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The role of the cloud radiative effect in the sensitivity of the Intertropical Convergence Zone to convective mixing.

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ABSTRACT

Studies have shown that the location and structure of the simulated Intertropical Convergence Zone (ITCZ) is sensitive to the treatment of sub-gridscale convection and cloud-radiation interactions. This sensitivity remains in idealised aquaplanet experiments with fixed surface temperatures. However, studies have not considered the role of cloud-radiative effects (CRE, atmospheric heating due to cloud-radiation interactions) in the sensitivity of the ITCZ to the treatment of convection. We use an atmospheric energy input (AEI) framework to explore how the CRE modulates the sensitivity of the ITCZ to convective mixing in aquaplanet simulations. Simulations show a sensitivity of the ITCZ to convective mixing, with stronger convective mixing favoring a single ITCZ. For simulations with a single ITCZ, the CRE maintains the positive, equatorial AEI. To explore the role of the CRE further, we prescribe the CRE as either zero or a meridionally and diurnally varying climatology. Removing the CRE is associated with a reduced equatorial AEI and an increase in the range of convective mixing rates that produce a double ITCZ. Prescribing the CRE reduces the sensitivity of the ITCZ to convective mixing by 50%. In prescribed-CRE simulations, other AEI components, in particular the surface latent heat flux, modulate the sensitivity of the AEI to convective mixing. Analysis of the meridional moist static energy transport shows that a shallower Hadley circulation can produce an equatorward energy transport at low latitudes even with equatorial ascent.

1. Introduction

Tropical rainfall is often associated with a discontinuous zonal precipitation band commonly known as the Intertropical Convergence Zone (ITCZ). The ITCZ migrates between the Northern and Southern Hemispheres with the seasonal cycle, with a zonal-, time-mean position of approximately 6°N (Schneider et al. 2014). The ITCZ is co-located with the ascending branch of the Hadley circulation, where strong moist convection leads to high rainfall. The upper branches of the Hadley circulation typically transport energy poleward, away from the ITCZ. Recent studies have associated characteristics of the ITCZ with the energy transport by the Hadley circulation (Frierson and Hwang 2012; Donohoe et al. 2013; Adam et al. 2016; Bischoff and Schneider 2016).

A double ITCZ bias is prominent in current and previous generations of coupled general circulation models (GCMs; Li and Xie 2014; Oueslati and Bellon 2015). The ITCZ is too intense in the Southern Hemisphere (Lin 2007), resulting in two annual-, zonal-mean tropical precipitation maxima, one in each hemisphere. A bias remains in atmosphere-only simulations with prescribed sea surface temperatures (SSTs) (Li and Xie 2014). Aquaplanet simulations provide an idealised modelling environment in which some complex boundary conditions in tropical circulation such as land/sea contrasts and orography are removed. However aquaplanet configurations of GCMs coupled to a slab ocean produce a broad range of tropical precipitation mean states (Voigt et al. 2016); even prescribing zonally uniform SSTs does not resolve the inter-model variability (Blackburn et al. 2013).

a. Modelling studies

Characteristics of the simulated ITCZ are sensitive to the representation of cloud-radiation interactions (Fermepin and Bony 2014; Li et al. 2015; Harrop and Hartmann 2016). In the deep tropics the cloud radiative effect (CRE) warms the atmosphere (Allan 2011), with important effects on tropical circulation (Slingo and Slingo 1988; Crueger and Stevens 2015). The CRE is associated with a more prominent single ITCZ (Crueger and Stevens 2015; Harrop and Hartmann 2016; Popp and Silvers 2017). Both Harrop and Hartmann (2016) and Popp and Silvers (2017) investigated the association between the Hadley circulation and CRE in a range of aquaplanet simulations with and without the CRE. In all GCMs used, the CRE is associated with increased equatorial rainfall, an equatorward contraction of the ITCZ, and a strengthening of the mean meridional circulation. The authors emphasise different mechanisms by which the CRE promotes a single ITCZ. Harrop and Hartmann (2016) propose that the CRE warms the upper tropical troposphere, which reduces the convective available potential energy and restricts deep convection to the region of warmest SSTs, whilst Popp and Silvers (2017) argue that the CRE strengthens the Hadley circulation and moves the ITCZ equatorward, associated with increased moist static energy (MSE) advection by the lower branches of the Hadley circulation. The strengthening of the mean circulation is associated with the CRE meridional gradient, as the CRE is positive in the tropics and negative in the extra-tropics ($\geqslant \pm 45°$ latitude; Allan 2011). However, it should be noted that the CRE reduces total tropical-mean ($\leqslant \pm 30°$ latitude) precipitation due to reduced radiative cooling (Harrop and Hartmann 2016).

Across a hierarchy of models it has been shown that the simulation of tropical precipitation is sensitive to the representation of convection (Terray 1998; Frierson 2007; Wang et al. 2007; Chikira 2010; Mobis and Stevens 2012; Oueslati and Bellon 2013; Bush et al. 2015; Nolan et al.

2016). For example, variations in lateral entrainment and detrainment rates, which alter the representation of deep convection, affect the diurnal cycle of precipitation over the Maritime Continent (Wang et al. 2007) and South Asian monsoon precipitation rates (Bush et al. 2015). Increasing convective mixing strengthens deep convection in convergence zones, associated with an increased moisture flux from subsidence regions (Terray 1998; Oueslati and Bellon 2013).

In full GCMs, complex surface characteristics and boundary conditions including land-sea contrasts, orography and SST gradients, make it challenging to understand the sensitivity of tropical precipitation to the representation of convection (Oueslati and Bellon 2013; Bush et al. 2015). Even in the absence of complex surface topography, aquaplanet studies have also shown that characteristics of tropical precipitation, in particular the location and intensity of the ITCZ, are sensitive to the sub-gridscale treatment of convection (Hess et al. 1993; Numaguti 1995; Chao and Chen 2004; Liu et al. 2010; Mobis and Stevens 2012). Mobis and Stevens (2012) studied the sensitivity of the ITCZ location to the choice of convective parameterisation scheme in an aquaplanet configuration of the ECHAM GCM by comparing the Nordeng (1994) and Tiedtke (1989) schemes, which vary in their formulations of entrainment, detrainment and cloud base mass flux for deep convection. The Nordeng scheme, with a higher lateral entrainment rate, produced a single ITCZ, whilst the Tiedtke scheme produced a double ITCZ. The authors associate the location of maximum boundary layer MSE with the ITCZ location; they argue that mechanisms that control the boundary layer MSE are important to the sensitivity of the ITCZ to the representation of convection. The boundary layer MSE distribution is predominantly controlled by the surface winds, which are influenced by convective heating, allowing variations in convective heating to influence the ITCZ structure. The importance of the surface winds is further emphasised by simulations with prescribed surface winds in the computation of the surface fluxes (Mobis and Stevens 2012). These simulations lead to the conclusion that there is a strong association between surface turbulent fluxes and the ITCZ.

While the ITCZ has been shown to be sensitive to the CRE and the convective parameterisation scheme, no study has separated these effects. This paper will analyse the sensitivity of the ITCZ to convective mixing in aquaplanet simulations using the Met Office Unified Model (MetUM), and the role of the CRE in this sensitivity.

b. Atmospheric Energy framework

Literature based on a hierarchy of models, as well as reanalysis data and observations, concludes that the northward displacement of the ITCZ from the equator is anti-correlated with the northward cross-equatorial atmospheric energy transport (Kang et al. 2008; Frierson and Hwang 2012; Donohoe et al. 2013). Bischoff and Schneider (2014) developed a diagnostic framework to relate the location of the ITCZ to this energy transport.

The zonal-mean atmospheric MSE budget is (Neelin and Held 1987):

$$[AEI] = \partial_t[\hat{h_e}] + \partial_y[\widehat{vh}] \tag{1}$$

where AEI is the atmospheric energy input (AEI); vh is the meridional MSE flux, (v is meridional wind; h is MSE); h_e is the moist enthalpy; [] denotes a zonal- and time-mean; ˆ represents a mass weighted vertical integral; ∂_y is the meridional derivative; and ∂_t is the time derivative. Local Cartesian coordinates are printed with $y = a\phi$, (where a is Earth's radius and ϕ is latitude,) but all calculations are performed in spherical coordinates. Bischoff and Schneider (2014) assume a statistically steady state ($\partial_t[\hat{h_e}]$=0) and that $[\widehat{vh}]$ in the tropics is dominated by the zonal-mean circulation and therefore $[\widehat{vh}]$ equals zero at the ITCZ. Through performing a first-order Taylor expansion of the equatorial $[\widehat{vh}]$, Bischoff and Schneider (2014) derive the dependence of the ITCZ location on the equatorial MSE flux and equatorial AEI:

$$\delta \approx -\frac{1}{a}\frac{[\widehat{vh}]_0}{[AEI]_0} \tag{2}$$

with the AEI defined as:

$$[AEI] = [S] - [L] - [O] \tag{3}$$

where subscript 0 denotes the equatorial value, S is the net incoming shortwave radiation at the top of the atmosphere (TOA), L is the outgoing longwave radiation at the TOA, and O is the net downward flux at the surface. Bischoff and Schneider (2016) retain higher order terms in the Taylor expansion to derive a framework for negative $[AEI]_0$. A negative $[AEI]_0$ is associated with a double ITCZ as $[\hat{vh}]$ no longer increases with latitude; energy is transported equatorward at low latitudes to achieve equilibrium. A double ITCZ is associated with two off-equatorial energy flux equators, where the total meridional energy flux equals zero. Bischoff and Schneider (2016) derive an expression for the locations of a double ITCZ:

$$\delta \approx \pm\frac{1}{a}\left\{-\frac{6([AEI])_0}{\partial_{yy}([AEI])_0}\right\}^{\frac{1}{2}} + \frac{[\widehat{vh}]_0}{2a([AEI])_0} \tag{4}$$

Note equation 4 is from a corrigendum for the original paper.

Bischoff and Schneider (2014) explore the relationship derived in (2) using an idealised slab-ocean GCM with a prescribed oceanic heat transport. They investigate the effects of the $[AEI]_0$ and the $[\widehat{vh}]_0$ through varying the imposed equatorial ocean heat flux and the atmospheric longwave absorption. Changes in both $[AEI]_0$ and $[\widehat{vh}]_0$ affect the latitude of the ITCZ; this theoretical relationship is supported in observations and reanalyses (Adam et al. 2016). Bischoff and Schneider (2016) examine the double ITCZ framework (4) using a slab-ocean GCM and varying the tropical and extra-tropical components of the imposed ocean energy flux divergence. An increased tropical ocean energy flux divergence decreases the $[AEI]_0$. For double ITCZ scenarios and when $[\widehat{vh}]_0$ is negligible, decreasing the $[AEI]_0$ shifts the energy flux equator poleward. The diagnosed energy flux equators from (2) and (4) are close to the simulated precipitation maxima, highlighting the association between the AEI and ITCZ.

However, Bischoff and Schneider (2014)'s definition of the $[AEI]$ (3) is chosen as their simulations prescribe O, which allows only the TOA energy budget $(S-L)$ to vary. This constrains the AEI response to model perturbations, as surface radiation and turbulent fluxes are constrained at equilibrium, which could reduce the impact of surface-flux feedbacks on the ITCZ. We use atmosphere-only simulations with prescribed SSTs, allowing variations in the components of O. As our experiments do not have a closed surface energy balance and we are interested in cloudy-sky radiation AEI components, we choose to write the AEI as:

$$[AEI] = [SW] + [LW] + [H] \tag{5}$$

where SW and LW represent the net atmospheric heating from shortwave and longwave radiation, respectively, and H denotes the atmospheric heating from surface sensible and latent heat fluxes. Both fixed SST and prescribed O frameworks misrepresent the real climate system by restricting air-sea coupled feedbacks (discussed further in section 4). From an AEI perspective, Mobis and Stevens (2012) severely constrain H in a subset of experiments by prescribing the surface winds when computing the surface fluxes. This reduces the sensitivity of the ITCZ to the convective parameterisation scheme.

Previous research on the response of the simulated ITCZ to variations in the sub-gridscale representation of convection have not considered the role of the CRE or used an energy budget framework like that proposed by Bischoff and Schneider (2014). We hypothesise that the sensitivity of the ITCZ to these factors can be linked to variations in AEI and $[\widehat{vh}]$.

2. Methodology

We use variations of an N96 (1.25° latitude × 1.875° longitude) aquaplanet configuration of the Met Office Unified Model (MetUM) Global Atmosphere 6.0 (GA6.0) configuration (Walters et al. 2017). The deep convective parameterisation scheme is an altered form of the mass flux scheme in Gregory and Rowntree (1990), including a convective avaliable potential energy closure based on Fritsch and Chappell (1980) and a mixing detrainment rate dependent on the relative humidity (Derbyshire et al. 2004). Unless noted, all simulations are run for three years with a "Qobs" SST profile (Neale and Hoskins 2001), with the first sixty days discarded as spin-up.

a. Simulations performed

To explore the sensitivity of the simulated ITCZ to convective mixing, we perform five simulations varying the lateral entrainment (ε) and detrainment (d_m) rates for deep-level convection (Table 1). In GA6.0 these rates are:

$$\varepsilon = 4.5 f_{dp} \frac{p(z)\rho(z)g}{p_*^2} \tag{6}$$

$$d_m = 3.0(1-RH)\varepsilon \tag{7}$$

Both ε and d_m are given as a fractional mixing rate per unit length (m^{-1}). In (6) and (7), p and p_* are pressure and surface pressure (Pa); ρ is density ($kg\ m^{-3}$); g is gravitational acceleration ($m\ s^{-2}$); f_{dp} is a constant with the default value of 1.13; RH is relative humidity. We control ε and d_m by scaling f_{dp} to five values between 0.25 and 1.5 × the default value: 0.28 (F0.28), 0.57 (F0.57), 0.85 (F0.85), 1.13 (F1.13) or 1.70 (F1.70).

To explore the influence of the CRE on the sensitivity of the ITCZ to convective mixing we perform a companion set of experiments with cloud-radiation interactions removed: F0.28NC, F0.57NC, F0.85NC, F1.13NC and F1.70NC (Table 1). Cloud-radiation interactions are removed by setting the cloud liquid and cloud ice to zero in the radiation scheme.

Finally, a third set of simulations use a prescribed CRE (Table 2) to investigate the relative importance of f_{dp} and the CRE to characteristics of the ITCZ. The four simulations have a prescribed, diurnally varying CRE vertical profile computed from a single-year simulation with f_{dp} equal to 0.57 or 1.13 (PC0.57 and PC1.13, respectively). The CRE is prescribed using cloudy-sky upward and downward fluxes at each model level at every model timestep. The diurnally varying CRE profile is computed as a hemispherically symmetric and zonally uniform composite of the climatological diurnal cycle at each grid point, referenced to local solar time. Two of the four simulations prescribe a CRE at a different f_{dp} constant from that in the simulation (F1.13PC0.57, F0.57PC1.13), whilst the other two simulations use a CRE from the same f_{dp} value to assess the sensitivity to prescribing cloud-radiation interactions (F1.13PC1.13, F0.57PC0.57).

3. Results

a. Sensitivity of the ITCZ to the convective mixing.

Figure 1a shows the sensitivity of the ITCZ to f_{dp} with a single ITCZ at higher values (F1.13, F1.70). Reducing f_{dp} promotes a double ITCZ, with peak precipitation further away from the equator (F0.28, F0.57). F0.85 has a marginal double ITCZ with no substantial difference between equatorial and off-equatorial precipitation. Decreasing f_{dp} is associated with a weaker horizontal gradient of the mass meridional streamfunction (Figure 2). F0.28 is the only simulation to show a reversed Hadley circulation in the deep tropics (Figure 2e), associated with upper-level zonal-mean equatorial subsidence, typical of a double ITCZ. F0.57 meanwhile has a typical double ITCZ structure in precipitation but not in the mass meridional streamfunction (Figure 1a and 2d), which we refer to as a “split ITCZ”: two off-equatorial precipitation maxima and two ascending branches of the Hadley circulation, without any substantial zonal-mean subsidence equatorward of the precipitation maxima.

Convective mixing reduces the difference in MSE between a convective plume, determined by the boundary layer MSE, and the free-troposphere (Mobis and Stevens 2012), which reduces the buoyancy of the convective plume. Assuming the sensitivity of the environmental saturated MSE to f_{dp} is small, the depth of convection will depend on the boundary layer MSE and f_{dp}. Decreasing f_{dp} will deepen convection for a constant boundary layer MSE, and reduce the minimum boundary layer MSE at which deep convection occurs. Following weak-temperature gradient arguments (e.g. Sobel et al. 2001) and assuming a small meridional gradient in free-tropospheric tropical temperature, and hence a small gradient in the saturated MSE across the deep tropics, the reduced minimum boundary layer MSE needed for deep convection strengthens convection in off-equatorial tropical latitudes over cooler SSTs. Stronger off-equatorial deep convection decreases equatorward low-level winds in the deep tropics, reducing equatorial boundary layer MSE. Hence, decreasing f_{dp} is associated with a poleward ITCZ shift and promotes a double ITCZ. Similar arguments can be made for higher f_{dp} promoting a single ITCZ.

The sensitivity of the ITCZ to f_{dp} is associated with AEI changes (Figure 1b), with a change from a single (F1.13) to a double/split ITCZ (F0.28/F0.57) associated with a decrease in the $[AEI]_0$ (Figure 3d and e). Simulations with a single/double ITCZ in precipitation have a positive/negative $[AEI]_0$ (Figure 1b), in agreement with Bischoff and Schneider (2014). Changes in cloudy-sky radiation and latent heat flux are the dominant components of AEI changes (blue and orange lines, respectively, in Figure 3). In F1.13 the total CRE peaks at approximately 60 Wm^{-2} at the equator and reduces to zero around 15° latitude (blue line in Figure 3b). This equatorial warming comes almost entirely from the longwave CRE, which dominates the total CRE equatorward of 10° latitude (not shown). In the subtropics, 20° to 30° latitude, low clouds contribute to a negative CRE of ≈ 2 Wm^{-2}, as longwave cooling from boundary layer clouds is greater than the shortwave heating. Without the CRE contribution to the $[AEI]_0$ in F1.13, $[AEI]_0$ would be negative, suggesting that the CRE maintains the single ITCZ. Removing the CRE from the AEI in F1.13 would give an $[AEI]_0$ of -25.7 Wm^{-2}, assuming that no other AEI components change. Using Bischoff and Schneider (2016)'s framework, (4), with values for AEI once removing the CRE and assuming that $[\widehat{vh}]_0 \simeq 0$ Wm^{-1}, (associated with an hemispherically symmetric atmospheric circulation), predicts a double ITCZ at $\pm$ 5.6° latitude.

The split ITCZ in F0.57 is associated with a substantially reduced equatorial CRE and an increased off-equatorial CRE (Figure 3d). We chose CRE profiles from one year of F0.57 and F1.13 for our prescribed CRE simulations (Table 2), as these two simulations show CRE profiles typical of a double and single ITCZ, respectively; these simulations are analysed in section 3d. As the Hadley circulation and ITCZ are associated with the AEI, and the CRE plays a substantial role in AEI changes when varying f_{dp}, we hypothesize that prescribing the CRE will reduce or remove the sensitivity of the AEI and ITCZ to f_{dp}.

b. Sensitivity of the ITCZ to convective mixing with no cloud radiative effect

To test our hypothesis above, we first analyse simulations with the CRE removed (Table 1), similar to Harrop and Hartmann (2016). Figure 4a and Figure 5 show the zonal-mean precipitation and mass meridional streamfunction respectively in simulations with no CRE (Table 1). Removing the CRE at f_{dp} = 1.13 (F1.13NC) leads to a switch from a single to a split ITCZ, and a $\approx$ 20% weakening of the Hadley circulation (Figure 4a and 5b).

Similar to Harrop and Hartmann (2016), removing the CRE cools the tropical ($\leqslant 30°$ latitude) upper-troposphere, destabilizing the atmosphere and reducing the environmental saturated MSE. For a fixed boundary layer MSE and convective mixing rate, removing the CRE deepens convection as the buoyancy of a convective plume increases relative to the saturated MSE of the environment. Hence, removing the CRE reduces the minimum boundary layer MSE for deep convection, strengthening off-equatorial convection over cooler SSTs. Stronger off-equatorial convection decreases equatorward low-level winds in the deep tropics, reducing equatorial boundary layer MSE and promoting a double ITCZ. This mechanism is similar to that proposed for the sensitivity of the ITCZ to f_{dp} (section 3a). However, when removing the CRE changes in the environmental saturated MSE play the dominant role, whilst for the sensitivity of the ITCZ to f_{dp}, changes in the convective parcel MSE dominate.

The weaker Hadley circulation and double ITCZ in precipitation in F1.13NC is consistent with AEI changes. In F1.13NC removing CRE reduces the $[AEI]_0$ by ≈ 45 Wm^{-2}, leading to a negative $[AEI]_0$, and increases the subtropical AEI by up to 15 Wm^{-2} (20 to 45° latitude) (Figure 6f). Across the deep tropics the AEI change is not equal to the CRE diagnosed from F1.13, due to increased turbulent and clear-sky fluxes. These increased fluxes, associated with an equatorward shift of the ITCZ, partially offset the reduction in $[AEI]_0$. Hence, the predicted location of the double ITCZ in section 3a when removing the CRE overestimated the poleward shift of the ITCZ. Removing the CRE reduces tropical-domain ($\leqslant 30°$ latitude) AEI, which is associated with increased AEI at higher latitudes to maintain equilibrium. Our simulations are consistent with the suggested mechanisms proposed by Popp and Silvers (2017): the ITCZ is located at the maximum boundary layer MSE, and a weaker meridional circulation is associated with a reduced AEI gradient.

At all f_{dp} removing the CRE reduces the maximum precipitation rate, weakens the Hadley circulation (comparing Figure 1a and 4a), and moves the latitude of peak precipitation poleward

(Figure 7a). The sensitivity of the ITCZ structure to removing the CRE depends on the convective mixing rate: either a broader single ITCZ (F1.70NC), a poleward shift of a double/split ITCZ (F0.28NC and F0.57NC), or a switch from a single to a split/double ITCZ (F0.85NC and F1.13NC). Removing the CRE cools the upper troposphere and reduces the boundary layer MSE required for deep convection. This increases the f_{dp} value at which the ITCZ transitions from single to split/double.

Removing the CRE changes, but does not remove, the sensitivity of the ITCZ to f_{dp}. Quantifying the apparent effect of the CRE on the sensitivity of the ITCZ to f_{dp} is difficult, as the effect depends on both the range of f_{dp} considered and the metric used (Figure 7). When an off-equatorial ITCZ is simulated in CRE-off simulations ($0.28 \leqslant f_{dp} \leqslant 1.13$), including the CRE increases the sensitivity of the ITCZ location to f_{dp} by $\approx 30\%$ (comparing the slopes of the solid regression lines in Figure 7a). However, because F1.70NC has a single ITCZ, including the CRE cannot shift the ITCZ equatorward. Hence, when $0.28 \leqslant f_{dp} \leqslant 1.70$ the change in sensitivity reduces to nearly zero (comparing the slopes of the dashed lines). The reduction in sensitivity also depends on the chosen metric; for instance, the maximum precipitation rate has a negligible sensitivity to f_{dp} in CRE-off simulations but a substantial sensitivity in CRE-on simulations (Figure 7b), highlighting that the CRE has a positive feedback on convection as increasing f_{dp} is associated with an increased CRE (Figure 8).

Increasing f_{dp} is associated with an increased tropical-domain CRE (Figure 8), which is counter-intuitive as one might expect that increasing f_{dp} will lead to lower cloud tops and hence a reduced CRE. However, the maximum cloud top height at the ITCZ is insensitive to f_{dp} (not shown), but the minimum temperature where the cloud fraction goes to zero (cloud top temperature) is sensitive to f_{dp} in both CRE-on and CRE-off simulations (Figure 8). The cloud top temperature decreases as f_{dp} increases (Figure 8), associated with a cooler upper-troposphere.

Furthermore, the increase in SST at the ITCZ location, associated with equatorward contraction of the ITCZ, also contributes to an increased CRE at higher f_{dp}.

Removing the CRE decreases the sensitivity of the AEI to f_{dp} (comparing Figure 1b and Figure 4b). The reduced sensitivity of the AEI is associated with a reduced sensitivity of the ITCZ. Latent heat flux variations account for most of the remaining AEI sensitivity to f_{dp} (Figure 6). In simulations with a double ITCZ (F0.28NC, F0.57NC and F0.85NC), changes in the latent heat flux and AEI have a bi-modal structure, indicating reduced latent heat flux at the location of maximum precipitation in F1.13NC (Figure 6c-e). Changes in the latent heat flux are predominantly controlled by alterations in near-surface wind speed rather than changes in near-surface specific humidity (not shown).

Simulations so far agree with the association in Bischoff and Schneider (2016) between a negative $[AEI]_0$ and a double ITCZ. However, the negative $[AEI]_0$ in F0.57, F0.85NC and F1.13NC requires an equatorward transport of energy at low latitudes, but the mean mass meridional streamfunction suggests a poleward transport of energy (Figure 2b, 5c, 5d). In the following subsection we discuss mechanisms for an equatorward energy transport.

c. Mechanisms responsible for an equatorward energy transport

To better understand the response of the mean circulation, associated with ITCZ changes, to varying f_{dp} and removing the CRE, we partition the divergence of the MSE flux ($\partial_y[\widehat{vh}]$) into two components: the mean circulation ($\partial_y([\hat{v}][\hat{h}])$) and the eddy contribution ($\partial_y[\widehat{vh}] - \partial_y([\hat{v}][\hat{h}])$. In these simulations it has not been possible to close the atmospheric energy budget (1) due to local energy conservation issues (discussed further in section 4), however the sign of the $[AEI]_0$ is consistent with the sign of the $\partial_y[\widehat{vh}]$ in simulations so far. In all simulations the eddy contribution to the meridional MSE flux is substantial across the tropics highlighting that the mean atmospheric circulation is not solely responsible for transporting energy. Furthermore, one should not necessarily assume a correspondence between the required MSE transport and the transport by the mean meridional circulation. In simulations with a single/double ITCZ, both the mean circulation and eddies transport energy poleward/equatorward at low latitudes. In F0.57, which has a negative $[AEI]_0$ and a split ITCZ, equatorward transport of energy at low latitudes is achieved solely by eddies. When f_{dp} equals 0.85 and 1.13, a change in the sign of the energy transport by the mean circulation ($\partial_y([\hat{v}][\hat{h}])$) occurs at low latitudes when removing the CRE, however there is still equatorial ascent across most of the troposphere (Figure 5b, c). To understand the sensitivity of the mean circulation to removing the CRE at these convective mixing rates, we partition the change in the MSE flux ($[\hat{v}][\hat{h}]$) into mean circulation changes and MSE variations.

First, the meridional mass flux, denoted by V, in F1.13NC (V_e) is partitioned into two components:

$$
\begin{aligned}
V_e &= V_c(1+\alpha) + V_r \\
\text{where } \alpha &= \frac{V_e \cdot V_c}{V_c \cdot V_c} - 1
\end{aligned} \tag{8}
$$

Subscripts c and e represent the zonal-, time-mean value of the control and experiment simulation (in this case F1.13 and F1.13NC respectively). α is a globally uniform scaling term calculated using the dot product of the meridional mass fluxes in the tropics (30°N to 30°S). We account for variations in density in V. $V_c(1+\alpha)$ represents a change in strength of the control circulation; V_r represents a change in circulation structure. Next, the MSE, ($c_pT + gz + Lq$), in the experiment simulation (h_e) is written as:

$$h_e = h_c + h_p \tag{9}$$

where subscript p represents the zonal-, time-mean difference between the two simulations. The change in the MSE flux between the experiment and control simulation can therefore be written as:

$$V_e h_e - V_c h_c = \\ \alpha V_c h_c + V_r h_c + V_c h_p + (\alpha V_c + V_r) h_p \tag{10}$$

Each term in (10) represents a mechanism by which vh can vary: $\alpha V_c h_c$ represents circulation intensity changes; $V_r h_c$ represents changes in circulation structure; $V_c h_p$ represents MSE profile changes; and $(\alpha V_c + V_r) h_p$, represents MSE profile changes correlated with changes in circulation structure and strength.

Three out of the four mechanisms are important in reducing the poleward MSE transport by the Hadley circulation in F0.85NC and F1.13NC (Figure 9): a reduction in Hadley circulation strength (Figure 9e); a shallower mean circulation (Figure 9f); and a reduced MSE export at the top of the Hadley circulation due to lower MSE associated with upper-tropospheric cooling (Figure 9g). MSE profile changes correlated with changes in circulation strength and intensity [$(\alpha V_c + V_r) h_p$] are small compared to the other three mechanisms (Figure 9h). As changes in circulation strength ($\alpha V_c h_c$) cannot change the direction of energy transport, the reduced upper-tropospheric MSE ($V_c h_p$) and shallower Hadley circulation ($V_r h_c$) must be responsible for the change in energy transport direction by the mean circulation. At the equator, circulation strength changes ($\alpha V_c h_c$) contribute ≈16% of the reduced $\partial_y([\widehat{v}][\widehat{h}])$; reduced MSE export by the upper branch of the mean circulation ($V_c h_p$) and a shallower Hadley circulation ($V_r h_c$) contribute ≈34% and 50% respectively (not shown). Therefore, at certain convective mixing rates, in our case when f_{dp} = 0.85 and 1.13, removing the CRE is not associated with a substantial double ITCZ in the mass meridional streamfunction, even though MSE is transported equatorward at low latitudes and the $[AEI]_0$ is negative. Similar behaviour has also been concluded by Popp and Silvers (2017) who found that in certain simulations the zero mass meridional streamfunction remained at the equator even when the $[AEI]_0$ was negative.

Removing the CRE and varying f_{dp} are associated with substantial AEI changes which require MSE transport variations. In the two sets of simulations discussed so far, we identified three mechanisms to transport MSE equatorward at low latitudes; which mechanisms dominates depends on the CRE and f_{dp}. First, in F0.28, F0.28NC and F0.57NC, subsidence across the equatorial region is associated with an equatorward MSE flux at low latitudes (Figure 2e and Figure 5d, e). Secondly, eddy energy transport plays a role in the equatorward MSE flux in F0.28, F0.57, F0.28NC, F0.57NC, F0.85NC. Thirdly, in F0.85NC and F1.13NC a shallower Hadley circulation and reduced upper-tropospheric MSE reduces the MSE exported in the upper branches of the mean circulation, resulting in a net equatorward MSE transport. All other simulations (F0.85, F1.13, F1.70 and F1.70NC) have a single ITCZ associated with a positive $[AEI]_0$ and poleward MSE transport at low latitudes.

d. Sensitivity of the ITCZ to convective mixing with a prescribed cloud radiative effect.

To further understand the role of the CRE on the sensitivity of the ITCZ to convective mixing, we perform prescribed-CRE simulations and vary f_{dp} (Table 2). The prescribed CRE is diagnosed from single-year simulations with f_{dp} equal to 1.13 or 0.57 (section 2). The effect of prescribing the diurnal cycle of the CRE in a simulation with the same f_{dp} is minimal; for example, the ITCZ is similar in F1.13PC1.13 and F1.13 (Figure 1 and 10). Hence, we only discuss the mean circulation in F1.13PC0.57 and F0.57PC1.13 (Figure 11a and c).

Similar to CRE-off simulations, the sensitivity of the ITCZ to f_{dp} reduces in prescribed CRE simulations (Figure 10a) compared to CRE-on simulations (Figure 1a), associated with a reduced sensitivity of the AEI to f_{dp} (Figure 10b, 12a and c). The prescribed CRE heating acts as a fixed MSE source, which requires an increase in MSE export and hence increased convective activity. In PC1.13 simulations the CRE maximises at the equator, which is associated with increased equatorial convective activity and a single ITCZ. In PC0.57 simulations on the other hand, the CRE peaks off the equator and promotes a double ITCZ. The root mean squared difference of tropical precipitation and the mass meridional streamfunction illustrates that prescribing the CRE reduces the sensitivity of the ITCZ and Hadley circulation to f_{dp} by $\approx$50% (Table 3). Whilst the CRE plays a role in the sensitivity of the ITCZ to convective mixing (for example, comparing F1.13PC1.13 and F1.13PC0.57 in Figure 10a), the ITCZ and Hadley circulation are still sensitive to f_{dp}. For example, reducing f_{dp} (F0.57PC1.13) leads to a weakening in the upper branch of the mean circulation whilst changing the prescribed CRE (F1.13PC0.57) intensifies the upper branch of the Hadley circulation as the higher f_{dp} value is associated with a cooler upper-troposphere, hence, an intensified upper branch of the mean circulation is required for similar MSE transport (comparing F1.13 in Figure 2b to F0.57PC1.13 and F1.13PC0.57 in Figure 11c and a, respectively). The response of convection to changes in convective mixing is partially offset by the effect of prescribing the location of the CRE.

As in CRE-off simulations, AEI changes in prescribed CRE simulations when varying f_{dp} are predominantly driven by latent heat flux variations. For example, between F1.13PC1.13 and F0.57PC1.13, the equatorial latent heat flux reduces whilst the off-equatorial latent heat flux increases (Figure 12a). These changes are partially offset by changes in the clear-sky radiation, associated with a decrease in the TOA outgoing longwave radiation, due to an increase in atmospheric water vapour content. As changes in the ITCZ are associated with AEI changes, we conclude that the remaining sensitivity of the ITCZ to f_{dp} in prescribed CRE simulations is associated with latent heat flux variations. In simulations where the prescribed CRE is varied but the same f_{dp} value is used, AEI changes are mostly associated with cloudy-sky radiation (Figure 12b, d). However, latent heat flux variations are of the same order of magnitude as when varying f_{dp}. Using the same technique described in section 3c, we conclude that a shallower, weaker Hadley circulation is primarily responsible for changes in the MSE transport by the mean circulation when reducing f_{dp} or changing the prescribed CRE from PC1.13 to PC0.57 (not shown).

F1.13PC0.57 and F0.57PC1.13 have similar, split ITCZs (Figure 10a), yet very different AEI profiles (Figure 10b, Figure 11b and d). F0.57PC1.13 highlights that a double ITCZ in precipitation does not require a negative $[AEI]_0$ or an equatorward MSE transport (green and black line respectively in Figure 11d), illustrating that a double ITCZ in precipitation is not necessarily associated with an equatorward MSE flux at low latitudes. Instead a negative $[AEI]_0$ is a sufficient but not a necessary condition for a double ITCZ in precipitation. Due to local energy conservation issues, which are discussed further in section 4, it is challenging to understand F1.13PC0.57, which shows a negative $[AEI]_0$ and a positive equatorial $\partial_y[\widehat{vh}]$ (Figure 11b), (contradicting (1) as steady-state has been reached).

4. Discussion

We have analysed aquaplanet simulations with variations to convective mixing to show an association between resultant variations in the AEI and characteristics of the ITCZ. Using the AEI framework we have shown the importance of the CRE in the sensitivity of the ITCZ to convective mixing. In a single ITCZ scenario (F0.85, F1.13 and, F1.70), the CRE is critical in maintaining a positive $[AEI]_0$. For example, the $[AEI]_0$ would be negative without the CRE in F1.13 and F1.70, associated with a double ITCZ. Changes in cloudy-sky radiation are the dominant cause of AEI changes when varying the convective mixing rate, leading to our hypothesis that prescribing the CRE would remove or reduce the sensitivity of the ITCZ to convective mixing. The fact that the sensitivity of the ITCZ to f_{dp} remains in CRE-off and prescribed CRE simulations highlights the importance of other AEI components, in particular the latent heat flux. All simulations, with the exception of F0.57PC1.13, are consistent with Bischoff and Schneider (2016): a positive $[AEI]_0$ is associated with a single ITCZ and a negative $[AEI]_0$ with a double ITCZ.

CRE-off simulations illustrate that the CRE plays a substantial role in the structure and intensity of the ITCZ. Similar to Harrop and Hartmann (2016), we observe that removing the CRE cools the tropical upper-troposphere, reducing atmospheric stability and resulting in deep convection over cooler SSTs. Stronger convection at higher latitudes reduces equatorial moisture convergence and is associated with a double ITCZ. Removing the CRE also weakens the Hadley circulation which is associated with a reduced AEI gradient between the tropics and sub-tropics, in agreement with Popp and Silvers (2017). The sensitivity of the ITCZ to f_{dp} reduces when removing the CRE, agreeing with our hypothesis that prescribing the CRE would either remove or reduce the sensitivity of the ITCZ to convective mixing. Quantifying the reduction in sensitivity of the ITCZ to f_{dp} when removing the CRE remains a challenge due to strong dependence on the chosen metric and range of f_{dp}. It should also be noted that when removing the CRE other AEI components change, such that the AEI change is not equal to the total CRE that is removed.

In prescribed CRE simulations, ITCZ characteristics are sensitive to both the prescribed CRE and f_{dp}, however the sensitivity of the ITCZ to f_{dp} reduces by $\approx$50% (Table 3). In prescribed CRE simulations the response of convection to changes in convective mixing is offset by the effect of prescribing the location of the CRE. Heating associated with the prescribed CRE is a MSE source, therefore to increase the MSE exported, convective activity increases. The reduction in sensitivity compliments work by Voigt et al. (2014), who found that prescribing the CRE reduced the sensitivity of the ITCZ to hemispheric albedo perturbations to a similar degree. Thus, the role of the CRE in the sensitivity of the ITCZ to both variations in the convection scheme and boundary forcing appear similar, based on these two studies.

In both CRE-off and prescribed CRE simulations, latent heat flux alterations, associated with circulation changes, are the predominant cause of AEI changes when varying f_{dp}. Circulation changes when varying f_{dp} in CRE-off simulations are not associated with clear-sky flux variations, consistent with Harrop and Hartmann (2016), which concluded that changes in the clear-sky radiative cooling do not change the modelled circulation. Mobis and Stevens (2012) highlighted the importance of surface fluxes in reducing the sensitivity of the ITCZ to the convective parameterisation scheme when prescribing the wind speeds in the computation of surface fluxes. Numaguti (1993) and Liu et al. (2010) also concluded that variations in surface evaporation are associated with the ITCZ structure. We highlight that the sensitivity of the ITCZ to convective mixing is predominantly associated with the surface fluxes in the absence of cloud feedbacks.

As noted earlier in sections 3c and 3d, the balance between the diagnosed AEI and diagnosed $\partial_y[\widehat{vh}]$, equation (1), does not hold locally in MetUM. The mean of the maximum absolute diagnosed imbalance across the tropics amongst simulations is 13.4 Wm^{-2}. More importantly, the diagnosed equatorial energy imbalance ranges from 6.94 Wm^{-2} in F0.28NC to -20.63 Wm^{-2} in F1.70 with a mean absolute error of 9.89 Wm^{-2}. For all of our simulations apart from F1.13PC0.57, the sign of the equatorial $d_y[vh]$ and $[AEI]_0$ are the same, and therefore using $[AEI]_0$ as a proxy for the direction of energy transport at low latitudes is still valid. In F1.13PC0.57 the difference between the diagnosed $d_y[vh]$ and $[AEI]$ is -16.9 Wm^{-2}; the equatorial $d_y[vh]$ is positive and $[AEI]_0$ is slightly negative (Figure 11b). Whilst the local energy imbalance is a concern for F1.13PC0.57, we argue that in all other simulations the local energy imbalance does not affect our conclusions. There are a number of possible reasons for the localised imbalance of the AEI budget including: non-conservation associated with the semi-Lagrangian advection scheme in MetUM; the use of dry and moist density in different components of the MetUM dynamics and physics; errors in our diagnosis of the MSE budget, for example, not considering density changes within a timestep; or, using an Eulerian approach for diagnosing the energy transport which is inconsistent with the semi-Lagrangian advection scheme. It is worth noting that other studies using the AEI framework have not shown that the MSE energy budget is locally closed, and this problem may not be unique to our study. Nevertheless, the local energy imbalance has challenged our interpretation of some simulations, and highlights that future modelling studies using an atmospheric MSE budget should be cautious.

Variations in the CRE when varying f_{dp} can lead to a negative $[AEI]_0$ associated with a net equatorward MSE energy transport at low latitudes. Whilst the predominant response to a negative $[AEI]_0$ is a double ITCZ associated with equatorward energy transport at low latitudes by the mean circulation (F0.28, F0.28NC and F0.57NC), F0.57, F0.85NC and F1.13NC have shown that a net equatorward MSE transport can occur at low latitudes even with a poleward energy transport by the mean flow at the tropopause. Two mechanisms can lead to this. Firstly, the MSE flux due to eddies contributes a substantial proportion to the total MSE flux (as seen in Figure 11 12b and d), and this can support equatorward MSE transport. In F0.57, the MSE flux due to eddies is responsible for a net equatorward energy transport in the deep tropics. This invalidates the assumption that the energy flux equator is associated with zero MSE transport by the mean circulation, as in Bischoff and Schneider (2016). This is also supported by the equatorward displacement of the energy flux equator (from 2 and 4) relative to maximum precipitation in all simulations except for F0.85NC and F1.70NC (Table 4). The second mechanism (F0.85NC and F1.13NC) is a change in the MSE transport direction due to a shallower Hadley circulation and a lower MSE in the upper-troposphere (section 3c). These changes reduce the MSE export in the upper branch of the Hadley circulation, resulting in an equatorward MSE transport by the mean circulation at low latitudes.

In our aquaplanet configuration SSTs are fixed which implies an arbitrary but varying oceanic heat transport to maintain SSTs given a net surface heat flux imbalance. Thus, our aquaplanet experiments may be viewed as energetically inconsistent. In Bischoff and Schneider (2014) and Voigt et al. (2016) ocean heat transport, and hence the net downward flux at the surface, is fixed, constraining the response of AEI components and potentially reducing the sensitivity of the ITCZ to model perturbations. In reality the ocean circulation, and thus ocean heat transport, is sensitive to changes in the surface wind stress. Therefore, both the SST and ocean heat transport could change in response to tropical circulation changes from variations to f_{dp} or the prescribed CRE. Recent work has shown that the ocean circulation plays an important role in the meridional transport of energy (Green and Marshall 2017), and that sensitivities of the ITCZ found in atmosphere-only simulations do not necessarily hold in a fully coupled model. For example, coupling reduces the sensitivity of the ITCZ to an interhemispheric albedo forcing (e.g. comparing Kay et al. (2016) and Hawcroft et al. (2017) to Voigt et al. (2014)). The radiative effect of clouds on the surface and Ekman heat transport associated with a single ITCZ would be expected to reduce the equatorial SST gradient, which would promote a double ITCZ (Numaguti 1995; Mobis and Stevens 2012) and may reduce the sensitivity of the ITCZ to convective mixing. Coupled simulations with an interactive ocean are required to further investigate the sensitivity of the ITCZ to the CRE and convective mixing.

5. Conclusions

The double ITCZ bias is a leading systematic error across a hierarchy of models (Li and Xie 2014; Oueslati and Bellon 2015). Inter-model variability in the ITCZ structure persists even in a highly-idealised framework such as an aquaplanet with prescribed SSTs (Blackburn et al. 2013). This study confirms and extends previous research that variations in the convective parameterisation scheme and convective mixing can alter the ITCZ (Figure 1a; Hess et al. 1993; Numaguti 1995; Chao and Chen 2004; Liu et al. 2010; Mobis and Stevens 2012). Higher convective mixing rates are associated with a single ITCZ whilst lower rates are associated with a double ITCZ. As the convective mixing rate reduces, convection at higher latitudes strengthens, decreasing equatorward low-level winds at low latitudes, promoting a double ITCZ structure. The sensitivity of the ITCZ to convective mixing is associated with AEI changes, predominantly caused by CRE variations. For example, the CRE plays an important role in maintaining a positive equatorial AEI, and is therefore associated with a single ITCZ structure (consistent with Harrop and Hartmann (2016) and Bischoff and Schneider (2016)'s framework). When removing the CRE, the response of the ITCZ depends on the convective mixing rate. At low convective mixing rates, where a double ITCZ is simulated with the CRE, precipitation bands shift poleward. At high convective mixing rates the ITCZ broadens, whilst at certain convective mixing rates the ITCZ structure changes from single to double. Quantifying whether the sensitivity of the ITCZ to convective mixing reduces when removing the CRE is challenging, as the sensitivity depends on the range of convective mixing rates and the chosen metric. Prescribing the CRE reduces the sensitivity of the ITCZ to convective mixing by $\approx$50%. When removing or prescribing the CRE other AEI components, in particular the latent heat flux, play a role in the sensitivity of the ITCZ to convective mixing. Hence, simulations where the ocean heat transport is fixed, thereby constraining surface fluxes, may underestimate the sensitivity of the ITCZ to changes in model formulation. We have also shown two mechanisms responsible for a net equatorward MSE transport even with no equatorial subsidence: MSE transport by eddies; and a reduced MSE export in the upper branch of the mean circulation due to a shallower Hadley circulation. These mechanisms highlight that caution should be taken when associating changes in the AEI to the ITCZ structure.

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LIST OF TABLES

TABLE 1. Simulations varying f_{dp} with cloud-radiation interactions on (CRE-on) and off (CRE-off). F1.13 is the default integration for GA6.0.

f_{dp}	CRE-on	CRE-off
0.28	F0.28	F0.28NC
0.57	F0.57	F0.57NC
0.85	F0.85	F0.85NC
1.13	F1.13	F1.13NC
1.70	F1.70	F1.70NC

TABLE 2. Simulations with a prescribed climatology of the CRE diurnal cycle. PC1.13 and PC0.57 represent the prescribed CRE diurnal cycle from a one-year simulation where f_{dp} equals 1.13 or 0.57 (respectively).

f_{dp}	PC1.13	PC0.57
1.13	F1.13PC1.13	F1.13PC0.57
0.57	F0.57PC1.13	F0.57PC0.57

TABLE 3. Root mean squared difference for tropical precipitation and mass meridional streamfunction between two simulations. Tropical domain defined as 30°N to 30°S. Percentage value is the percentage reduction compared to F0.57 and F1.13.

Simulations	Precipitation (mm day^{-1})	Mass Meridional Streamfunction ($\times 10^{10}$ kg s^{-1})
F0.57 & F1.13	2.84	1.78
F0.57PC1.13 & F1.13PC1.13	1.18 (58%)	0.67 (62%)
F0.57PC0.57 & F1.13PC0.57	1.65 (42%)	0.96 (46%)

TABLE 4. AEI_0, location of ITCZ and approximate energy flux equator (δ) using equation 2 or 4 in each simulation. A single/double ITCZ is assumed when AEI_0 is positive/negative, respectively. Not applicable (N/A) occurs when AEI_0 and $\partial_{yy}([AEI])_0$ are both negative and therefore the square root of $-\frac{6([AEI])_0}{\partial_{yy}([AEI])_0}$ has an imaginary component.

Simulation	AEI_0 (W m^{-2})	ITCZ location ($^\circ$)	Energy Flux Equator (δ) location ($^\circ$)
F0.28	-18.1	8.13/-8.13	6.85/-7.06
F0.57	-5.9	4.38/-4.38	0.84/-2.87
F0.85	33.4	1.88	-0.41
F1.13	36.7	0.63	0.22
F1.70	33.7	0.63	0.30
F0.28NC	-4.9	9.38/-9.38	N/A
F0.57NC	-12.2	8.13/-8.13	N/A
F0.85NC	-18.3	6.88/-5.63	6.48/-6.80
F1.13NC	-5.9	4.38/-4.38	3.21/-3.58
F1.70NC	2.0	1.88	2.73
F1.13PC1.13	33.6	0.63	0.16
F1.13PC0.57	-1.7	3.13/-3.13	0.19/-1.75
F0.57PC1.13	20.6	3.13	-0.12
F0.57PC0.57	-14.2	4.38/-4.38	2.70/-2.64

LIST OF FIGURES

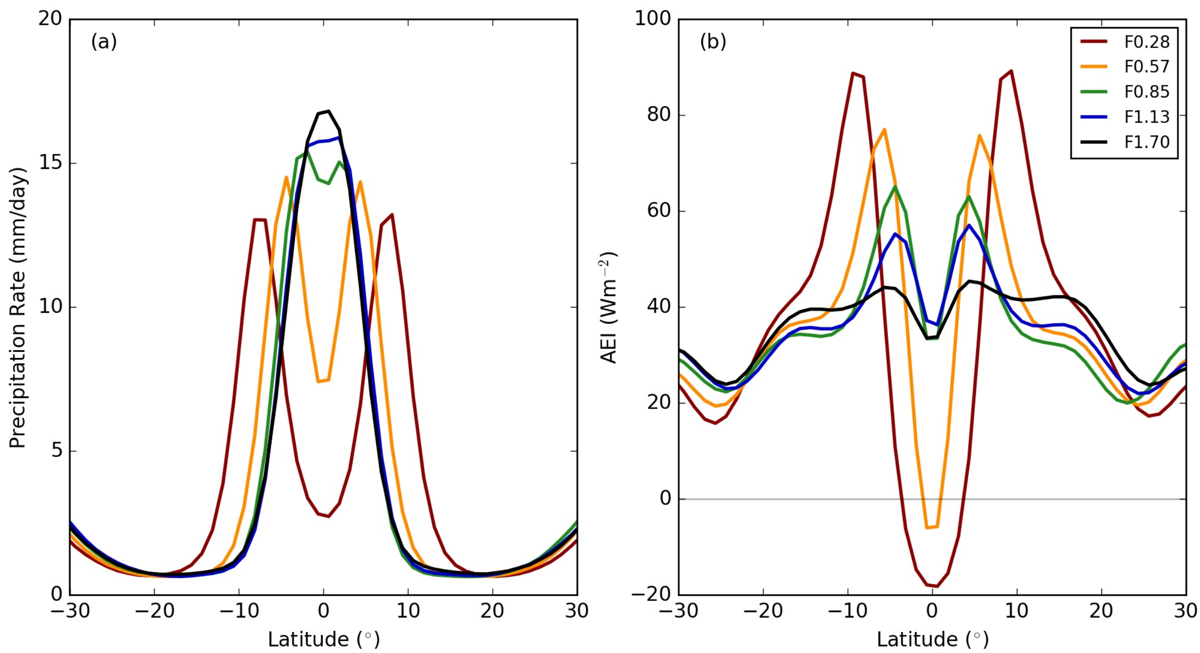
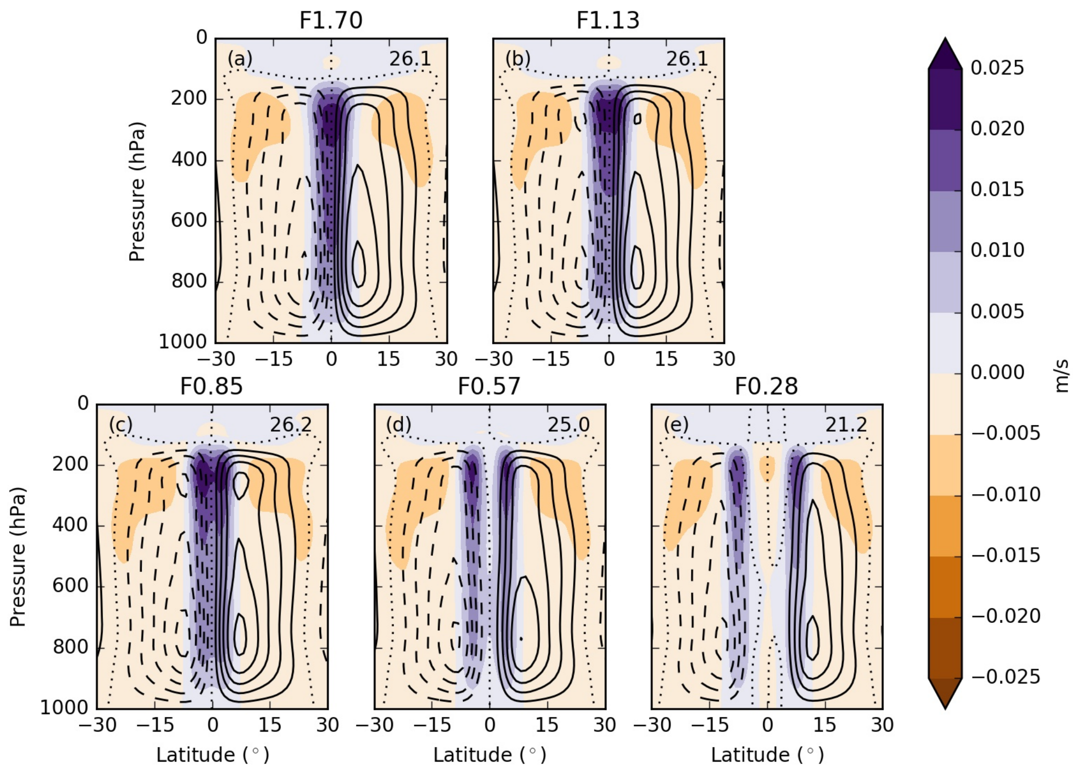


FIG. 1. Zonal-, time-mean (a) precipitation rates (mm day^{-1}) and (b) AEI (W m^{-2}) in simulations where f_{dp} is varied.



772 FIG. 2. Zonal-, time-mean mass meridional streamfunction (kg s^{-1}) (lined contours) and vertical wind speed

773 (m s^{-1}) (filled contours). Lined contours are in intervals of 5×10^{10}, with dashed contours representing negative

774 values. Dotted contour is zero value. (a): F1.70, (b): F1.13, (c): F0.85, (d): F0.57, (e): F0.28. Maximum value

775 of the mass meridional streamfunction printed in top right-hand corner of each subplot.

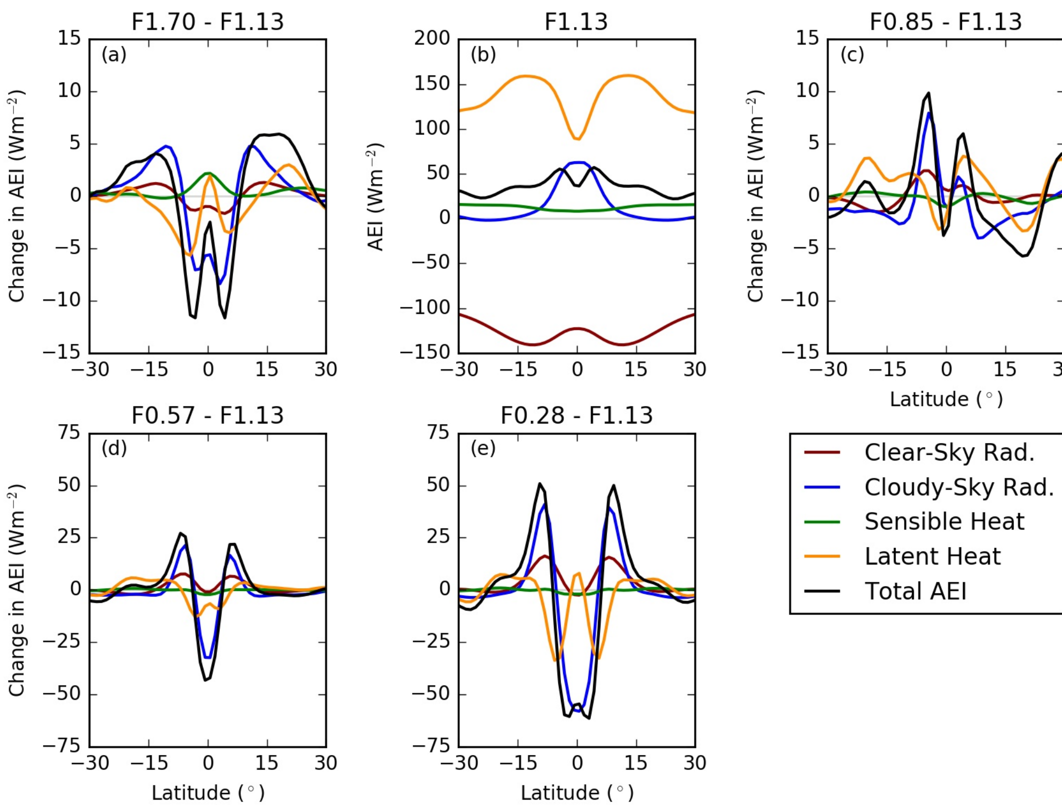


FIG. 3. Zonal, time-mean AEI components (W m^{-2}). (b): F1.13 and (a),(c)-(e): Change in AEI components compared to F1.13 for (a) F1.70; (c) F0.85, (d) F0.57, (e) F0.28. Red line is the clear-sky component, blue line is the cloudy-sky component. Green and orange lines represent the sensible and latent heat flux, respectively, and the black line is the total change in AEI. Note, (a) and (c) have axis limits -15 and 15 W m^{-2}, whilst (d) and (e) have limits -75 and 75 W m^{-2}.

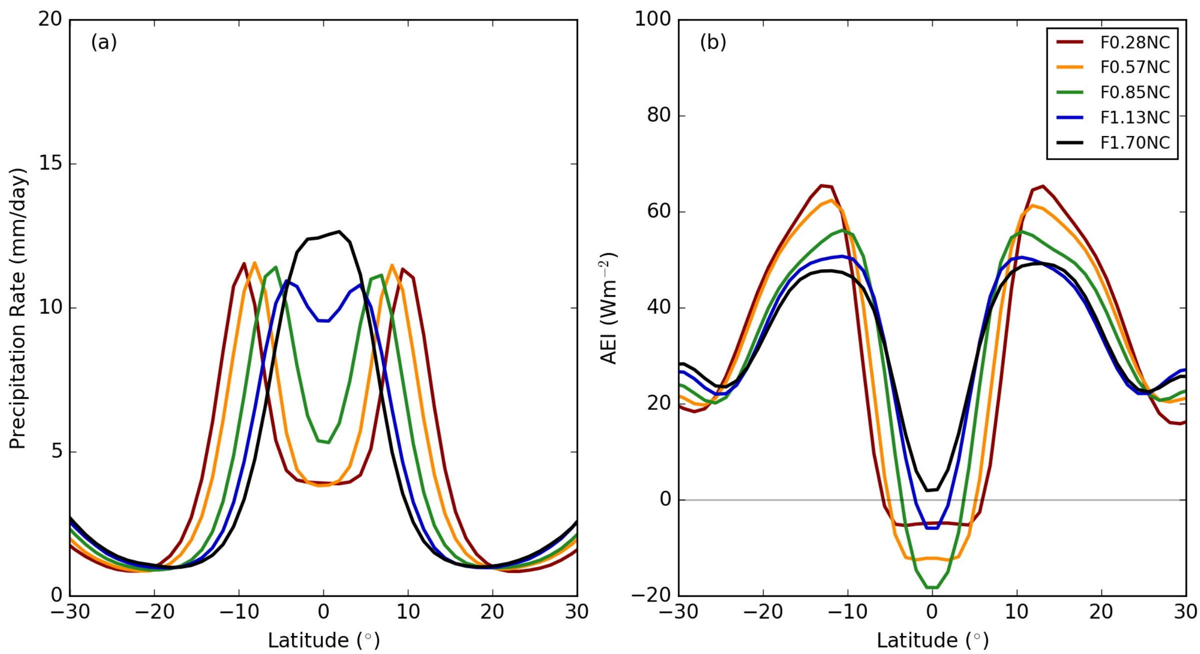


FIG. 4. Zonal-, time-mean (a) precipitation rates (mm day^{-1}) and (b) AEI (W m^{-2}) in CRE-off simulations where f_{dp} is varied.

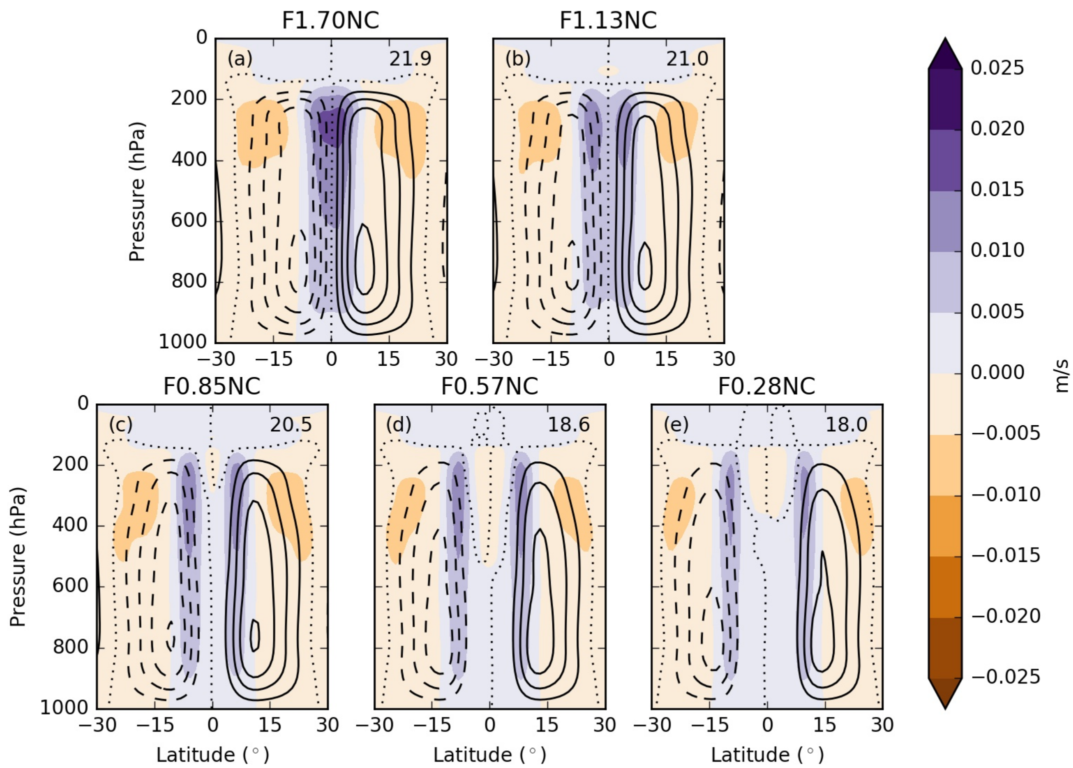


FIG. 5. Zonal-, time-mean mass meridional streamfunction (kg s^{-1}) (lined contours) and vertical wind speed (m s^{-1}) (filled contours). Lined contours are in intervals of 5×10^{10}, with dashed contours representing negative values. Dotted contour is zero value. (a) F1.70NC, (b) F1.13NC, (c) F0.85NC, (d) F0.57NC, (e) F0.28NC. Maximum value of the mass meridional streamfunction printed in top right-hand corner of each subplot.

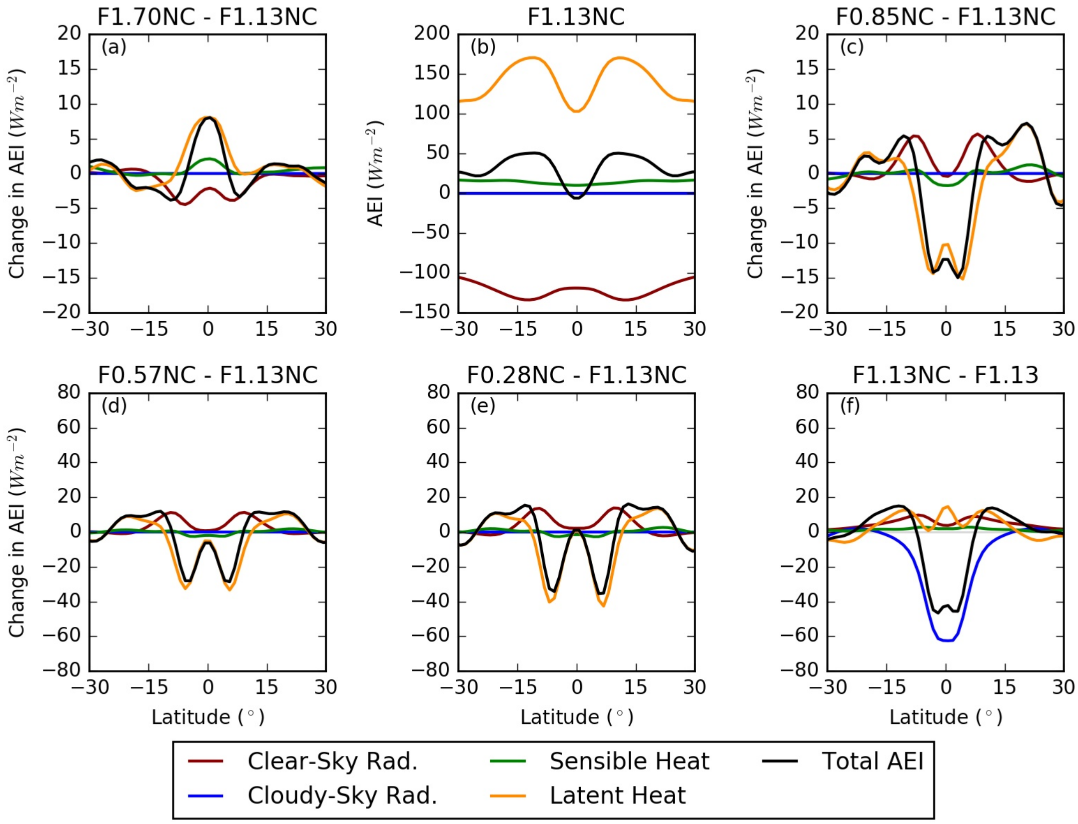


FIG. 6. Zonal, time-mean AEI components (W m^{-2}). (b) F1.13NC (a),(c)-(e): Change in AEI components compared to F1.13NC for (a) F1.70NC; (c) F0.85NC, (d) F0.57NC, (e) F0.28NC and (f): Change in AEI components when removing cloud-radiation interactions at f_{dp} equals 1.13 (F1.13NC - F1.13). Note, (a) and (c) have axis limits -20 and 20 W m^{-2} whilst (d)-(f) have limits -80 and 80 W m^{-2}.

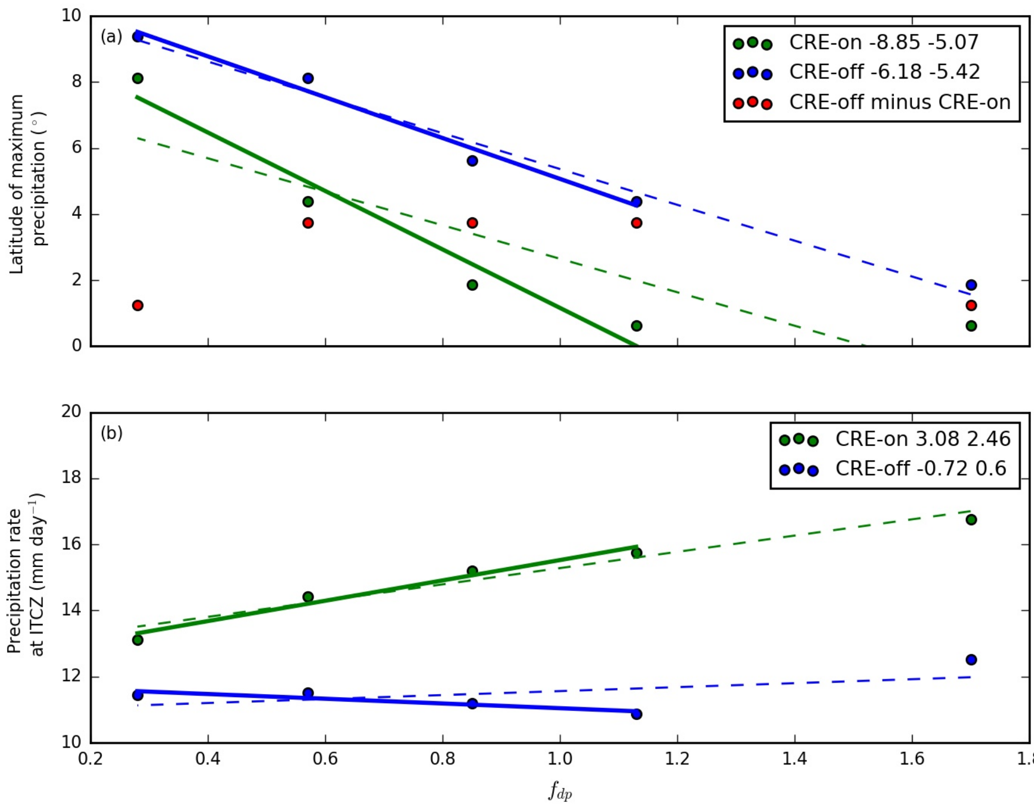


FIG. 7. Diagnostics for determining the sensitivity of the ITCZ to f_{dp} in CRE-on (green) and CRE-off (blue) simulations. Top (a): Latitude of maximum precipitation (°), bottom (b): Precipitation rate at ITCZ (mm day^{-1}). Four regression lines are plotted in each subplot. Solid lines where $0.28 \leqslant f_{dp} \leqslant 1.13$ and dashed lines where $f_{dp} \leqslant 1.70$. The slope of each regression line is printed in the legend. First value where $0.28 \leqslant f_{dp} \leqslant 1.13$ and second value where $f_{dp} \leqslant 1.70$.

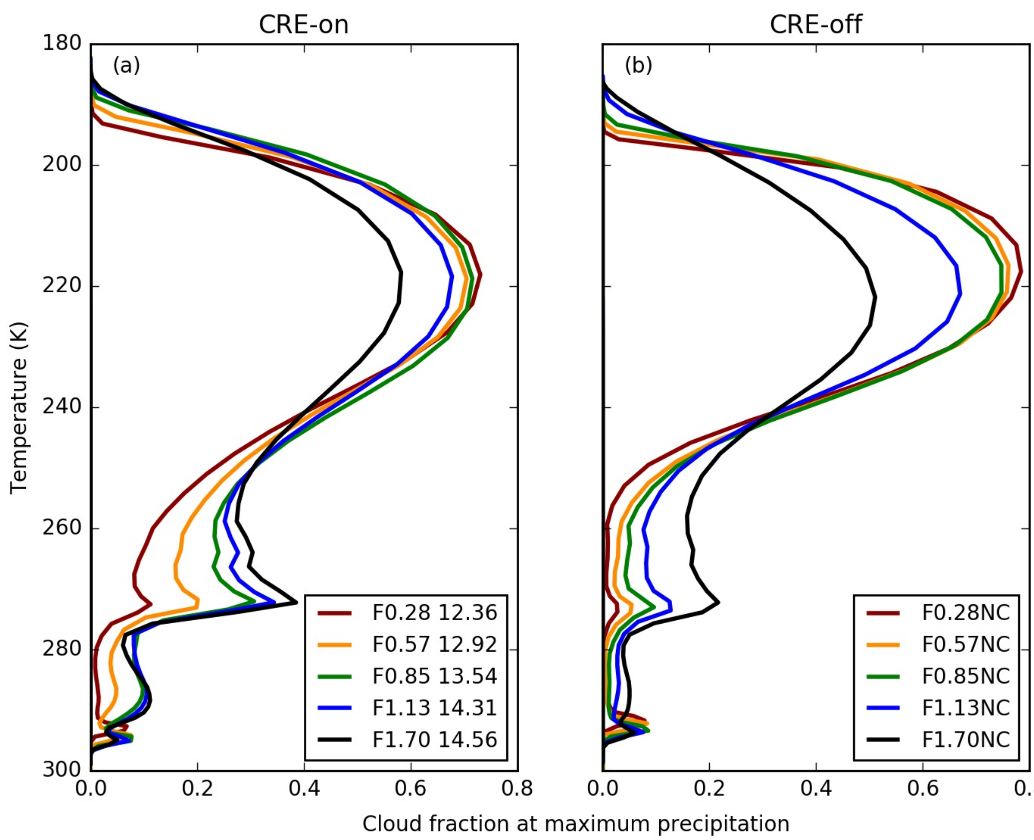


FIG. 8. Zonal-, time-mean cloud fraction against temperature (K) at latitude of maximum precipitation. Left (a): CRE-on simulations, right (b): CRE-off simulations. Printed in legend, the tropical-domain average CRE (W m^{-2}) for CRE-on simulations.

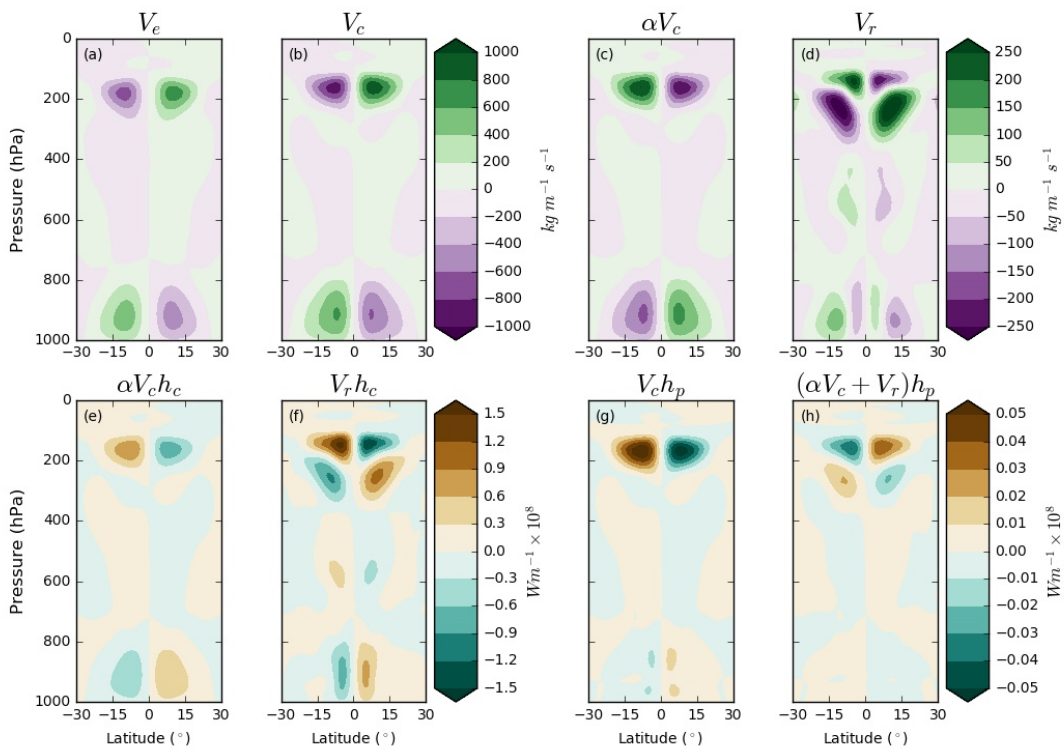
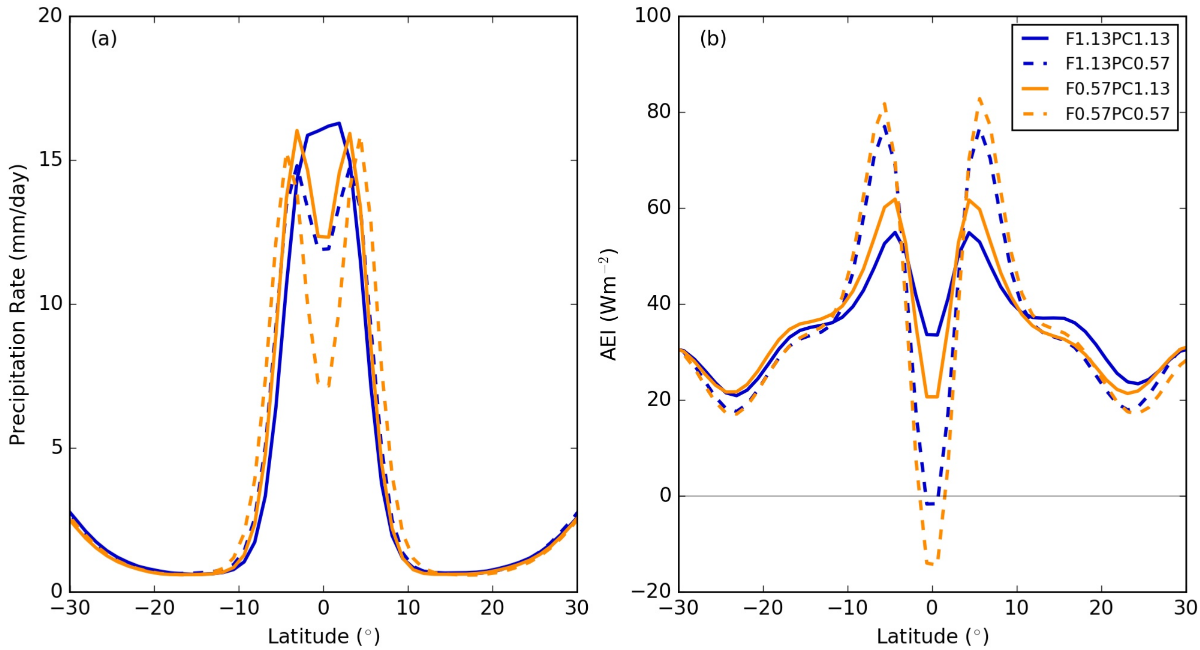


FIG. 9. Top row: (a) and (b): Meridional mass flux (kg m^{-1} s^{-1}) in F1.13NC and F1.13 respectively, (c) and (d): Change in meridional mass flux due to change in circulation strength and change in meridional wind, respectively. Bottom row: Components of MSE flux change (W m^{-1}), equation (10), due to (e), circulation intensity changes $\alpha V_c h_c$, (f), changes in circulation structure $V_r h_c$, (g), MSE profile changes $V_c h_p$, and, (h), MSE changes correlated with changes in circulation structure and strength $(\alpha V_c + V_r) h_p$. Analysis explained in Section 3c.



805 FIG. 10. Zonal-, time-mean (a) precipitation rates (mm day^{-1}) and (b) AEI (W m^{-2}) in simulations with a
806 prescribed CRE.

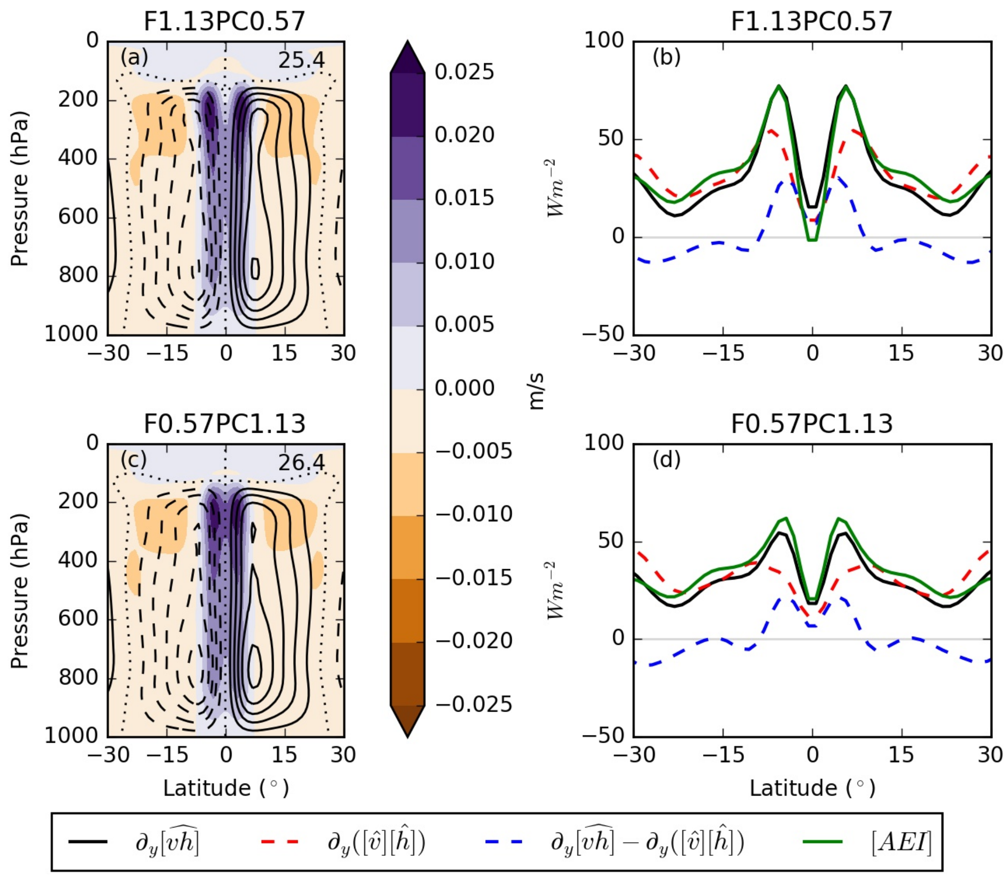


FIG. 11. Left: Zonal-, time-mean mass meridional streamfunction (kg s^{-1}) (lined contours) and vertical wind speed (m s^{-1}) (filled contours) for (a) F1.13PC0.57 and (c) F0.57PC1.13. Lined contours are in intervals of 5 $\times$ 10^{10}, with dashed contours representing negative values. Dotted contour is zero value and maximum value of the mass meridional streamfunction printed in top right-hand corner of each subplot. Right: Divergence of the MSE flux (W m^{-2}) and AEI for (b) F1.13PC0.57 and (d) F0.57PC1.13. Solid black line - Divergence of total MSE flux $\partial_y[\widehat{vh}]$, red dotted line - MSE flux due to mean circulation $\partial_y[\hat{v}][\hat{h}]$, blue line - $\partial_y[\widehat{vh}] - \partial_y[\hat{v}][\hat{h}]$, green line - $[AEI]$.

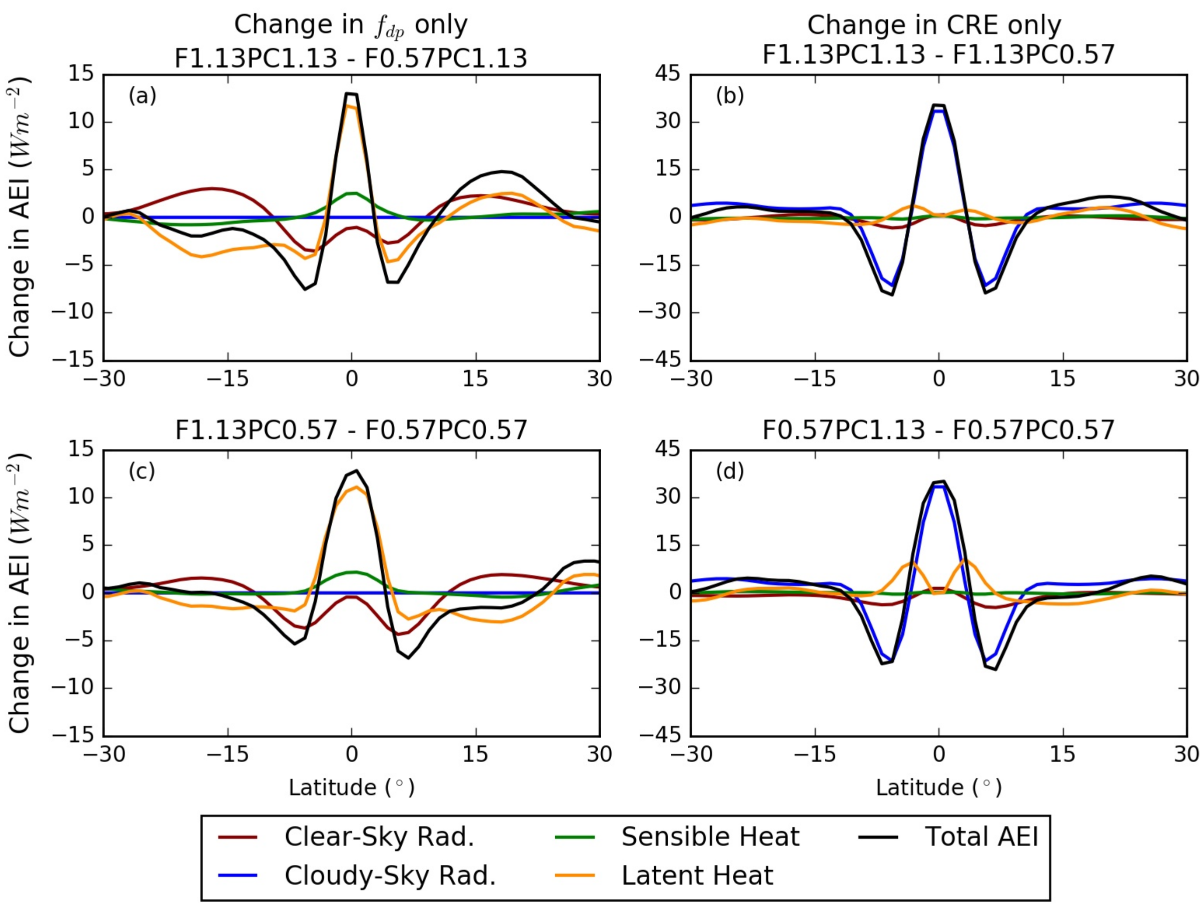


FIG. 12. Changes in zonal-, time-mean AEI contributions (W m^{-2}) for prescribed CRE simulations. Comparison of simulations with same f_{dp} constant (a, c) have y-axis limits of -15 to 15 W m^{-2}, whilst those with a different prescribed CRE (b, d) have y-axis limits -45 to 45 W m^{-2}.