, 2002). Intensifying land use can potentially reduce soil fertility
316 (without additional inputs) and its ability to sustain high production, as well as soil resilience

317 to extreme weather under climate change, pests and biological invasion, environmental
318 pollutants and other pressures. Some key management practices and consequences are
319 highlighted below and summarised in Table 3.

320

321 [Table 3 here]

322

323 **3.1 Nutrient management**

324

325 Cultivation of soils results in a decline in soil nutrients (nutrient mining). Nutrient inputs,
326 from both natural and synthetic sources, are needed to sustain soil fertility and supply nutrient
327 requirements for crop production. Nutrient supply can improve plant growth which increases
328 organic matter returns to the soil, which in turn can improve soil quality (see section 3.5), so
329 balanced nutrient supply has a positive impact on soils (Smith *et al.*, 2015). Overuse,
330 however, has negative environmental consequences. Annual global flows of nitrogen and
331 phosphorus are now more than double natural levels (Matson *et al.*, 1997; Smil, 2000; Tilman
332 *et al.*, 2002). In China, for example, N input in agriculture in the 2000's was twice that in
333 1980's (State Bureau of Statistics-China, 2005).

334

335 Between 50-60% of nutrient inputs remain in agricultural soils after harvest (West *et al.*,
336 2014) and can enter local, regional, and coastal waters becoming a major source of pollution
337 such as eutrophication leading to algal blooms (Carpenter *et al.*, 1998). In many places
338 around the world, over-use of synthetic nitrogen fertilizers is causing soil acidification and
339 increased decomposition of soil organic matter, leading to loss of soil function in over-
340 fertilized soils (Ju *et al.*, 2009; Tian *et al.*, 2012).

341

342 Use of fertilisers and manures contributes to climate change through their energy intensive
343 production and inefficient use (Tubiello *et al.*, 2015). Globally, approximately 3-5% of
344 nitrogen additions are released as nitrous oxide (N₂O) to the atmosphere when both direct
345 (from soils) and indirect (e.g. downstream from nitrate leaching) emissions are considered
346 (Galloway *et al.*, 2004), and N₂O has ~300 times the radiative forcing of carbon dioxide
347 (IPCC, 2007). Recent research indicates that the relationship between nitrogen application
348 and N₂O emissions is non-linear, resulting in an increasing proportion of added N being
349 emitted, as application rate increases (Philibert *et al.*, 2013; Shcherbak *et al.*, 2014). China,

350 India, and the United States account for ~56% of all N₂O emissions from croplands, with
351 28% from China alone (West *et al.*, 2014). Overuse of nitrogen and phosphorus fertilizer can
352 contribute to eutrophication of water bodies, adversely affecting water quality and
353 biodiversity (Galloway *et al.*, 2003, 2004, 2008).

354

355 Nutrient use-efficiency can be significantly increased, and nitrate losses to water and N₂O
356 emissions can be reduced, through changes in rate, timing, placement, and type of
357 application, as well as balancing fertilization (Venterea *et al.*, 2012; Snyder *et al.*, 2014). It
358 has been estimated that current levels of global cereal production could be maintained while
359 decreasing global nitrogen application by 50% (Mueller *et al.*, 2014).

360

361 **3.2 Carbon management: reduced disturbance and organic matter additions**

362

363 Agricultural soils have the potential to store additional carbon than at present if best
364 management practices are used (Paustian *et al.*, 1997; Smith, 2008; Smith, 2012). Soil
365 organic matter content of soils can be increased through use of improved crop varieties or
366 grassland species mixtures with greater root mass or deeper roots (Kell, 2012), improved crop
367 rotations in which C inputs are increased over a rotation (Burney *et al.*, 2010), greater residue
368 retention (Wilhelm *et al.*, 2004), and use of cover crops during fallow periods to provide
369 year-round C inputs (Burney *et al.*, 2010; Poepflau & Don, 2015). Several studies report that
370 soil carbon increases in croplands under no-till management (West & Post, 2002; Ogle *et al.*,
371 2005). However, the carbon benefits of no-till may be limited to the top 30cm of soil
372 (Blanco-Canqui & Lal 2008; Powlson *et al.*, 2014). Baker *et al.* (2007) found that total soil
373 carbon was similar in non-till and conventional systems, suggesting that carbon accumulation
374 is occurring at different depths in the soil profile under different management schemes. Given
375 the larger variability in sub-surface horizons and lack of statistical power in most studies,
376 more research is needed on soil carbon accumulation at depth under different tillage regimes
377 (Kravchenko & Robertson, 2010).

378

379 Adding plant-derived carbon from external sources such as composts and biochar can
380 increase soil carbon stocks. Composts and biochars are more slowly decomposed compared
381 to fresh plant residues, with mean residence times several (composts) to 10-100 (biochars)
382 longer than un-composted organic materials (Ryals *et al.*, 2015; Lehmann *et al.*, 2015).
383 Recent developments suggest that biochar, from the pyrolysis of crop residues or other

384 biomass, can consistently increase crop N use efficiency while greatly (over 25%) reducing
385 direct N₂O emissions from N fertilizers (Liu *et al.*, 2012; Huang *et al.*, 2012), as well as
386 enhancing soil fertility (Woolf *et al.*, 2010).

387

388 **3.3 Water management**

389

390 The amount of irrigated croplands has doubled in the last 50 years and now accounts for 70%
391 of all water use on the planet (Gleick, 2003). While irrigated crops cover 24% of all cropland
392 area, they account for 34% of all production (Siebert & Döll, 2010). Irrigation is concentrated
393 in precipitation-limited areas such as India, China, Pakistan, and the USA, which account for
394 72% of irrigation water use (West *et al.*, 2014). Agricultural water-use competes with uses
395 for human and natural ecosystems exacerbating water stress in dry regions. Increased
396 irrigation has occurred in many areas of world agriculture due to the increasing frequency of
397 drought under the climate change (West *et al.*, 2014). Where irrigation increases productivity
398 (e.g. in drought prone areas), organic carbon inputs to the soils would be expected to
399 increase, increasing soil organic matter content (section 3.2).

400

401 Irrigation can increase soil salinity in dry regions with high salt content in the subsoil
402 (Ghassemi *et al.*, 1995; Setia *et al.*, 2011). Where salinization occurs, additional irrigation is
403 needed to “flush” the salts beyond the root zone of the crops, which can further exacerbate
404 stress on water resources, particularly when using underground water sources. Saline soils,
405 which have a high concentration of soluble salts, occupy approximately 3.1% (397 Mha) of
406 the world’s land area (FAO, 1995). Climate change (need for more frequent irrigation) and
407 increases in human population (increasing demand for more production) are likely to increase
408 the extent of saline soils (Rengasamy, 2008). The energy required by plants or soil organisms
409 to withdraw water from the soil or retain it in cells increases with decreasing osmotic
410 potential. As soils dry out, the salt concentration in the soil solution increases (decreasing
411 osmotic potential), so two soils of different texture may have the same electrical conductivity,
412 but the osmotic potential is lower in the soil with low water content (Setia *et al.*, 2011a;
413 Chowdhury *et al.*, 2011; Ben-Gal *et al.*, 2009). The accumulation of salts in the root zone has
414 adverse effects on plant growth activity, not only due to negative osmotic potential of the soil
415 solution resulting in decreased availability of water to plants, but also ion imbalance and
416 specific ion toxicity (Chowdhury *et al.*, 2011). Salinity affects microorganisms mainly by
417 decreasing osmotic potential, which affects a wide variety of metabolic activities and alters

418 the composition and activity of the microbial community (Chowdhury *et al.*, 2011) and
419 thereby soil organic matter decomposition.

420

421 In saline soils, SOC content is influenced by two opposing factors: reduced plant inputs
422 which may decrease SOC, and reduced rates of decomposition (and associated mineralisation
423 of organic C to CO₂) which could increase SOC content if the C input were unchanged.
424 Using a modified Rothamsted Carbon model (RothC) with a newly-introduced salinity
425 decomposition rate modifier and a plant input modifier (Setia *et al.*, 2011b, 2012), Setia *et al.*
426 (2013) estimated that, historically, world soils that are currently saline have lost an average of
427 3.47 t SOC ha⁻¹ since they became saline. With the extent of saline soils predicted to increase
428 under the future climate, Setia *et al.* (2013) estimated that world soils may lose 6.8 Pg SOC
429 due to salinity by the year 2100. Soil salinization is difficult to reverse, but salt tolerant plant
430 species could be used to rehabilitate salt affected soils (Setia *et al.*, 2013).

431

432 Water efficiency can be improved through management practices that reduce water
433 requirement and evaporation from the soil (such as adding mulch as groundcover), more
434 precise irrigation scheduling and rates, fixing leaks in dryland irrigation systems, improved
435 application technology (e.g., drip irrigation) and use of intermittent irrigation in rice paddies.
436 Given that water limitation is projected to become even more limiting in several semi-arid
437 regions, e.g. Sub-Saharan Africa, where the human population will probably increase most in
438 the future, and climate change impacts are projected to be severe, improved water harvesting
439 methods, e.g. storage systems, terracing and other methods for collecting and storing runoff,
440 are required to make best use of the limited water resource.

441

442 **3.4 Harvest frequency**

443

444 Approximately 9% of crop production increases from 1961-2007 was from increasing the
445 harvest frequency (Alexandratos & Bruinsma, 2012). The global harvested area (i.e. counting
446 each time an area is harvested) increased four times faster than total cropland area between
447 2000 and 2011 (Ray & Foley, 2013). The fraction of net primary production (NPP) extracted
448 by humans is increasing (Haberl *et al.*, 2007). Global warming is increasing the total area
449 suitable for double or even triple cropping in subtropical and warm temperate regions (Liu *et al.*
450 *et al.*, 2013). The increase results from fewer crop failures, fewer fallow years, and an increase
451 in multi-cropping.

452

453 Increasing harvest frequency can reduce soil quality by e.g. continuously removing soil
454 nutrients and increasing soil compaction through greater soil traffic, but if legumes are
455 included in rotations as harvest frequency increases, soil quality could be improved.
456 Increasing harvest frequency may require increasing pesticide and herbicide use, and
457 increased use of fertilisers contributing to pollution (section 3.1). The net effect will depend
458 on the effectiveness of the management practices followed.

459

460 ***3.5 Soil compaction***

461

462 Soil compaction causes degradation of soil structure by increasing soil bulk density or
463 decreasing porosity through externally or internally applied loads, as air is displaced from the
464 pores between the soil grains (McCarthy, 2007; Alakukku, 2012). It is the most important
465 subtype of physical soil deterioration, covering 68 Mha globally when first mapped in the
466 1990s (Oldeman et al., 1991). Compaction of agricultural soils often results from heavy
467 machinery or from animal trampling, so is more likely to occur in intensive agricultural
468 systems (machinery use and high stocking densities), and affects physical, chemical and
469 biological properties of soil. Top soil compaction can be reversed and controlled, but when
470 compaction creates impermeable layers in the subsoil, this is less easily reversed.

471

472 Subsoil compaction can disrupt nutrient water flows, which in turn can lead to reduced crop
473 yields, poorer crop quality and can give rise to increased GHG emissions, water and nutrient
474 run-off, erosion, reduced biodiversity and reduced groundwater recharge (Batey, 2009).

475 Where compaction cannot be avoided, mitigation is necessary. Biological approaches to
476 mitigation include planting deep rooted plants such as agroforestry; chemical methods
477 include fertilization (to overcome yield penalty, though not to remedy compaction); and
478 technical measures include machinery in which planting does not coincide with wheel tracks,
479 wide tyres / reduced tyre pressures to reduce pressure per unit area, and precision farming to
480 retain the same wheel tracks each year (Hamza & Anderson, 2005).

481

482 ***3.6 Livestock density***

483

484 Livestock production is projected to increase significantly in order to meet the growing
485 demand from a growing population and increase in per-capita meat consumption, with total

486 demand for meat expected to grow by more than 200 Mt by 2050 (Alexandratos & Bruinsma,
487 2012). The greatest increases in per-capita consumption are projected to be in developing and
488 transition countries (Alexandratos & Bruinsma, 2012). Since the 1970s, most increased
489 livestock production has resulted from intensification: increasing livestock density and
490 shifting to a greater fraction of livestock raised in industrial conditions (Bouwmann *et al.*,
491 2006). For example, 76-79% of pork and poultry production is industrialized (Herrero &
492 Thornton, 2013). Manure, inputs for growing feed, and soil loss from intensively managed
493 areas can be major sources of water pollution to local and downstream freshwater
494 ecosystems. Clearing natural ecosystems for new pastures, particularly in arid and semi-arid
495 regions, typically occurs on low-productivity lands with a much higher risk of soil erosion
496 and soil carbon/nutrient depletion (Alexandratos & Bruinsma, 2012), and negatively impacts
497 water storage and biodiversity. The impacts of livestock production are particularly prevalent
498 for beef production, which has a least an order of magnitude greater impact on land, water,
499 GHGs, and reactive nitrogen compared to other livestock (Eshel *et al.*, 2014; Ripple *et al.*,
500 2014). Moreover, industrial livestock production had led to an increased use of veterinary
501 medicines, antibiotics and hormones, posing potential risks to soil, water, ecosystems and
502 human health. Improved grazing management (e.g. optimised stocking density) can reduce
503 soil degradation, and thereby maintain and enhance organic matter content (McSherry &
504 Ritchie, 2013; see sections 3.2 and 4.3), and can reduce soil compaction, thereby increasing
505 infiltration and water storage and reduce risk of runoff and flooding downstream (Marshall *et*
506 *al.*, 2009).

507

508 **4. Other land management**

509

510 ***4.1 Forest management***

511

512 Logging and fire are the major causes of forest degradation in the tropics (Bryan *et al.*, 2013).
513 Logging removes nutrients and negatively affects soil physical properties and nutrient levels
514 (soil and litter) in tropical (e.g. Olander *et al.*, 2005; Villela *et al.*, 2006; Alexander, 2012)
515 and temperate forests (Perez *et al.*, 2009). Forest Fires affect many physical, chemical,
516 mineralogical, and biological soil properties, depending on fire regime (Certini, 2005).
517 Increased frequency of fires contributes to degradation, and reduces the resilience of the
518 biomes to natural disturbances. A meta-analysis of 57 publications (Nave *et al.*, 2011)
519 showed that fire caused a significant decrease in soil C (-26%) and N (-22%). Fires reduced

520 forest floor storage (pool sizes only) by an average of 59% (C) and 50% (N), but the relative
521 concentrations of these two elements did not change. Prescribed fires caused smaller
522 reductions in C and N storage (-46% and -35%) than wildfires (-67% and -69%). These
523 differences are likely because of lower fuel loads or less extreme weather conditions in
524 prescribed fires, both factors that result in lower fire intensity. Burned forest floors recovered
525 their C and N pools in an average of 128 and 103 years, respectively. Among mineral soil
526 layers, there were no significant changes in C or N storage, but C and N concentrations
527 declined significantly (-11% and -12%, respectively). Mineral soil C and N concentrations
528 were significantly reduced in response to wildfires, but not after prescribed burning.

529

530 Forest fires produce charcoal, or black carbon, some of which can be preserved over centuries
531 and millennia in soils. Dissolved black carbon (DBC) from burning of the Brazilian Atlantic
532 forest continued to be mobilized from the watershed each year in the rainy season, despite the
533 fact that widespread forest burning ceased in 1973 (Dittmar *et al.*, 2012).

534

535 A large field study in the Amazon (225 forest plots) on the effects of anthropogenic forest
536 disturbance (selective logging, fire, and fragmentation) on soil carbon pools showed that the
537 first 30 cm of the soil pool did not differ between disturbed primary forests and undisturbed
538 areas of forest, suggesting a resistance to impacts from selective logging and understory fires
539 (Berenguer *et al.*, 2014). As with deforestation, impacts of human disturbances on the soil
540 carbon are of particular concern in tropical forests located on organic soils and on steep
541 easily-eroded slopes.

542

543 **4.2 Shifting cultivation**

544

545 Shifting cultivation practices, where land is cleared through fire, have been practiced for
546 thousands of years, but recent increasing demographic pressure has reduced the duration of
547 the fallow period, affecting the system sustainability. Moreover, especially in Southeast Asia
548 where urbanisation is expanding in fertile plains, shifting cultivation is practiced in sloping
549 uplands, which are prone to soil and carbon loss by erosion (Chaplot *et al.*, 2005). A review
550 by Ribeiro Filho *et al.* (2013) reported negative impact on SOC associated with the
551 conversion stage, modified by the characteristics of the burning. Chop-and-mulch of enriched
552 fallows appears to be a promising alternative to slash-and-burn, conserving soil bulk density,
553 and significantly increasing nutrient concentrations and organic matter content compared to

554 burnt cropland, and a control forest in a study in the Amazon (Comtea *et al.*, 2012).

555

556 ***4.3 Grassland management and dryland degradation***

557

558 Grasslands, including rangelands, shrublands, pastureland, and cropland sown with pasture
559 and fodder crops, cover 26% of the global ice-free land area and 70% of the agricultural area,
560 and contain about 20% of the world's soil organic carbon (C) stocks. Grasslands on every
561 continent have been degraded due to human activities, with about 7.5% of grassland having
562 been degraded because of overgrazing (Conant, 2012). A meta-analysis (McSherry & Ritchie,
563 2013) of grazer effects on SOC density (17 studies that include grazed and un-grazed plots)
564 found higher grazing intensity was associated with increased SOC in grasslands dominated
565 by C4 grasses (increase of SOC by 6–7%), but with lower SOC in grasslands dominated by
566 C3 grasses (decrease of SOC by an average 18%). An increase in mean annual precipitation
567 of 600 mm resulted in a 24% *decrease* in the magnitude of the grazer effect on finer textured
568 soils, but on sandy soils the same increase in precipitation produced a 22% *increase* in the
569 grazer effect on SOC (McSherry & Ritchie, 2013).

570

571 Land use dynamics and climate change are the major drivers of dryland degradation with
572 important feedbacks, with changes in plant community composition (e.g. shrub encroachment
573 and decrease in vegetation cover; D'Odorico *et al.*, 2013). A review by Ravi *et al.* (2010)
574 indicated soil erosion as the most widespread form of land degradation in drylands, with wind
575 and water erosion contributing to 87% of the degraded land. Grazing pressure, loss of
576 vegetation cover, and the lack of adequate soil conservation practices increase the
577 susceptibility of these soils to erosion. The degree of plant cover is negatively related to
578 aridity, and an analysis of 224 dryland sites (Delgado-Baquerizo *et al.*, 2013) highlighted a
579 negative effect of aridity on the concentration of soil organic C and total N, but a positive
580 effect on the concentration of inorganic P, possibly indicating the dominance of physical
581 processes such as rock weathering, a major source of P to ecosystems, over biological
582 processes that provide more C and N through litter decomposition (Delgado-Baquerizo *et al.*,
583 2013).

584

585 Soil carbon dynamics in pastures strongly depend on management, with soil carbon increases
586 or decreases observed for different combinations of animal densities and grazing frequency
587 (Conant 2012; Machmuller *et al.* 2015). Different grazing strategies, especially in the semi-

588 natural dryland biomes, have large implications for vegetation and the carbon balance (Yates
589 *et al.* 2000). Under certain conditions, grazing can lead to increased annual net primary
590 production over un-grazed areas, particularly with moderate grazing in areas with a long
591 evolutionary history of grazing and low primary production but this does not always lead to
592 an increase in soil carbon (e.g. Badini *et al.* 2007); grazing, like crop harvest, fundamentally
593 leads to the rapid oxidation of carbon that would otherwise be eventually transferred to the
594 soil. It has long been recognised that the potential effects of management on carbon storage in
595 grassland and dryland soils are substantially greater than that of climate change or CO₂
596 enhancement (Ojima *et al.* 1993), and Henderson *et al.* (2015) estimated that the optimization
597 of grazing pressure could sequester 148 Tg CO₂ yr⁻¹.

598

599 **4.4 Artificial surfaces, urbanisation and soil sealing**

600

601 In 2014, 54% of the world's population was urban, and by 2050, two thirds of the global
602 population will be urban. Many regions in the world, (such as Europe and Asia) are affected
603 by migration of populations from rural area to large megacities. Africa and Asia have more
604 rural populations, but are urbanizing faster than the other regions (World Urbanization
605 Prospects, 2014). With urbanization comes land take (development of scattered settlements in
606 rural areas, the expansion of urban areas around an urban nucleus, and densification on land
607 within an urban area) and soil sealing. Soil sealing refers to the permanent covering of an
608 area of land and its soil by impermeable artificial material (e.g. asphalt and concrete), for
609 example through buildings and roads. The area actually sealed is only part of a settlement
610 area, and gardens, urban parks and other green spaces are not covered by an impervious
611 surface (Prokop *et al.*, 2011).

612

613 Sealing by its nature has a major effect on soil, diminishing many of its benefits (Tóth *et al.*,
614 2007). It is normal practice to remove the upper layer of topsoil, which delivers most of the
615 soil-related ecosystem services, and to develop a strong foundation in the subsoil and/or
616 underlying rock to support the building or infrastructure. Loss of ecosystem and social
617 services (mainly on high-quality soils) include impacts on water resources (e.g. reduction of
618 rainfall absorbed by the soil, reduction of soil water holding capacity affecting flooding), on
619 soil biodiversity when sealing prevents recycling of dead organic material (Marfenina *et al.*
620 2008), on the carbon cycle due to topsoil and vegetation removal (Davies *et al.*, 2011).

621 Sealing through expansion of urban areas can also lead to agricultural land becoming more

622 marginal since the best agricultural land around settlements is lost as they expand, with
623 agricultural land displaced to more marginal land.

624

625 Appropriate mitigation measures can be taken in order to maintain some of the soil functions.
626 In urban planning management, objectives to reduce the impact of soil sealing include: i)
627 preventing the conversion of green areas, ii) re-use of already built-up areas (e.g. brownfield
628 sites Meuser, 2010; Hester & Harrison, 2001 – though remediation of contaminated sites can
629 be costly; Maderova & Paton, 2013), iii) using (where appropriate) permeable cover materials
630 instead of concrete or asphalt supporting green infrastructure, and iv) implementation of
631 compensation measures. In order to deliver this mitigation a number of actions are necessary,
632 e.g. reduction of subsidies that act as drivers for unsustainable land take and soil sealing
633 (Prokop *et al.*, 2011), and strong collaboration between relevant public authorities and
634 governance entities (Siebielec *et al.*, 2010). Development impacts can be reduced by
635 inclusion of green infrastructure, a network of high-quality green spaces and other
636 environmental features that have a positive effect on well-being (Gill *et al.*, 2007) as well as
637 soils. In some regions, urban sprawl is exacerbated by insufficient incentives to re-use
638 brownfield (derelict, underused or abandoned former industrial or commercial) sites, putting
639 increasing pressure on greenfield land take.

640

641 Actions to alleviate pressures on soils driven by sealing fall into three categories: limiting,
642 mitigating and compensating. Actions to limit soil sealing centre around reduction of land
643 take through development of spatial urban planning and environmental protection. Mitigation
644 of soil sealing entails use of strategic environmental assessment for plans and programmes,
645 use of permeable materials and surfaces, green infrastructure within built and urban
646 environments, and natural water harvesting. Compensating soil sealing entails reclamation of
647 degraded land, re-use of extracted topsoil, de-sealing and is incentivised by land take fees and
648 application of environmental cost calculations.

649

650 **5. Anthropogenic environmental change pressures that interact with land** 651 **management pressures on soils**

652

653 In addition to the direct impacts of humans on soils *via* land use change and land
654 management, anthropogenic activity has indirect impacts through human-induced
655 environmental change, such as pollution and climate change. These interact with land

656 management. Soils provide a temporary but labile store for pollutants (Meuser, 2010).
657 Natural processes can release pollutants back to the atmosphere, make them available to be
658 taken up by plants and organisms, leached in to surface waters (Galloway *et al.*, 2008) and/or
659 transported to other areas by soil erosion (Ravi *et al.*, 2010). Pollutants disrupt natural
660 biogeochemical cycles by altering both soil quality and function through direct changes to the
661 nutrient status, acidity and bioavailability of toxic substances and also by indirect changes to
662 soil biodiversity, plant uptake and litter inputs (EEA, 2014). Soil sensitivity to atmospheric
663 pollution varies with respect to key properties influenced by geology (cation exchange
664 capacity, soil base saturation, aluminium), organic matter, carbon to nitrogen ratio (C:N) and
665 water table elevation (EEA, 2014).

666

667

668 Atmospheric pollutant deposition impacts on soils vary with respect to soil sensitivity to a
669 specific pollutant and the actual pollutant load. Sulphur, nitrogen and heavy metals are
670 released in to the atmosphere by fossil fuel combustion (e.g. power generation, industry and
671 transport) and non-combustion processes (e.g. agricultural fertilizers, waste). These pollutants
672 are transported off-site and deposited as either dry or wet deposition, which can cross
673 national borders. Deposition is enhanced in forests and with altitude because of reduced wind
674 speeds and greater precipitation, respectively, impacting remote areas. Harmful effects to soil
675 function and structure occur where deposition exceeds the ‘critical load’ that a particular soil
676 can buffer (Nilsson & Grennfelt, 1988). Spatial differences in soil sensitivity (commonly
677 defined by the ‘crucial load’) and pollutant deposition result in an uneven global distribution
678 of impacted soils (Figure 3). For instance, global emissions of sulphur and nitrogen have
679 increased 3-10 fold since the pre-industrial period (van Aardenne *et al.*, 2001), yet only 7-
680 17% of the global land area sensitive to acidification is in a region where deposition exceeds
681 the critical load (Bouwman *et al.*, 2002).

682

683 Emissions of pollutants, notably sulphur, across Europe and North America have declined
684 since the 1980s following protocols established under the 1979 Convention on Long-range
685 Transboundary Air Pollution (LRTAP) and the 1990 US Clean Air Act Amendments
686 (CAAA) (Greaver *et al.*, 2012; Reis *et al.*, 2012; EEA, 2014). Conversely, emissions are
687 likely to increase in response to industrial and agricultural development in south and east
688 Asia, sub-Saharan Africa and South America (Kuylenstierna *et al.*, 2001; Dentener *et al.*,

689 2006). Further emission increases are occurring in remote areas due to mining activity, such
690 as oil sand extraction in Canada (Kelly *et al.*, 2010; Whitfield *et al.*, 2010).

691

692 **5.1 Sulphur deposition**

693

694 Sulphur emissions are primarily from combustion of coal and oil, typically associated with
695 power generation and heavy industry. In 2001, regions with deposition in excess of 20 kg S
696 ha⁻¹ yr⁻¹ where China and Republic of Korea, western Europe and eastern North America
697 (Vet *et al.*, 2014; Figure 3a). Deposition in un-impacted areas is <1 kg S ha⁻¹ yr⁻¹ (Figure 3a).
698 Pollution control measures have seen an 80% reduction in pollutant sulphur deposition across
699 Europe between 1990 and 2010 (Reis *et al.*, 2012), and emissions in China have declined
700 since 2005 (Fang *et al.*, 2013).

701

702 Soil acidification is a natural process that is altered and accelerated by sulphur and nitrogen
703 deposition (Greaver *et al.*, 2012). Sulphur oxides (SO_x) react with water to form sulphuric
704 acid (H₂SO₄). Excess inputs of acidity (H⁺) displace soil base cations (e.g. calcium (Ca²⁺) and
705 magnesium (Mg²⁺)) from soil surfaces into solution, which are subsequently lost by leaching
706 (Reuss & Johnson, 1986). Mineral soils can buffer base cation losses if inputs from rock
707 weathering and/or atmospheric dust deposition exceed the amount lost. Therefore, the global
708 distribution of acid sensitive soils is associated with conditions that favour development of
709 soils with low cation exchange capacity and base saturation (Bouwman *et al.*, 2002; Figure
710 3c). Wetland can also buffer inputs of acidity through biological sulphate reduction, although
711 acidity can be mobilised again following drought and drainage (Tipping *et al.*, 2003; Laudon
712 *et al.*, 2004; Daniels *et al.*, 2008). Organic acids can also buffer mineral acidity in naturally
713 acidic organic soils (Krug and Frink, 1983).

714

715 Decreased soil fertility or 'sterilisation' due to loss of nutrients and mobilisation of toxic
716 metals, particularly Al, are caused by acidification. Impacts in Scandinavia over the 1960s-
717 80s included declines in freshwater fish populations and damage to forests (EEA, 2014).
718 Sulphur can also stimulate microbial processes that make mercury bioavailable, leading to
719 bioaccumulation in the food chain (Greaver *et al.*, 2012). In agricultural soils in Europe,
720 however, fertilizer inputs of sulphur have increased to combat crop sulphur deficiencies as a
721 result of sulphur emission controls (Bender & Weigel, 2011).

722

723 Acidification is reversible, evident by increases in soil pH following decreased sulphur
724 emissions, although the recovery time varies; some areas with organic soils where deposition
725 has declined are showing either slow or no recovery (Greaver *et al.*, 2012; Lawrence *et al.*,
726 2012; RoTAP, 2012). On agricultural soils, lime can be applied to increase soil pH.
727 However, 50-80% of sulphur deposition on land is on natural, non-agricultural land
728 (Dentener *et al.*, 2006). Application of lime to naturally acidic forest soils can cause further
729 acidification of deep soil layers whilst increasing decomposition in surface litter, with no
730 improvement in tree growth (Lundström *et al.*, 2003).

731

732 Wider effects of acidification are starting to be understood through long-term monitoring.
733 Decreased organic matter decomposition due to acidification has increased soil carbon
734 storage in tropical forests (Lu *et al.*, 2014). However, in temperate forest soils acidification
735 can lead to reduced C:N ratios in soil, which in turn increases nitrification (Aber *et al.*, 2003),
736 but on already acidic soils reduces nitrification. In wetland soils, methane (CH₄) emissions
737 have also been suppressed by sulphur deposition (Gauci *et al.*, 2004). Conversely, declining
738 sulphur deposition has been associated with increased dissolved organic carbon fluxes from
739 organic soils (Monteith *et al.*, 2007), and decreased soil carbon stocks in temperate forest
740 soils (Oulehle *et al.*, 2011; Lawrence *et al.*, 2012).

741

742 **5.2 Nitrogen deposition**

743

744 Nitrogen deposition covers a wider geographical area than sulphur, as the sources are more
745 varied, including extensive agriculture fertilizer application, ammonia derived from livestock
746 operations, and biomass burning in addition to fossil fuel combustion (Figure 3b). Regions
747 with deposition in excess of 20 kg N ha⁻¹ yr⁻¹ in 2001 were western Europe, South Asia
748 (Pakistan, India, Bangladesh) and eastern China (Vet *et al.*, 2014); although extensive areas
749 with 4 kg N ha⁻¹ yr⁻¹ were found across North, Central and South America, Europe and Sub-
750 Saharan Africa. By contrast, 'natural' deposition in un-impacted areas is around 0.5 kg N ha⁻¹
751 yr⁻¹ (Dentener *et al.*, 2006). While emissions related to fossil fuel combustion have declined
752 along with sulphur across Europe, agricultural sources of nitrogen are likely to stay constant
753 in the near future across Europe (EEA, 2014), whilst overall global emissions are likely to
754 increase (Galloway *et al.*, 2008). Nitrogen deposition in China's industrialized and
755 intensively managed agricultural areas in the 2000s was similar to peaks in Western Europe
756 during the 1980s before mitigation (Liu *et al.*, 2013).

757

758 Deposition of nitrogen induces a ‘cascade’ of environmental problems, including both
759 acidification and eutrophication that can have both positive and negative effects on ecosystem
760 services (Galloway *et al.*, 2003). Excluding agricultural areas where nitrogen is beneficial,
761 11% of land surface received nitrogen deposition above $10 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ (Dentener *et al.*,
762 2006; Bouwman *et al.* 2002; Figure 3d). In Europe, eutrophication has and will continue to
763 impact a larger area than acidification (EEA, 2014).

764

765 Nitrogen fertilisation can increase tree growth (Magnani *et al.*, 2007) and cause changes in
766 plant species and diversity (Bobbink *et al.*, 2010), which in turn will alter the amount and
767 quality of litter inputs in to soils, notably the C:N ratio and soil-root interactions (RoTAP,
768 2012). However, increased carbon sequestration (Reay *et al.*, 2008) may be offset by
769 increased emissions of the greenhouse gases N_2O and CH_4 (Liu & Greaver, 2009). Long-term
770 changes caused by nitrogen deposition are uncertain as transport times vary between
771 environmental systems; and the only way to remove excess nitrogen is to convert it to an
772 unreactive gas (Galloway *et al.*, 2008).

773

774 [Figure 3 here]

775

776 **5.3 Heavy metal deposition**

777

778 Heavy metal emissions are associated with coal combustion and heavy industry. In the UK,
779 deposition is responsible for 25-85% of inputs to UK soils (Nicholson *et al.*, 2003). In
780 Europe, the areas at risk from cadmium, mercury and lead deposition in 2000 were 0.34%,
781 77% and 42% respectively, although emissions are declining (Hettelingh *et al.*, 2006).
782 Tighter legislation to control industrial emissions of heavy metals are helping to reduce the
783 environmental load of heavy metals in many regions, though rapid industrial growth in some
784 regions such as East Asia is increasing pressures on soil from heavy metal deposition. Global
785 heavy metal emissions and deposition are poorly understood in comparison to sulphur and
786 nitrogen; although the on-site impact of heavy metal contamination has been well studied
787 (Guo *et al.*, 2014). Metals in bioavailable form have toxic effects on soil organisms and
788 plants, influencing the quality and quantity of plant inputs to soils, rate of decomposition and,
789 importantly, can bio-accumulate in the food chain. Some heavy metals will persist for
790 centuries as they are strongly bound to soil organic matter (RoTAP, 2012), although they can

791 be mobilised to bioavailable form following drought-induced acidification, drainage and soil
792 erosion (Tipping *et al.*, 2003; Rothwell *et al.*, 2005).

793

794 Whilst the direct impacts of sulphur, nitrogen and heavy metals on inorganic soil chemical
795 processes are generally well understood, many uncertainties still exist about pollutant impacts
796 on biogeochemical cycling, particularly interactions between organic matter, plants and
797 organisms in natural and semi-natural systems (Greaver *et al.*, 2012). Process understanding
798 is dominated by research in Europe and North America (e.g. Bobbink *et al.*, 2010). Research
799 is needed across Asia, Africa and South and Central America where soil properties and
800 environmental conditions differ. Models need to be developed to examine the combined
801 effects of air pollutants and their interactions with climate change impacts and feedbacks on
802 greenhouse gas balances and carbon storage (Spranger *et al.*, 2008; RoTAP, 2012). Air
803 quality, biodiversity and climate change policies all impact on soils. A more holistic approach
804 to protecting the environment is needed, particularly as some climate change policies (e.g.
805 biomass burning, carbon capture and storage) have potential to impact air quality and,
806 therefore, soil quality (Reis *et al.*, 2012; RoTAP, 2012; Aherne & Posch, 2013).

807

808 Indirect impacts on soils can be addressed largely by preventing the pollution at source, or by
809 mitigating the adverse effects where these have already occurred. Air pollution control on
810 coal burning and increased car and fleet efficiency standards has been effective in reducing
811 sulphur deposition in many areas of the world, particularly in Europe since the 1970s.
812 Substitution of coal with bioenergy might also reduce sulphur emissions, but unless burned
813 cleanly in a controlled way, can also release pollutants to the air. In terms of nitrogen,
814 ammonia abatement techniques when fertilizers are spread (e.g. slurry injection) are helping
815 to reduce N deposition (Sutton *et al.*, 2007).

816

817 **6. Existing policies and practices that alleviate global change pressures on soils** 818 **from land use and management**

819

820 The previous text has highlighted specific anthropogenic activities that exert or alleviate
821 pressures on soils. Actions that alleviate pressures on soils driven by land use change and
822 land management can be broadly divided into three categories, those that:

823

- 824 1) Prevent conversion of natural ecosystems to other uses (e.g. protected areas, reduced
825 deforestation, prevention of wetland drainage, intensification rather than extensification);
826 2) Prevent soil degradation (erosion control, fire management, reduced tillage / conservation
827 agriculture, long term fallows, flood protection, use of organic amendments, intercropping,
828 improved rotations); and
829 3) Result in soil / ecosystem restoration (e.g. peatland rewetting, afforestation, re-vegetation
830 on degraded lands, improved grass varieties, appropriate animal stocking densities,
831 bioremediation).

832

833 Policies to encourage such actions were recently reviewed by Bustamante *et al.* (2014) and
834 include:

835

- 836 a) Economic incentives, e.g., special credit lines for low carbon agriculture and forestry
837 practices and projects, payment for ecosystem services (such as carbon storage) and tradable
838 credits such as carbon,
839 b) Regulatory approaches, e.g. enforcement of environmental law to protect natural areas, set-
840 aside policies,
841 c) Research, development and diffusion investments, e.g. increase of resource use-efficiency,
842 livestock improvement,
843 d) Information and certification schemes, e.g. in China, forest certification to promote
844 sustainable forest management, state regulation for protecting mandatory arable lands,
845 protection projects on Tibetan grasslands, a national wetland protection programme, and the
846 “grain for green” programme.

847

848 Many of these actions and policies are not directed at soil conservation, but nevertheless have
849 an effect on soil quality. Two of the main pieces of international policy that have served to
850 reduce pressures on soils, directly and indirectly, are the United Nations Convention to
851 Combat Desertification (CCD) and the United Nations Framework Convention on Climate
852 Change (UNFCCC). In general, policies and actions are important at all scales from
853 international conventions to local action, and local activity is encouraged by policies at
854 regional, national and global level. Policies to sustainably increase land productivity, for
855 example, can prevent land use change, and there are various other supporting actions that can
856 help deliver improvements, e.g. agricultural research, technology transfer, knowledge transfer

857 and improved rural infrastructure. Some examples of policies that impact on land
858 management and soil quality are given below.

859

860 ***6.1 United Nations Framework Convention on Climate Change (UNFCCC) and other***
861 ***climate specific policies***

862

863 Soil carbon storage and nutrient cycling as climate services are being increasingly recognised
864 e.g. under UNFCCC as part of national reporting and accounting, as part of life-cycle
865 greenhouse gas assessments for biofuels, in various regional initiatives and national efforts.
866 The UNFCCC is an international treaty, which came into force in 1994, setting an overall
867 framework for intergovernmental efforts to tackle the challenge posed by climate change.
868 The requirements for the 196 country Signatories (or 'Parties') to the UNFCCC include
869 adopting national mitigation policies and publishing national inventories of anthropogenic
870 emissions and sinks of greenhouse gases including activities on the land such as afforestation,
871 deforestation, agricultural management and wetland drainage and rewetting. Developed
872 country signatories have legally binding targets under the Kyoto Protocol and can count land
873 based emissions or sinks towards meeting these targets, thus incentivising activities that
874 protect soil carbon. Developing countries currently have voluntary targets and several
875 countries have made pledges that include reduced deforestation (e.g. Brazil and Indonesia) or
876 afforestation (e.g. 400000 km² in China). Under the Clean Development Mechanism (CDM)
877 developed countries can fund projects in developing countries that generate certified emission
878 reduction credits (CERCs). China, for example, has the largest number of CERCs in the
879 world (IFPRI, 2011). Brazil also has 180 CDM projects, the third largest number of CERCs
880 after China and India (Cole & Liverman, 2011). A number of projects in Africa, North
881 America and South Asia have a significant component for soil greenhouse gas emission
882 reduction of soil carbon sequestration, financed through the Verified Carbon Standard or the
883 American Carbon Registry.

884

885 As part of negotiations leading to the new climate treaty in Paris in December 2015, all
886 parties will be required to submit INDCs (Intended Nationally determined Contributions).
887 The new treaty will also include provision for REDD+ (reduced Emissions from
888 Deforestation and Degradation, including management of forests and enhancement of forest
889 carbon stocks). This could go some way to protecting forest soils, and negotiations have
890 been intense around methods for monitoring reporting and verification, with key issues such

891 as permanence (the risk the forest may be lost at a later date due to management or
892 environmental change) and leakage (displacement of land use change to other areas), and
893 how to finance such activities.

894

895 ***6.2 United Nations Convention to Combat Desertification (CCD)***

896

897 The CCD entered into force in December 1996; today 179 countries acknowledge it as a
898 legally binding framework to tackle land degradation and promote sustainable development
899 in fragile ecosystems. The Global Mechanism was established under the convention to
900 "promote actions leading to the mobilization and channelling of substantial financial
901 resources, including for the transfer of technology, on a grant basis, and/or on concessional or
902 other terms, to affected developing country Parties". In September 2011 the United Nations
903 General Assembly declared a goal of building a world with no land degradation. In October
904 2011 parties to the CCD issued a declaration calling for zero land degradation and for
905 adopting sustainable land management as a way to achieve sustainable development.

906

907 ***6.3 Millennium Development Goals (MDGs)***

908

909 Of the eight MDGs (UNDP, 2014a), soil protection is most relevant to the goal to ensure
910 environmental sustainability, since soils are critical in underpinning environmental
911 sustainability (Smith *et al.*, 2015). A complementary MDG, to develop a global partnership
912 for development, will improve the governance structure to deliver soil security. The other
913 MDG to which soils plays a critical contribution is the goal to eradicate extreme poverty and
914 hunger, with the role of soils in supporting food provision critical for the latter part of this
915 MDG (Smith *et al.*, 2015). The MDGs are currently being revisited to set a post-2015
916 development agenda (UNDP, 2014b), with discussion around the themes of localising the
917 post-2015 development agenda, helping to strengthen capacities and build effective
918 institutions, participatory monitoring for accountability, partnerships with civil society,
919 engaging with the private sector, and culture and development. The key emerging principles
920 from the dialogue are participation, inclusion, and the need for strengthened capacities and
921 partnerships (UNDP, 2014b). It is important that soils play their role in delivering this post-
922 2015 agenda.

923

924 ***6.4 Protected areas and the Convention on Biological Diversity (CBD)***

925

926 Many measures to protect biodiversity and vulnerable habitats also protect the soils
927 underpinning them, so numerous conservation actions around the world serve to protect soils,
928 even if this was not the primary aim (Smith *et al.*, 2013). Between 1990 and 2010, the
929 amount of forest land designated primarily for the conservation of biological diversity
930 increased by 35 percent, indicating a political commitment to conserve forests. These forests
931 now account for 12 percent of the world's forests (FAO, 2010). In India, a Supreme Court
932 ruling in 2011 on effective self-governance of "common" or communal land by local
933 communities may help to protect these valuable resources, and thereby the soils that underpin
934 them. Soil biodiversity is known to be important for soil function (Bodelier, 2011), yet it
935 rarely receives the attention enjoyed by larger flora and fauna within the ecosystem.

936

937 ***6.5 Reduced deforestation and forest management***

938

939 Various actions have been implemented to reduce deforestation (Bustamante *et al.*, 2014),
940 and to reduce the impact of forestry activities, such as reduced impact logging. UNFCCC,
941 carbon markets and other international environmental programs have contributed to global
942 efforts to reduce deforestation in addition to other sustainable natural resource management
943 programs in countries and by industry. For example, zero deforestation commitments made
944 by several companies (many made in 2014), and activities from bodies such as the
945 Roundtable for Sustainable Palm Oil (RSPO) and the Forest Stewardship Council (FCO)
946 certification scheme. Land improvement has increased in East Asia between 1981 and 2006
947 despite population increase, attributed largely to policies promoting tree planting and forest
948 plantation programs in China and Korea. In Brazil, deforestation was rapidly reduced after
949 national laws and regulations were enacted to protect forests in the 1990s and early 2000s
950 (including the soy moratorium and the forest code), followed up by state and municipal
951 governments setting further by-laws enforcing the deforestation moratorium (Bustamante *et*
952 *al.*, 2014).

953

954 ***6.6 Agricultural policies and practices***

955

956 The pressures on soils imposed by land use intensity change can be mitigated by regulation of
957 over-grazing and reduction of over-stocking on grazed grasslands, return of crop residues to
958 the soil, reduced tillage, best management practices, targeted nutrient management and

959 precision farming on croplands, and wetland / floodplain restoration. These actions have been
960 encouraged by various policies. Some examples include: The EU set-aside programme of the
961 1990s encouraged less intensive use of agricultural land where production is low and
962 environmental impacts are high. The EU Common Agricultural Policy ties agricultural
963 subsidies to implementation of best management practices and environmental protection, for
964 example through pillar 2 (rural development programmes) providing crop insurance for lower
965 fertilizer application rates; in Africa, policies for integrated land management to help protect
966 vulnerable soils; China's conservation tillage program (2012-2030); the USA Conservation
967 Reserve Program (set aside marginal lands, steep slopes).

968

969 **7. Conclusion: Keeping soils central to the science and policy agendas**

970

971 The International Year of Soils in 2015 is an excellent opportunity to raise the profile of soils
972 in the minds of national and international policy makers, land managers, timber and agro-
973 industries, and the public. Ensuring that vulnerable and high environmental value soils (e.g.
974 peatlands) are considered when making policies and decisions about which habitats and
975 ecosystems to convert or to protect, will help to reduce the pressure on soils particularly
976 vulnerable to global change drivers such as land use and land management, and maintain
977 important ecosystem services. This is in part happening with agendas around valuation of
978 ecosystem services and life-cycle assessments of impacts of land use change that include soil
979 carbon. At a time when governments are negotiating a legally binding climate change treaty
980 and making national targets for greenhouse gas reduction, and revisiting the Millennium
981 Development Goals, keeping soil carbon and nitrogen central to land based greenhouse gas
982 monitoring and reporting will maintain awareness with policy makers and industries with
983 emissions reduction targets. Both science and policy agendas are increasingly concerned with
984 long-term food security, ensuring that soils are central to considerations of how to achieve
985 on-going increases in production will enable those increases to be more sustainable into the
986 future.

987

988 Research and policy regarding soil quality and sustainability is abundant, but patchy and
989 disjointed. To ensure that soils are protected as part of on-going wider environmental and
990 sustainable production efforts, soils cannot, and should not, be considered in isolation of the
991 ecosystems that they underpin, but the role of soils in supporting ecosystems and natural
992 capital needs greater recognition (Robinson *et al.*, 2013, 2014). This can, in part, be enhanced

993 through education and awareness-raising which has started during the International Year of
994 the Soils in 2015. The time is ripe to consider a global soil resilience programme, under the
995 auspices of a global body such as the UN or one of its delivery agencies such as the FAO to
996 monitor, recover or sustain soil fertility and function, and to enhance the ecosystem services
997 provided by soils. The lasting legacy of the International Year of Soils in 2015 should be to
998 bring together robust scientific knowledge on the role of soils, and to put soils at the centre of
999 policy supporting environmental protection and sustainable development.

1000

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1002

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1009

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Tables

Table 1. Observed and modelled soil carbon change (%) when converting from land cover classes in the left hand column to land cover classes listed across the top. Results are from meta-analysis of observations from the sources listed below. Model results (range across three models) are shown for comparison in square brackets, range across the ISAM, LPJml, and LPJ_GUESS models (see text), although note this calculated as difference in soil carbon under the different land classes in 2010 and is thus not modelled loss/gain after a conversion. Negative numbers represent loss of soil carbon.

| | | Regrowth Forest | Tree plantation | Grassland | Pasture | Cropland |
|-----------|-------------------------------|----------------------|----------------------------------|-----------|---------------------------------|--|
| Forest | Global Trop. | -9% (2) | -13% (3) ^a | | +8% (3) -12% (2) | -42% (3) -41% (1) -25% (2) ^b -30% (2) ^c -24% (5) |
| | Temp. | | | | [-40 to -63%] | [-51 to -62%] -52% (1) -36% (4) |
| | Boreal | | | | [-52% to +17%] [-14 to -49%] | [-24 to -60%] -31% (1) [-63 to -65%] |
| Grassland | Global Trop | | | | [-1 to +15%] | [-2 to -6%] -32% (4) |
| | Temp Boreal | | | | [-28 to +3%] [-26 to -71%] | [-15 to -53%] [-70 to -79%] |
| Pasture | Global Trop Temp Boreal | | -10% (3) | | | -59% (3) [-19 to +0.5%] [-17 to -35%] [-28 to -59%] |
| Cropland | Global Trop Temp | +53% (3) +16% (4) | +18% (3) +29% (2) +20% (6) | +28% (4) | +19% (3) +26% (2) | |
| | Boreal | | | | | |

Footnotes: ^a Broadleaf tree plantations onto prior native forest or pasture did not affect soil C stocks whereas pine plantations reduced soil C stocks by -12 to -15%; ^b Annual crops; ^c Perennial crops; 1 Wei *et al.* (2014a); 2 Don *et al.* (2011); 3 Guo & Gifford (2002; tropical and temperate zones compiled); 4 Poepflau *et al.* (2011); 5 Murty *et al.* (2014); 6 Barcena *et al.* (2014).

Table 2. Soil carbon loss due to land use change 1860 to 2010 (PgCO₂)

| Model | Tropical | Temperate | Boreal | Global |
|------------------|-----------------|------------------|---------------|---------------|
| LPJ-GUESS | 46 | 55 | 1 | 109 |
| LPJmL | 128 | 95 | 0 | 227 |
| ISAM | 63 | 139 | 19 | 221 |
| Mean | 79 | 96 | 7 | 186 |

Table 3. Threats to soil resource quality and functioning under increasing intensity of agricultural management

| Agricultural management practice | Specific issue | Distribution | Major environmental consequence | Knowledge gap |
|---|------------------------|--|---|---------------------------------------|
| Cropping practice | Harvest frequency | Global | Soil quality and resilience | Impact on total C and nutrient cycles |
| | Monoculture | Global but particularly in developing and transition countries | Soil health, pesticide residue in intensively managed monocultures | Biological resilience |
| Use of agrochemicals | Over fertilization | Particularly in some developing countries | Soil acidification, water pollution, N ₂ O emission and nitrate accumulation | Rate reducing versus balancing |
| Irrigation | Submerged Rice | Developing countries, Asian | Water scarcity, methane emission | Trade-offs C and water, |
| | Arid/semi-arid regions | Arid/semi-arid regions | Secondary salinization, water scarcity | Competition use of water |
| Livestock management | Over-grazing | Largely in developing countries | Soil degradation, water storage, C loss | Forage versus feed crops? |
| | Industrial breeding | Largely in industrialized and transition countries | Waste pressure, water pollution, residue of veterinary medicine and antibiotics | Safe waste treatment and recycling |
| Agriculture in wetlands | Wetland drainage | Developing and transition countries | C loss | Agro-benefit versus natural value |

1 **Figure Legends**

2

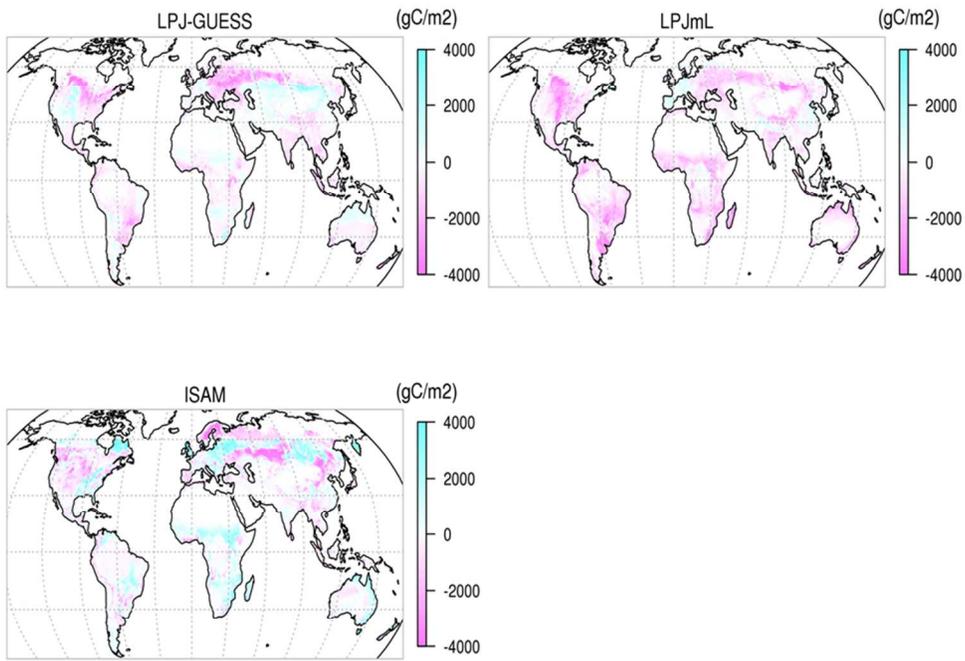
3 **Figure 1.** Maps of change in soil carbon due to land use change and land management
4 from 1860 to 2010 from three vegetation models. Pink indicates loss of soil carbon, blue
5 indicates carbon gain.

6

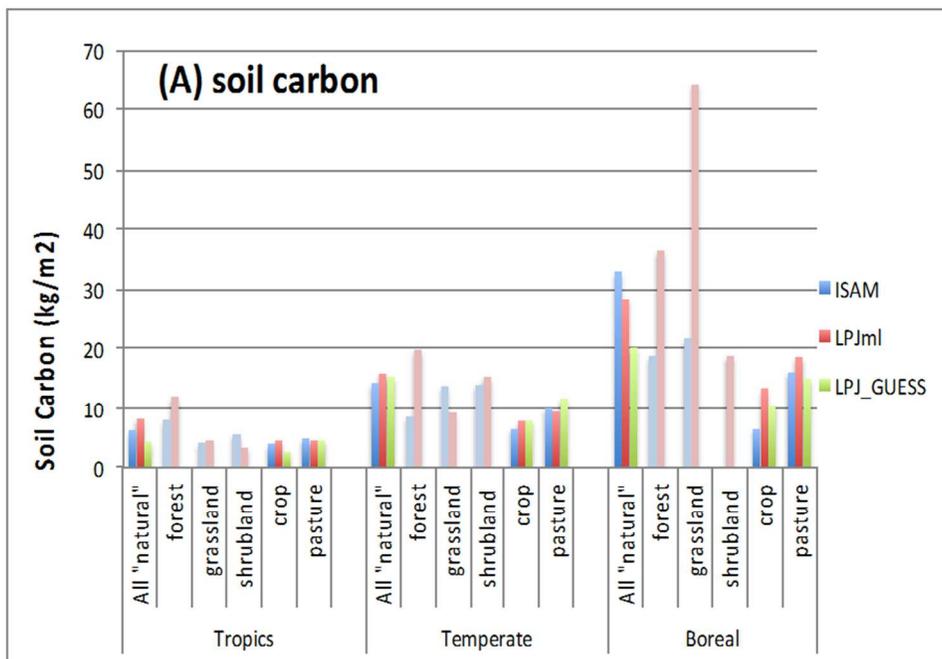
7 **Figure 2.** Soil carbon and nitrogen under different land cover types in three different
8 vegetation models (values are the annual average over the period 2001 to 2010).

9

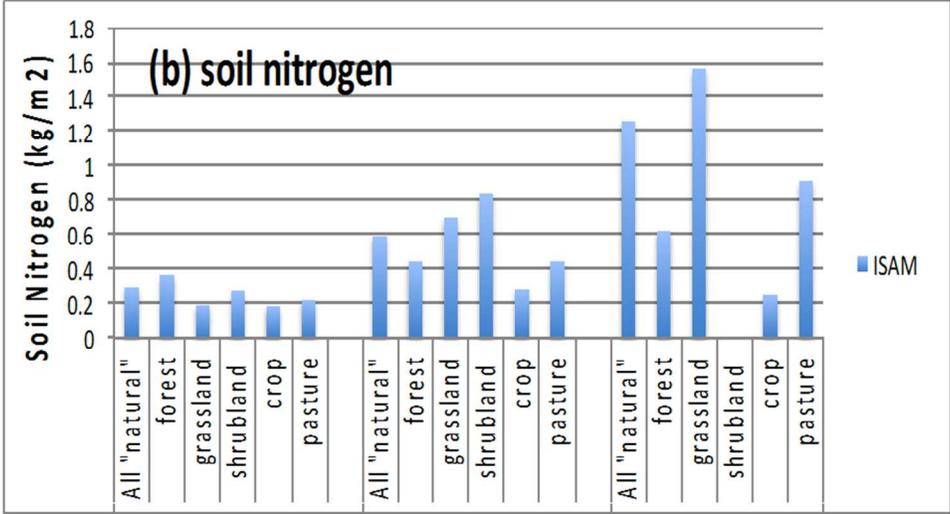
10 **Figure 3.** Uneven global distribution of soils sensitive to pollution by (a) acidification and (b)
11 eutrophication (measured by soil C:N) compared to uneven distribution of atmospheric (c)
12 sulphur and (d) nitrogen pollution. Soils most sensitive to acidification have low base
13 saturation and cation exchange capacity, as defined by (Kuylenstierna *et al.*, 2001).
14 Acidification is caused by both sulphur and nitrogen. Eutrophication is caused by nitrogen.
15 Soil data in (a) and (b) were produced using the ISRIC-WISE derived soil properties (ver 1.2)
16 (Batjes, 2012) and the FAO Digital Soil Map of the World. Atmospheric deposition data in
17 (c) and (d) were provided by the World Data Centre for Precipitation Chemistry
18 (<http://wdcpc.org>, 2014) and are also available in Vet *et al.* (2014). Data show the ensemble-
19 mean values from the 21 global chemical transport models used by the Task Force on
20 Hemispheric Transport of Air Pollution (HTAP) (Dentener *et al.*, 2006). Total wet and dry
21 deposition values are presented for sulphur, oxidized and reduced nitrogen.



254x190mm (96 x 96 DPI)

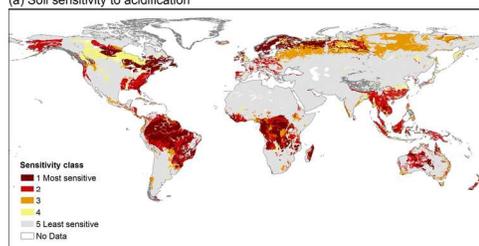


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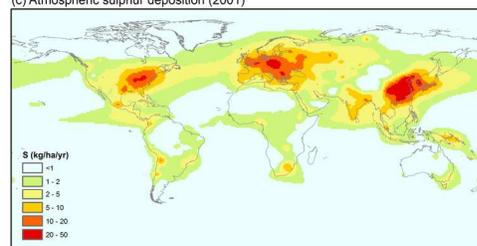


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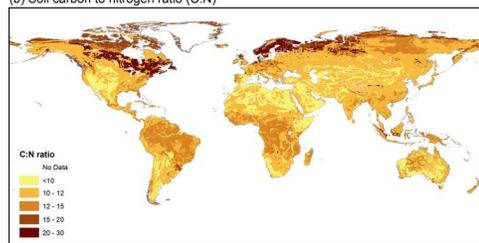
(a) Soil sensitivity to acidification



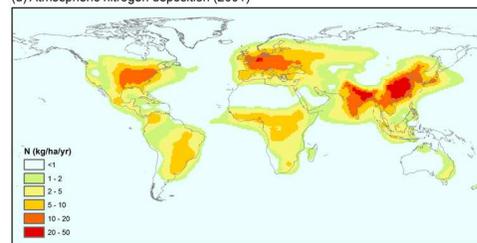
(c) Atmospheric sulphur deposition (2001)



(b) Soil carbon to nitrogen ratio (C:N)



(d) Atmospheric nitrogen deposition (2001)



168x101mm (300 x 300 DPI)