

# *Storyline approach to the construction of regional climate change information*

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**Storyline approach to the construction of regional climate-change information**

Theodore G. Shepherd, Department of Meteorology, University of Reading, PO Box 243, Earley Gate, Whiteknights, Reading RG6 6BB, U.K.

theodore.shepherd@reading.ac.uk, tel. 0118 378 8957

**Abstract**

Climate science seeks to make statements of confidence about what has happened, and what will happen (conditional on scenario). The approach is effective for the global, thermodynamic aspects of climate change, but is ineffective when it comes to aspects of climate change related to atmospheric circulation, which are highly uncertain. Yet atmospheric circulation strongly mediates climate impacts at the regional scale. In this way the confidence framework, which focuses on avoiding Type 1 errors (false alarms), raises the prospect of committing Type 2 errors (missed warnings). This has ethical implications.

At the regional scale, however, where information on climate change has to be combined with many other factors affecting vulnerability and exposure — most of which are highly uncertain — the societally relevant question is not “What will happen?” but rather “What is the impact of particular actions under an uncertain regional climate change?” This re-framing of the question can cut the Gordian Knot of regional climate-change information, provided one distinguishes between epistemic and aleatoric uncertainties — something that is generally not done in climate projections. It is argued that the storyline approach to climate change — the identification of physically self-consistent, plausible pathways — has the potential to accomplish precisely this.

**Keywords:** Climate change, climate ethics, uncertainty, atmospheric circulation, climate impacts

**Non-technical summary:**

This study addresses the challenge of how to construct useful climate-change information at the regional scale in the face of the large uncertainties that exist, whilst retaining the relevant information concerning climate risk. It is argued that the usual methods of constructing climate information are not as objective or value-free as they might seem to be. An alternative ‘storyline’ approach, which emphasizes plausibility over probability, has been proposed as a way to provide climate information relevant to decision-making. It is shown that the two approaches can be cast within a common framework.

## 1. Introduction

Although there is high confidence in thermodynamic aspects of climate change (global warming, sea-level rise, atmospheric moistening, melting of ice), the levels of confidence concerning dynamical aspects of climate change, such as the location and strength of storm tracks, are much lower [1]. None of the three key lines of evidence used in climate-change science — predicated by accepted theory, detected in observations, and consistently represented in climate models — apply to aspects of climate change that are closely related to large-scale atmospheric circulation. This includes, notably, mean precipitation changes over many of the most populated regions on Earth (Figure 1). It is in striking contrast to thermodynamic aspects of change, at least when sufficiently aggregated [3], where all three lines of evidence apply [2].

Lack of agreed-upon theoretical predictions is related to the fact that different drivers of change can act in opposite directions, so the result is often a small difference of large terms [4,5]. Lack of detection in observations is related to the small signal-to-noise ratio of forced circulation changes, reflecting the fact that climate variability is primarily a dynamical phenomenon [6]. Lack of model agreement is related to both these issues, and to the fact that circulation changes are often quite sensitive to model biases, which can be substantial [7-9].

Furthermore, thermodynamic aspects of climate can be described by extensive quantities (e.g. heat content or ocean volume), which can be readily aggregated, and strong conclusions can be drawn from thermodynamic principles alone, often in terms of global budgets [10]. In contrast, circulation aspects of climate are inherently regional, and involve dynamics (Newton's second law) as well as thermodynamics. Since dynamics is also inherently chaotic, the challenge of atmospheric circulation should come as no surprise.

Ways must therefore be found to construct useful scientific information on the regional scale, and even on the local scale, that reflect an appropriate level of uncertainty yet retain the relevant information about climate risk. It has recently been argued [11] that *storylines* — physically self-consistent unfoldings of past events, or of plausible future events or pathways — provide a potential way forward, both for the interpretation of the observed record and for the description of plausible futures. However, storylines are inherently subjective and thus would seem to be at odds with more probabilistic approaches, which give the appearance of objectivity. The purpose of this paper is to place storylines within a broader epistemological framework.

It is first shown (Section 2) how the standard, confidence-based framework for the construction of climate information prioritizes reliability (the avoidance of Type 1 errors, or false alarms) over informativeness (the

avoidance of Type 2 errors, or missed warnings), and thus has ethical implications. It follows that there is no such thing as value-free climate science. In Section 3, the difference between epistemic and aleatoric (random) uncertainty is shown to be critical to the treatment of climate risk. Since epistemic uncertainty is deterministic and inherently subjective, it follows that there is no objective basis for a probabilistic approach, and no such thing as objective climate information. This motivates a re-framing of the climate risk question from the ostensibly objective prediction space into the explicitly subjective decision space (Section 4). Finally, it is shown in Section 5 how such a re-framing can be cast within the mathematical framework of a causal network, thereby reconciling storyline and probabilistic approaches.

## **2. The confidence straightjacket**

The most authoritative statements on physical aspects of climate change come from Working Group I (WGI) of the Intergovernmental Panel on Climate Change (IPCC). In the Summary for Policymakers of the last (5<sup>th</sup>) IPCC WGI Assessment Report [2], atmospheric circulation is scarcely mentioned, and all the statements of confidence are based on thermodynamics. This remarkable fact evidences better than anything else the lack of scientific consensus on dynamical aspects of climate change. Moreover, the statements of confidence are crafted to be reliable, generally by emphasizing global rather than regional aspects of change. A good example is the headline statement on the water cycle:

“Changes in the global water cycle in response to the warming over the 21<sup>st</sup> century will not be uniform. The contrast in precipitation between wet and dry regions and between wet and dry seasons will increase, although there may be regional exceptions.” [2]

This statement is based on the sound physical principle that, all else being equal, a moister atmosphere will exhibit an accelerated hydrological cycle [10]. The statement achieves its reliability in the tropics by including oceanic regions (see Figure 1); indeed a key observation supporting the statement is the increased salinity in the subtropical upper oceans (due to increased evaporation). However, it is precipitation over land that matters for climate impacts, and there have been many studies showing that the “wet get wetter, dry get drier” paradigm does not hold over land regions [12-14], as is reflected in the general lack of stippling over these regions (apart from the high northern latitudes) in Figure 1. The statement is perfectly reliable as an explanation of how the global climate system works, but it does not provide useful information at the regional scale, as the final caveat makes clear. In this way, reliability is achieved at the price of informativeness.

To find a high-level statement on dynamical aspects of climate change in the IPCC WGI 5<sup>th</sup> Assessment Report, one must look one level down, in the Technical Summary [2]. The statements are uniformly characterized by low

125 levels of confidence and a lack of informativeness at the regional scale. An  
126 illustrative example is the statement on changes in Northern Hemisphere  
127 (NH) storm tracks, which are an important determinant of midlatitude  
128 weather:

129 “Substantial uncertainty and thus low confidence remains in projecting  
130 changes in NH storm tracks, especially for the North Atlantic basin.” [2]

131 Furthermore, IPCC WGI uses a likelihood scale in which the term “unlikely” is  
132 used to describe likelihoods of up to 33%. This terminology seems rather  
133 perverse from a lay perspective; in most areas of life, one would pay attention  
134 to likelihoods that high, especially if the consequences were serious — as they  
135 are with climate change. (Would you board an airplane if you were told that it  
136 had a 33% chance of crashing?) Yet in the WGI report, the term “unlikely” is  
137 generally used to dismiss rather than to highlight a possibility. Consider this  
138 example from the Technical Summary, again with reference to the North  
139 Atlantic storm track:

140 “...it is unlikely that the response of the North Atlantic storm track is  
141 a simple poleward shift” [2]

142 The context here is that, despite the lack of an agreed-upon theoretical  
143 explanation, the concept of a poleward storm track shift under climate change  
144 has become a general expectation [5]. However, projected changes in the  
145 North Atlantic storm track do not conform to that expectation [15]. An  
146 equivalent version of the statement would be “...it is likely that the response  
147 of the North Atlantic storm track [to climate change] is not a simple poleward  
148 shift”. Because in the present state of knowledge a consensus statement could  
149 not be crafted on what was likely to happen, the authors instead chose to  
150 emphasize what was not likely to happen. Yet there are several possibilities  
151 for what might happen, each with their own implications for climate risk,  
152 which could have been articulated (e.g. [16]). However, the simultaneous  
153 consideration of contradictory futures is not naturally expressed through  
154 statements of confidence. Thus, reliability is again achieved at the price of  
155 informativeness.

156 These examples illustrate the fact that by employing a confidence framework,  
157 which seeks to attribute what has happened and to predict what will happen  
158 (for a given climate forcing scenario), climate science winds up in something  
159 of a straightjacket when it comes to aspects of regional climate change that  
160 are closely related to large-scale atmospheric circulation, such as drought and  
161 storminess.

162 It is notable in this respect that IPCC Working Group II, which deals with  
163 impacts and adaptation, defines climate change as any observed change, not  
164 necessarily one that has been attributed to anthropogenic forcing [17]. This is

done to avoid the confidence straightjacket, but it creates a knowledge gap between the WGI and WGII science domains [18].

There is always a trade-off to be made between reliability and informativeness [19]. Yet a focus on reliability, guarding preferentially against Type 1 errors (false positives, i.e. false alarms), increases the likelihood of Type 2 errors (false negatives, i.e. missed warnings). It follows that much as though climate science might strive to be value-free, it cannot be: the way in which climate information is constructed has ethical implications [20]. Lloyd and Oreskes [20] raise the important question of why in climate science it has become normative that scientific rigour is associated with a focus on reliability. They point out that the decision on whether to preferentially guard against Type 1 or Type 2 errors is not a scientific one, but one of values. For example, in deciding whether to bring a new drug to market, one assesses both the drug's efficacy (guarding against Type 1 errors) and whether it has any unwanted side effects (guarding against Type 2 errors). Similarly, in deciding whether to issue an evacuation order for a city in the face of a forecasted storm, a balance of concern between Type 1 and Type 2 errors will be considered. Thus, there is nothing unscientific about seeking to guard against Type 2 errors.

It would seem entirely appropriate to preferentially guard against Type 1 errors when making high-level definitive statements concerning global climate change such as "Warming of the climate system is unequivocal" [2]. However, the framework is not so evidently appropriate when it comes to regional aspects of change (see also [21]). This situation seems to be an example of Kuhn's [22, p.37] important observation that "a paradigm can ... insulate the [scientific] community from those socially important problems that ... cannot be stated in terms of the conceptual and instrumental tools the paradigm supplies". Thus, it is imperative to find alternative paradigms.

### 3. Epistemic vs aleatoric uncertainty

Broadly, uncertainty in climate projections arises from three sources: uncertainty in future climate forcing, in the climate system response to that forcing (i.e. the change in climate), and in the actual realization of climate for a particular time window, which is subject to internal variability. The nature of these uncertainties is very different (e.g. [23]). The first depends primarily on human actions and is called the scenario, and the projections are normally made conditional on the scenario. The second is what is known as an *epistemic* uncertainty; there is only one truth, but we do not know what it is. The third is what is known as an *aleatoric* uncertainty; there is a random element to what will occur, whose probability is known to some extent. Any discussion of climate risk must address the central fact that the nature of the second and third uncertainties is fundamentally different. This is especially important for circulation-related aspects of climate change at the regional scale, for which these two sources of uncertainty tend to dominate the overall

uncertainty (see [24] for regional precipitation changes). Yet it is standard practice in climate science to mingle the two sources of uncertainty together, e.g. in the multi-model ensembles (with one realization taken from each model) that are in such widespread use [2]. In such ensembles the differences between the individual model projections include both the systematic differences between different model climates (epistemic) and the random differences that arise from the limited sampling of internal variability (aleatoric), which poses challenges in interpretation [25].

We first discuss the uncertainty arising from internal variability, since it is conceptually much easier to deal with. Internal variability is a property of the physical climate system, whose random character arises from the chaotic nature of atmospheric and oceanic dynamics, and which can be characterized from observations. Indeed, the definition of climate includes internal variability, which is characterized through statistical measures such as variances and co-variances of physical fields, as well as higher-order moments such as skewness or extremes, and includes coherent modes of variability such as the El Niño/Southern Oscillation phenomenon. The uncertainty from internal variability is fundamentally irreducible (leaving aside the possibility of finite-time prediction from specified initial conditions), and users of climate information need to understand that the mantra of “reducing uncertainty” is inappropriate in this case; rather, the scientific goal is to better quantify the uncertainty. The magnitude of the uncertainty for any particular quantity can be reduced by taking coarser spatial and temporal averages, but that operation changes and may simultaneously reduce the value of the information provided.

The concept of internal variability is not without ambiguity since climate has various sources of non-stationarity, and what is meant by internal variability is conditional on any non-stationary influence, including climate change itself. Furthermore, knowledge of internal variability is limited by the finite observational record, and there is uncertainty in how internal variability will respond to global warming. Nevertheless, in most cases, the main uncertainty in what climate conditions will be experienced at a particular place and time arising from internal variability can be considered to be aleatoric, and thus amenable to a straightforward (i.e. frequentist) probabilistic interpretation. The reliability of model simulations of internal variability can be similarly assessed, at least in principle.

The uncertainty in the climate response to forcing is conceptually very different. It is not a property of the physical climate system; rather, it is a property of a state of knowledge, or degree of belief, and it *can* be reduced as knowledge improves. In contrast to aleatoric uncertainty, which is objective, such epistemic uncertainty is *subjective* [26]. Therefore, treating epistemic uncertainty as if it were aleatoric, with a focus on the multi-model mean as a best estimate, has no epistemological justification. This has been recognized



for some time [27,28,21], but the practice continues to be normative (e.g. as in Figure 1). It is interesting to consider why this is so, since in most areas of science the essential distinction between systematic and random sources of uncertainty is well recognized. One of the reasons may be that the extent of the epistemic uncertainty is not particularly well known. First, climate models are imperfect representations of reality and share many deficiencies, thus may exhibit a collective bias and fail to explore important aspects of climate change. Second, even within the world represented by climate models, the forced circulation response of any particular model is obscured by internal variability.

As an example of the latter, Deser et al. [29] estimate that for NH wintertime midlatitude surface pressure (whose spatial gradient provides an indicator of circulation changes), ensemble sizes of around 30 are generally needed to determine the forced decadal changes of a given model over a 45-year period. This is in striking contrast to surface temperature changes, where the signal-to-noise ratio of the forced response is much larger, and even single simulations can be informative. One might be tempted to think that if such a large ensemble size is needed to detect the signal, then the signal must be small. However, Deser et al. [29] show that such a change in surface pressure patterns can alter the risk of regional drought or heavy precipitation by a factor of two, which is hardly negligible. Most climate model simulations are performed with much smaller ensemble sizes, although there is a growing interest in large single-model ensembles in order to better characterize the epistemic uncertainty within current models.

Another conceptual challenge in dealing with the epistemic uncertainty of climate change is that the concept of “error” is not well defined. Although in principle there may be one truth, it is not knowable: there will never be sufficient observations to define all relevant aspects of future climate; future climate will in any case be non-stationary; and model projections are based on climate forcing scenarios that will not be the ones actually realized. Thus, there has been interest in trying to understand the relationship between model errors in observable aspects of climate and the forced response simulated by that model — so-called “emergent constraints” (e.g. [30]). Such an approach permits a Bayesian probabilistic interpretation of epistemic uncertainty [31]. However, there is a danger that any such relationship is merely statistical and not causal, and many published emergent constraints have been subsequently debunked (see [32-34]). In any case, subjective choices are required in the application of any such constraints.

That an aleatoric interpretation of multi-model ensembles can blur the climate information contained within those ensembles is not difficult to appreciate. Circulation aspects of climate are related to features such as jet streams. Over Europe during wintertime, some models show an increase in jet strength under climate change and others a decrease (see Figure 4 of [1]),

moreover the location of the changes varies between models. Whilst all models predict a significant jet response somewhere, averaging over the models will lead to a washed-out response. Thus the multi-model mean may not only be unlikely, but even implausible. The situation is analogous to the idealized case of a bi-modal Probability Density Function, whose mean may not be a physically realizable state.

A related issue is apparent in Figure 1. Because precipitation increases in some regions and decreases in others, the multi-model mean change inevitably passes through zero, and will be small compared to internal variability on either side of that line. However, that does not mean that the change in those regions can be expected to be small compared to internal variability; it just reflects uncertainty in the sign of the change. When there are equally plausible futures that point in different directions, averaging those futures buries relevant information and underestimates risk.

The essential point is that epistemic uncertainties are deterministic, which means that they introduce correlations; unless those correlations are accounted for, inferences may be flawed. For example, Madsen et al. [35] show that the spread across CMIP5 model projections in temperature and precipitation changes at the gridpoint scale is significantly exaggerated when treating the gridpoints independently, as compared to when the models are ranked by the global mean changes (where the spread comes mainly from climate sensitivity). This illustrates the general point that with an inhomogeneous distribution of estimators, one should examine the distribution of responses to a perturbation rather than the overall response of the distribution to the perturbation.

#### **4. Re-framing the question**

If the construction of regional climate information inevitably involves ethical choices, then those choices should be made by the users of the climate information, based on their values. If the uncertainties in the climate information involve a significant epistemic component, then subjectivity is inevitable and the epistemic uncertainties similarly need to be understandable and assessable by the users of the climate information, within their particular context. Both imperatives move the climate risk problem outside the domain of pure climate science. Moreover, the recognition that epistemic uncertainties are deterministic removes the impulse to provide probabilities, which can give the illusion of objectivity and thereby reduce transparency. Instead, epistemic uncertainty can be represented through a discrete set of (multiple) storylines — physically self-consistent, plausible pathways, with no probability attached [11,36].

Rather than asking what will happen (as in the traditional, scenario-driven approach), which we may not be able to answer with any confidence, storylines allow us to ask what would be the effect of particular interventions

— e.g. different climate forcing scenarios, or different adaptation measures — across a range of plausible futures. The latter questions are in any case the societally relevant ones. This re-framing of the climate risk question from the prediction space to the decision space avoids the confidence straightjacket. Storylines have much in common with scenario planning and other methods of robust decision-making under uncertainty [37,38]. What is novel is their application to physical climate science, where, perhaps because the system obeys known physical laws, the operative paradigm up to now has been probabilistic, which gives the impression of objectivity.

The different uncertainties that are relevant to climate risk, and the different human decision points, can be broadly represented as follows. There is uncertainty in the future climate forcing, which is mainly anthropogenic in origin, and represents the mitigation options. This combines with the epistemic uncertainty in climate sensitivity to determine the global-mean warming level. Whilst there is some aleatoric uncertainty in the global-mean warming, it is small compared with the forced response on decadal or longer timescales. A given global-mean warming level will be associated with distinct patterns of regional warming (e.g. land warms more than ocean, the Arctic warms more than lower latitudes during the winter season), including changes in lapse rate [39]. These regional warming patterns are largely explainable from thermodynamic principles and thus are fairly well understood, though have substantial quantitative epistemic uncertainty (including the possibility of tipping points). A given global-mean warming level will also be associated with particular dynamical conditions in any specific region (including the circulation effects of coupled atmosphere-ocean variability), which have a very large aleatoric component but whose forced changes are also highly uncertain. The regional warming patterns and dynamical conditions together produce hazards such as weather or climate extremes, which then combine with the non-climatic anthropogenic factors of vulnerability and exposure to create climate impacts.

This representation of the climate risk problem provides a natural framework for storyline approaches. For example, from the perspective of the Paris Agreement, one may ask the question of what the climate impacts would be at different levels of global-mean warming, and what different mitigation pathways would lead to those warming levels [40]. The epistemic uncertainty in climate sensitivity now no longer affects the estimation of climate impacts, but is instead relevant to the carbon budget allowed by the given level of warming. The epistemic uncertainty in future dynamical conditions (for a given level of global-mean warming) can then be managed via storylines, the simplest of which is that the changes in hazard are dominated by the thermodynamic effects arising from the regional temperature changes, with the forced changes in dynamical conditions assumed to be negligible. Given the large uncertainties in the forced dynamical changes, this can be considered a reasonable null hypothesis for climate change [41,42], and it is

far from uninformative. It is in fact the basis for all of the predicted changes in extremes shown in Table SPM.1 of the IPCC AR5 [2]. It also underlies the “surrogate climate change” (also known as “pseudo-global warming”) methodology [43,39] which is widely used in regional climate change simulations, and the circulation-analogue methodology [44] which is widely used in extreme-event attribution. However, specific storylines of forced circulation change can also be considered [42,16].

Reframing the climate risk question in this way increases the signal-to-noise ratio of the climate information by explicitly accounting for the correlated nature of epistemic uncertainty. An example is provided by Figure 2. The Mediterranean region receives most of its precipitation during the winter season, so the predicted wintertime drying of the region, which is a robust feature of climate model projections (see Figure 1), has important consequences. The extent of the drying will depend on the global-warming level, and it is relevant to ask, for instance, what would be the difference between 1.5C and 2.0C of global warming. However the extent of the drying will also depend on the pattern of circulation change in the region — an epistemic uncertainty — which can be characterized by physically coherent storylines [16]. Considering just the range between the low-impact and high-impact storylines shown in Figure 2, the difference in drying between 1.5C and 2.0C of global-mean warming under the standard probabilistic framing is the difference between 0.09 [0.04 to 0.15] and 0.12 [0.05 to 0.20] mm/day (left panel), which would be considered indeterminate within the stated uncertainties. The storyline framing of the difference is, in contrast, a deterministic 0.04 vs 0.05 for the low-impact circulation storyline, 0.09 vs 0.12 for the median storyline, and 0.15 vs 0.20 for the high-impact storyline (right panel). This is a more informative way of representing the uncertainty, because it quantifies different plausible outcomes. For reference, 0.08 mm/day corresponds to a change that is statistically detectable, and 0.19 to one standard deviation of the interannual variability — quite a large change, likely requiring significant adaptation measures. The distinction between the two approaches is analogous to that between accuracy and precision; sometimes, the latter is all that is needed for decision-making.

Storylines are ideal vehicles for quantifying the impacts of climate change and adaptation measures. They provide a way of dealing with singular historical events, which within the probabilistic framework are merely accidents within a phase space of unrealized possibilities, yet often provide benchmarks for resilience; and with the local context, where the human element becomes part of the analysis rather than a confounding factor. For example, rather than seeking to determine the recurrence likelihood of a particularly damaging storm (an inherently fuzzy question since every storm is unique), one can ask how much worse the flooding would be in a warmer, moister climate [41], or under a particular urban development scenario. Such conditioning of the question enormously reduces the dimension of the problem and thereby

allows the use of much more realistic modelling tools, which users of climate information can relate to. In this way, the storyline approach addresses the needed re-framing of the climate risk problem whilst representing the epistemic uncertainties in a traceable manner.

That there is relevant information concerning climate risk contained even in single historical events is illustrated by Figure 3, which shows a small region in central France during one day in August 2000 and another day in August 2003 during the severe heat wave that affected Europe that summer [45]. From a statistical perspective, it may seem meaningless to compare two single days because they will each be strongly influenced by synoptic variability. However, the images show that the crops and grasses in the agricultural plots died out during the 2003 heat wave, and the surface temperature difference between the two days over those parts of the scene was 20 C, vs only 11 C in the forested region. Since a difference of 9 C over a distance of several hundred metres cannot be explained by synoptic variability (which has much larger correlation scales), this clearly shows the impact of land cover on the climate risk from heat waves. (Moreover, the average temperature difference in the agricultural plots rises to 24 C if the hedgerows are excluded, and the temperature difference in fields that were bare in both 2000 and 2003 is 11 C.) Whilst it may not be possible to predict the future statistics of heat waves in this region, it is possible to make informative statements about how those heat waves would be affected by land cover and thus inform adaptation strategies.

The tension between global and local descriptions (in time or space) is not unique to climate science, of course. It arises in any scientific context where statistical power is achieved by aggregation over an inhomogeneous population, and thus blurs information. There is a growing move in many fields towards analysis methods that aim to consider information in context rather than in aggregate, especially when that information is sparse (e.g. safety in health care: [46]). Storyline approaches to climate risk can be seen as part of that movement.

## 5. Causal networks

In the above, storylines have been presented in contrast to probabilistic representations of uncertainty. However, if storylines are to provide an alternative scientific paradigm for the construction of regional climate change information, they must be somehow reconcilable with the conventional, probabilistic approach, in order to effectively bridge between climate science and climate impacts, and from the global to the local scale.

The narrative description of the regional climate risk problem in the previous section is represented graphically in Figure 4. Figure 4 is a *directed acyclic graph*, which means that the climate risk problem can be represented mathematically as a causal network [47,48]. This observation provides the

key to reconciling storyline and probabilistic approaches. Following [48], a joint probability of  $n$  variables  $P(x_1, \dots, x_n)$  can be expressed as the product of conditional probabilities  $P(x_j | pa_j)$ , where  $pa_j$  are the ‘parent’ factors influencing  $x_j$ , according to

$$P(x_1, \dots, x_n) = \prod_j P(x_j | pa_j). \quad (1)$$

The representation (1) factorizes the uncertainty, which is extremely useful when the different uncertainties have rather different characteristics, as in the climate risk problem. A storyline  $x_i = x'_i$  for a particular  $i$  can be defined by imposing that particular condition within (1), represented symbolically by  $\hat{x}'_i$ , which leads to [48, pp. 72-73]

$$P(x_1, \dots, x_n | \hat{x}'_i) = \begin{cases} \prod_{j \neq i} P(x_j | pa_j) = \frac{P(x_1, \dots, x_n)}{P(x'_i | pa_i)} & \text{if } x_i = x'_i \\ 0 & \text{if } x_i \neq x'_i \end{cases}. \quad (2)$$

The expression (2) is thus a *truncated factorization* of the expression (1) for the unconditional probability, representing a blend of probabilistic and deterministic factors. Multivariate storylines can be treated by repeated application of this procedure. In this way, storylines can be cast within the context of a probabilistic framework.

We illustrate this for the system represented in Figure 4. The traditional scenario-driven prediction problem aims to estimate the joint probability of the climate state conditional only on the climate forcing  $F$ :

$$P(H, D, R, G, S | F). \quad (3)$$

According to the causal linkages represented in Figure 4, this factorizes to

$$P(H | D, R) P(D | G) P(R | G) P(G | S, F) P(S). \quad (4)$$

Within this perspective, it is necessary to have knowledge of the climate sensitivity  $S$ . However, from the perspective of the Paris Agreement, one can define a storyline consisting of a particular global warming level, say  $G = G_1$ , which specifies  $G$  deterministically. This condition blocks the influence of  $S$ , leaving the truncated factorization

$$P(H | D, R) P(D | G = G_1) P(R | G = G_1) \quad (5)$$

where now the hazard  $H$  depends only on the dynamical conditions  $D$  and the regional warming  $R$ . Note that (5) does not imply that  $D$  and  $R$  are independent; they share a common dependence through  $G$ , hence storylines of  $R$  may be correlated with storylines of  $D$ . This is precisely the basis of the approach of [16].

Interestingly, imposing a global-mean warming target builds in a relationship between the climate sensitivity  $S$  and the climate forcing  $F$ . This is in contrast

to the traditional scenario-driven formulation of climate risk, where these quantities are treated as independent. [In (3),  $S$ , as a property of the climate system, would be assumed independent of  $F$ .] Such a relationship expresses the policy-relevant information that society will need to act more aggressively on controlling emissions if climate sensitivity turns out to be high, but may allow itself more time if climate sensitivity turns out to be low.

If  $R$  is taken to be a deterministic function of  $G$ , i.e. the uncertainty in  $R$  is considered to be mainly epistemic, then (5) simplifies to

$$P(H \mid D, R = R_1) P(D \mid G = G_1) , \quad (6)$$

where  $R_1 = R(G_1)$ . The first term in (6) represents the thermodynamic effects of a particular regional warming  $R_1$  on  $H$ , given knowledge of  $D$ , whilst the second term represents the dynamical effects of climate change. As already discussed, the epistemic uncertainty in the latter can be very high, but is representable through storylines. The simplest storyline is that dynamics remains unchanged, in which case the conditionality in the second term drops out and we are left with

$$P(H \mid D, R = R_1) P(D) , \quad (7)$$

where  $P(D)$  can be based, for example, on observations. This is exactly the formulation of the “surrogate climate change” methodology mentioned earlier, which is widely used in regional climate change simulations. However, one can certainly also specify different dynamical storylines to represent plausible changes in dynamics. Since (6) essentially describes the regional climate modelling paradigm, it may provide a useful framework for the construction of regional climate-change information and the design of ensembles of simulations using regional climate models, including the representation of particularly extreme forms of internal variability.

Without this factorization of the probabilities, the regional climate risk problem for a given global warming level is representable instead in the form

$$P(H, D, R \mid G) , \quad (8)$$

which lends itself to a probabilistic interpretation of the dynamical aspects of climate change. This hides the implicit assumptions concerning the epistemic uncertainties that are made explicit in the representation (6). Moreover, the comparatively unconditional nature of (8) requires the use of global models, whereas (6) permits the use of regional models, which can provide a more physically realistic representation of regional climate risk [49-51].

By casting storylines within the context of a probabilistic framework, it becomes clear that there is nothing to prevent assigning probabilities to storylines, if the scientific basis exists to do that. At the very least, physically implausible behaviours could be excluded [50]. As epistemic uncertainties are

reduced, this knowledge can be immediately incorporated into a revised risk analysis. Thus, storylines provide a very flexible, transparent representation of epistemic uncertainty.

Not only do causal networks reconcile storyline and probabilistic approaches to climate risk, they are also ideally suited for moving the risk question into the decision space. That is because the calculus of causal networks explicitly allows the consideration of counter-factual outcomes [48], and decision-making is precisely the consideration of counter-factual outcomes. Within this context, storylines correspond to what Halpern and Pearl [52] define as explanations: “a fact that is not known for certain but, if found to be true, would constitute an actual cause of the fact to be explained, regardless of the agent’s initial uncertainty”.

More generally, causal networks are a way of combining expert knowledge with probability [47]. The factorization (1) allows for the ready incorporation of knowledge within a local semantics, and yields results that are comprehensible to humans [53]. In the published Discussion of Lauritzen and Spiegelhalter [47, p. 210] J. Pearl invokes the following statement (attributed to G. Halter): “Probability is not really about numbers; it is about the structure of reasoning.” Making the subjective assumptions explicit leads to transparency in the subsequent analysis [54] and provides an audit trail for decision-makers [55]. This is important since, as Beven [55] puts it, “Decision and policy makers are ... far more interested in evidence than uncertainty.”

The challenge for regional climate-change science then becomes that of constructing suitable causal networks. Causal networks are necessarily a simplification, because they entail the reduction of continuous fields to a finite-dimensional system. However, they very much correspond to how climate scientists reason. For example, the El Niño variability in tropical sea-surface temperatures drives a Rossby-wave teleconnection pathway which affects circulation and weather regimes in the mid-latitudes, and all these elements can be represented to a reasonable extent with physical climate indices. Thus, atmospheric dynamics already provides the building blocks for the construction of causal networks relevant to regional climate risk. (In practice, the “Dynamical conditions” node in Figure 4 could be expanded into a sub-network.) Comprehensive climate simulation models are still needed to explore uncertainty space, but causal networks can provide the diagnostic framework within which to extract the relevant climate information from those simulations, and combine it with other sources of information in a format that is suitable for decision-making.

The causal network depicted in Figure 4 incorporates two emergent aspects of climate change. Both aspects are simplifications, but they are extremely powerful and are widely used in the interpretation of climate information. The first is what is known as “pattern scaling” [56,57]: namely that regional climate change is a function of global-mean warming. In practice, the patterns



of regional warming are time-dependent [58] so are different for transient and equilibrated warming levels, and short-lived climate forcers such as aerosol can have distinct regional effects [59]. Such additional degrees of freedom, as well as global tipping points, could be incorporated by making the node  $G$  suitably multivariate. The second emergent aspect is the distinction between thermodynamic and dynamical aspects of regional climate change, which has already been discussed. Whilst the distinction is not precise and has its limitations, it is useful (e.g. [60]); it has even been used for the last two Dutch Climate Change Scenarios [61]. As with the other simplifications implicit in Figure 4, e.g. the lack of any arrows pointing back from the right to the left, the validity of all these simplifications can be assessed *a posteriori*.

Note that linearity is not assumed in causal networks. However, if certain relationships can be shown to be linear to a suitable level of approximation for the problem at hand, then the analysis is enormously simplified. This is generally necessary for any observational analysis, because of the limited sample size [62].

## 6. Discussion

This paper has argued that the storyline approach to regional climate-change information avoids the straightjacket that hampers the standard confidence-based approach, by allowing a reframing of the climate risk question from the prediction space into the decision space. Whilst in principle such a reframing is possible from probabilistic estimates of risk, the challenge for regional climate-change information is that the level of epistemic uncertainty is sufficiently high that subjective choices must inevitably be made, and the range of users sufficiently inhomogeneous that there is no consensus on values. Under such conditions, probabilistic ‘rational-choice’ approaches to decision-making are ineffective [63,64] and the decision framework needs to be one where the subjective and ethical choices are both flexible and transparent [65]. Since epistemic uncertainty is inherently deterministic and subjective, there is no imperative to represent it probabilistically [23], and probabilistic representations can give a false impression of objectivity.

The reframing of the risk question from the prediction space to the decision space may seem uncomfortable from a physical science perspective, but is in fact quite orthodox from the perspective of statistical inference. Despite the widespread use of p-values as an ostensibly objective measure of statistical significance, the inference derived from data concerning a particular hypothesis is far from a straightforward matter and involves many assumptions [66]. In the Neyman-Pearson framework, the inference problem is regularized by placing it in a decision context between two alternative hypotheses, which takes into account the possibility of both Type 1 and Type 2 errors [67]. In the Bayesian framework, the strength of evidence between these alternative hypotheses ( $H_1$  and  $H_2$ ) provided by the data  $D$  is given by

$$\frac{P(H_2 | D)}{P(H_1 | D)} = \frac{P(D | H_2) P(H_2)}{P(D | H_1) P(H_1)}, \quad (9)$$

which follows directly from Bayes' theorem. The Bayes factor  $P(D|H_2)/P(D|H_1)$  is independent of the prior likelihoods  $P(H_2)$  and  $P(H_1)$ , so can be considered objective, but it does not represent any sort of absolute knowledge — only an increment in knowledge, relative to the prior beliefs.

Moving the climate risk problem out of the domain of pure climate science requires humility on the part of climate scientists. To quote Funtowicz and Ravetz [63] — who used sea-level rise as an example — “the traditional domination of ‘hard facts’ over ‘soft values’ [is] inverted... traditional scientific inputs... become ‘soft’ in the context of the ‘hard’ value commitments that will determine the success of policies for mitigating the effects of [climate change]”. Indeed, it has been argued that humility is one of the four core elements — the others being integrity, transparency, and collaboration — that should be intrinsic to the production of regional climate information [68]. In this way, the goal is not so much to be authoritative, which has something of a gatekeeper connotation, but to be trustworthy [69]. This involves a loss of control, because one's trustworthiness is a judgement made by others.

This perspective also involves an acknowledgement that climate-relevant decisions, especially at the local scale, are not usually made on the basis of climate change alone but involve many other changing factors, most of which are highly uncertain. If climate impacts  $I$  are a product of hazard  $H$ , vulnerability  $V$  and exposure  $E$ , then, conceptually, the anthropogenic changes in  $I$  can be represented as

$$\delta I = \delta(HVE) = HV\delta E + HE\delta V + VE\delta H. \quad (10)$$

It may well be that the largest terms on the right-hand side of (10) are the first two, where it is the combination of climate and weather *variability* with changing vulnerability and exposure that is the main determinant of climate risk [70]. In this case the decision framework is not so much that of dealing with climate change as it is that of bringing climate information into decisions that need to be made in any case. There are calls for this sort of complex-systems thinking in other areas of science, such as public health [71]: “Instead of asking whether an intervention works to fix a problem, researchers should aim to identify if and how it contributes to reshaping a system in favourable ways.”

To return to Kuhn [22], the construction of regional climate-change information is not most usefully viewed as a search for an objective truth, but rather as a search for more complete descriptions of the realities that people have experienced and may experience in the future, and how those depend on contingent factors that are under human control. Kuhn's version of the

665 Bayesian perspective described above, and the cutting of the Gordian Knot it  
666 enables, is as follows [22, p. 170]: “If we can learn to substitute evolution-  
667 from-what-we-know for evolution-toward-what-we-wish-to-know, a number  
668 of vexing problems may vanish in the process.” In such an enterprise, physical  
669 knowledge of the climate system provides the foundation for the construction  
670 of regional climate information.

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## References

1. Shepherd TG. 2014 Atmospheric circulation as a source of uncertainty in climate change projections. *Nature Geosci.* **7**, 703–708. (doi:10.1038/NGEO2253)
2. IPCC. 2014a *Climate Change 2013: The Physical Basis*. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change (Stocker TF, et al., eds.). Cambridge, UK: Cambridge University Press.
3. Fischer EM, Beyerle U, Knutti R. 2013 Robust spatially aggregated projections of climate extremes. *Nature Clim. Change* **3**, 1033–1038. (doi:10.1038/NCLIMATE2051)
4. Hoskins B, Woollings T. 2015 Persistent extratropical regimes and climate extremes. *Curr. Clim. Change Rep.* **1**, 115–124. (doi:10.1007/s40641-015-0020-8)
5. Shaw TA, et al. 2016 Storm track processes and the opposing influences of climate change. *Nature Geosci.* **9**, 656–664. (doi:10.1038/NGEO2783)
6. Deser C, Terray L, Phillips AS. 2016 Forced and internal components of winter air temperature trends over North America during the past 50 years: Mechanisms and implications. *J. Clim.* **29**, 2237–2258. (doi:10.1175/JCLI-D-15-0304.1)
7. Bony S, Bellon G, Klocke D, Sherwood S, Fermepin S, Denvil S. 2013 Robust direct effect of carbon dioxide on tropical circulation and regional precipitation. *Nature Geosci.* **6**, 447–451. (doi:10.1038/NGEO1799)
8. Simpson IR, Seager R, Ting M, Shaw TA. 2016 Causes of change in Northern Hemisphere winter meridional winds and regional hydroclimate. *Nature Clim. Change* **6**, 65–70. (doi:10.1038/nclimate2783)
9. van Niekerk A, Scinocca JF, Shepherd TG. 2017 The modulation of stationary waves, and their response to climate change, by parameterized orographic drag. *J. Atmos. Sci.* **74**, 2557–2574. (doi:10.1175/JAS-D-17-0085.1)
10. Held IM, Soden BJ. 2006 Robust responses of the hydrological cycle to global warming. *J. Clim.* **19**, 5686–5699. (doi:10.1175/JCLI3990.1)
11. Shepherd TG, et al. 2018 Storylines: an alternative approach to representing uncertainty in physical aspects of climate change. *Climatic Change* **151**, 555–571. (doi:10.1007/s10584-018-2317-9)

- 710 12. Scheff J, Frierson D. 2012 Twenty-first-century multimodel subtropical  
711 precipitation declines are mostly midlatitude shifts. *J. Clim.* **25**, 4330–4347.  
712 (doi:10.1175/JCLI-D-11-00393.1)
- 713 13. Chadwick R, Boutle I, Martin G. 2013 Spatial patterns of precipitation  
714 change in CMIP5: Why the rich do not get richer in the tropics. *J. Clim.* **27**,  
715 3803–3822. (doi:10.1175/JCLI-D-12-00543.1)
- 716 14. Byrne MP, O’Gorman PA. 2015 The response of precipitation minus  
717 evapotranspiration to climate warming: Why the “wet-get-wetter, dry-get-  
718 drier” scaling does not hold over land. *J. Clim.* **28**, 8078–8092.  
719 (doi:10.1175/JCLI-D-15-0369.1)
- 720 15. Zappa G, Shaffrey LC, Hodges KI, Sansom PG, Stephenson DB. 2013 A  
721 multimodel assessment of future projections of North Atlantic and  
722 European extratropical cyclones in the CMIP5 climate models. *J. Clim.* **26**,  
723 5846–5862. (doi:10.1175/JCLI-D-12-00573.1)
- 724 16. Zappa G, Shepherd TG. 2017 Storylines of atmospheric circulation change  
725 for European regional climate impact assessment. *J. Clim.* **30**, 6561–6577.  
726 (doi:10.1175/JCLI-D-16-0807.1)
- 727 17. IPCC. 2014b *Climate Change 2014: Impacts, Adaptation, and Vulnerability*.  
728 Contribution of Working Group II to the Fifth Assessment Report of the  
729 Intergovernmental Panel on Climate Change (Field CB, et al., eds.).  
730 Cambridge, UK: Cambridge University Press.
- 731 18. Coughlan de Perez E, Monasso F, van Aalst M, Suarez P. 2014 Science to  
732 prevent disasters. *Nature Geosci.* **7**, 78–79. (doi:10.1038/ngeo2081)
- 733 19. Yaniv I, Foster DP. 1995 Graininess of judgment under uncertainty: An  
734 accuracy-informativeness trade-off. *J. Exp. Psych.: Gen.* **124**, 424–432.  
735 (doi:10.1037/0096-3445.124.4.424)
- 736 20. Lloyd EA, Oreskes N. 2018 Climate change attribution: When is it  
737 appropriate to accept new methods? *Earth’s Future* **6**, 311–325.  
738 (doi:10.1002/2017EF000665)
- 739 21. Beven K. 2011 I believe in climate change but how precautionary do we  
740 need to be in planning for the future? *Hydrol. Process.* **25**, 1517–1520.  
741 (doi:10.1002/hyp.7939)
- 742 22. Kuhn TS. 2012 *The Structure of Scientific Revolutions*, 50<sup>th</sup> anniversary  
743 edition. Chicago, USA: The University of Chicago Press.
- 744 23. Dessai S, Hulme M. 2004 Does climate adaptation policy need  
745 probabilities? *Climate Policy* **4**, 107–128.  
746 (doi:10.1080/14693062.2004.9685515)

- 747 24. Hawkins E, Sutton R. 2011 The potential to narrow uncertainty in  
748 projections of regional precipitation change. *Clim. Dyn.* **37**, 407–418.  
749 (doi:10.1007/s00382-010-0810-6)
- 750 25. Tebaldi C, Knutti R. 2007 The use of the multi-model ensemble in  
751 probabilistic climate projections. *Phil. Trans. R. Soc. A* **365**, 2053–2075.  
752 (doi:10.1098/rsta.2007.2076)
- 753 26. Kahneman D, Tversky A. 1982 Variants of uncertainty. *Cognition* **11**, 143–  
754 157. (doi:10.1016/0010-0277(82)90023-3)
- 755 27. Smith LA. 2002 What might we learn from climate forecasts? *Proc. Natl.*  
756 *Acad. Sci. USA* **99**, 2487–2492. (doi:10.1073/pnas.012580599)
- 757 28. Oppenheimer M, O’Neill BC, Webster M, Agrawala S. 2007 The limits of  
758 consensus. *Science* **317**, 1505–1506. (doi:10.1126/science.1144831)
- 759 29. Deser C, Phillips A, Bourdette V, Teng HY. 2012 Uncertainty in climate  
760 change projections: the role of internal variability. *Clim. Dyn.* **38**, 527–546.  
761 (doi:10.1007/s00382-010-0977-x)
- 762 30. Hall A, Qu X. 2006. Using the current seasonal cycle to constrain snow  
763 albedo feedback in future climate change. *Geophys. Res. Lett.* **33**, L03502.  
764 (doi:10.1029/2005GL025127)
- 765 31. Sexton DMH, Murphy JM, Collins M, Webb MJ. 2012 Multivariate  
766 probabilistic projections using imperfect climate models. Part I: Outline of  
767 methodology. *Clim. Dyn.* **38**, 2513–2542. (doi:10.1007/s00382-011-1208-  
768 9)
- 769 32. Pithan F, Mauritsen T. 2013 Comments on “Current GCMs’ Unrealistic  
770 Negative Feedback in the Arctic”. *J. Clim.* **26**, 7783–7788.  
771 (doi:10.1175/JCLI-D-12-00331.1)
- 772 33. Simpson IR, Polvani L. 2016 Revisiting the relationship between jet  
773 position, forced response, and annular mode variability in the southern  
774 midlatitudes. *Geophys. Res. Lett.* **43**, 2896–2903.  
775 (doi:10.1002/2016GL067989)
- 776 34. Caldwell PM, Zelinka MD, Klein SA. 2018 Evaluating emergent constraints  
777 on equilibrium climate sensitivity. *J. Clim.* **31**, 3921–3942.  
778 (doi:10.1175/JCLI-D-17-0631.1)
- 779 35. Madsen MS, Langen PL, Boberg F, Christensen JH. 2017 Inflated  
780 uncertainty in multimodel-based regional climate projections. *Geophys. Res.*  
781 *Lett.* **44**, 11,606–11,613. (doi:10.1002/2017GL075627)

- 782 36. Hazeleger W, van den Hurk BJJM, Min E, van Oldenborgh GJ, Petersen AC,  
783 Stainforth DA, Vasileiadou E, Smith LA. 2015 Tales of future weather.  
784 *Nature Clim. Change* **5**, 107–113. (doi:10.1038/NCLIMATE2450)
- 785 37. Prudhomme C, Wilby RL, Crooks S, Kay AL, Reynard NS. 2010 Scenario-  
786 neutral approach to climate change impact studies: application to flood  
787 risk. *J. Hydrol.* **390**, 198–209. (doi:10.1016/j.jhydrol.2010.06.043)
- 788 38. Lempert R. 2013 Scenarios that illuminate vulnerabilities and robust  
789 responses. *Climatic Change* **117**, 627–646. (doi:10.1007/s10584-012-  
790 0574-6)
- 791 39. Kröner N, Kotlarski S, Fischer E, Lüthi D, Zubler E, Schär C. 2017  
792 Separating climate change signals into thermodynamic, lapse-rate and  
793 circulation effects: theory and application to the European summer climate.  
794 *Clim. Dyn.* **48**, 3425–3440. (doi:10.1007/s00382-016-3276-3)
- 795 40. IPCC. 2018 *Global warming of 1.5°C*. An IPCC Special Report on the impacts  
796 of global warming of 1.5°C above pre-industrial levels and related global  
797 greenhouse gas emission pathways, in the context of strengthening the  
798 global response to the threat of climate change, sustainable development,  
799 and efforts to eradicate poverty (Masson-Delmotte V, et al., eds.). Geneva,  
800 CH: World Meteorological Organization.
- 801 41. Trenberth KE, Fasullo JT, Shepherd TG. 2015 Attribution of climate  
802 extreme events. *Nature Clim. Change* **5**, 725–730.  
803 (doi:10.1038/NCLIMATE2657)
- 804 42. Shepherd TG. 2016 A common framework for approaches to extreme  
805 event attribution. *Curr. Clim. Change Rep.* **2**, 28–38. (doi:10.1007/s40641-  
806 016-0033-y)
- 807 43. Schär C, Frei C, Lüthi D, Davies HC. 1996 Surrogate climate-change  
808 scenarios for regional climate models. *Geophys. Res. Lett.* **23**, 669–672.  
809 (doi:10.1029/96GL00265)
- 810 44. Cattiaux J, Vautard R, Cassou C, Yiou P, Masson-Delmotte V, Codron F.  
811 2010 Winter 2010 in Europe: A cold extreme in a warming climate.  
812 *Geophys. Res. Lett.* **37**, L20704. (doi:10.1029/2010GL044613)
- 813 45. Zaitchik BF, Macalady AK, Bonneau LR, Smith RB. 2006. Europe's 2003  
814 heat wave: a satellite view of impacts and land-atmosphere feedbacks. *Int.*  
815 *J. Clim.* **26**, 743–769. (doi:10.1002/joc.1280)
- 816 46. Wears RL. 2003 Still learning how to learn. *Qual. Saf. Health Care* **12**, 471–  
817 472.

818 47. Lauritzen SL, Spiegelhalter DJ. 1988 Local computations with probabilities  
819 on graphical structures and their application to expert systems. *J. R. Stat.*  
820 *Soc. B* **50**, 157–224. (<http://www.jstor.org/stable/2345762>)

821 48. Pearl J. 2009 *Causality*, 2<sup>nd</sup> ed. Cambridge, UK: Cambridge University Press.

822 49. Hall A. 2014 Projecting regional change. *Science* **346**, 1461–1462.  
823 (doi:10.1126/science.aaa0629)

824 50. Bukovsky MS, McCrary RR, Seth A, Mearns LO. 2017 A mechanistically  
825 credible, poleward shift in warm-season precipitation projected for the U.S.  
826 Southern Great Plains? *J. Clim.* **30**, 8275–8298. (doi:10.1175/JCLI-D-16-  
827 0316.1)

828 51. Maraun D, et al. 2017 Towards process-informed bias correction of  
829 climate change simulations. *Nature Clim. Change* **7**, 764–773.  
830 (doi:10.1038/NCLIMATE3418)

831 52. Halpern JY, Pearl J. 2005 Causes and explanations: A structural-model  
832 approach. Part II: Explanations. *Brit. J. Phil. Sci.* **56**, 889–911.  
833 (doi:10.1093/bjps/axi148)

834 53. Binder J, Koller D, Russell S, Kanazawa K. Adaptive probabilistic networks  
835 with hidden variables. *Machine Learning* **29**, 213–244. (doi:  
836 10.1023/A:1007421730016)

837 54. Chandler RE. 2013 Exploiting strength, discounting weakness: combining  
838 information from multiple climate simulators. *Phil. Trans. R. Soc. A* **371**,  
839 20120388. (doi:10.1098/rsta.2012.0388)

840 55. Beven K. 2016 Facets of uncertainty: epistemic uncertainty, non-  
841 stationarity, likelihood, hypothesis testing, and communication. *Hydrol. Sci.*  
842 *J.* **61**, 1652–1665. (doi: 10.1080/02626667.2015.1031761)

843 56. Mitchell TD. 2003 Pattern scaling: An examination of the accuracy of the  
844 technique for describing future climates. *Climatic Change* **60**, 217–242.  
845 (doi:10.1023/A:1026035305597)

846 57. Tebaldi C, Arblaster J. 2014 Pattern scaling: Its strengths and limitations,  
847 and an update on the latest model simulations. *Climatic Change* **122**, 459–  
848 471. (doi:10.1007/s10584-013-1032-9)

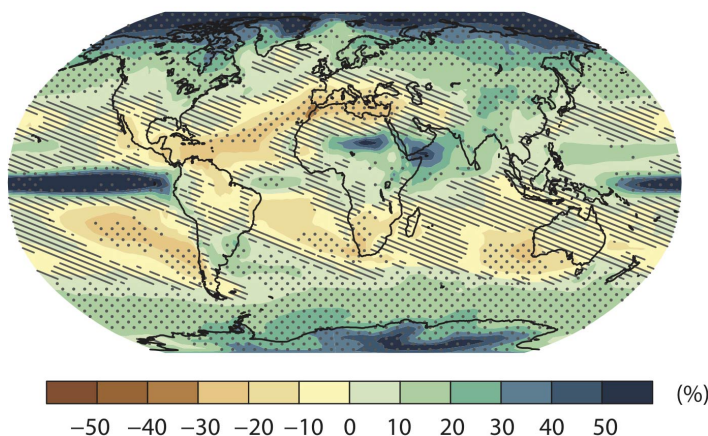
849 58. Ceppi P, Zappa G, Shepherd TG, Gregory, JM. 2018 Fast and slow  
850 components of the extratropical atmospheric circulation response to CO<sub>2</sub>  
851 forcing. *J. Clim.* **31**, 1091–1105. (doi:10.1175/JCLI-D-17-0323.1)

852 59. Ming Y, Ramaswamy V, Chen G. 2011 A model investigation of aerosol-  
853 induced changes in boreal winter extratropical circulation. *J. Clim.* **24**,  
854 6077–6091. (doi:10.1175/2011JCLI4111.1)

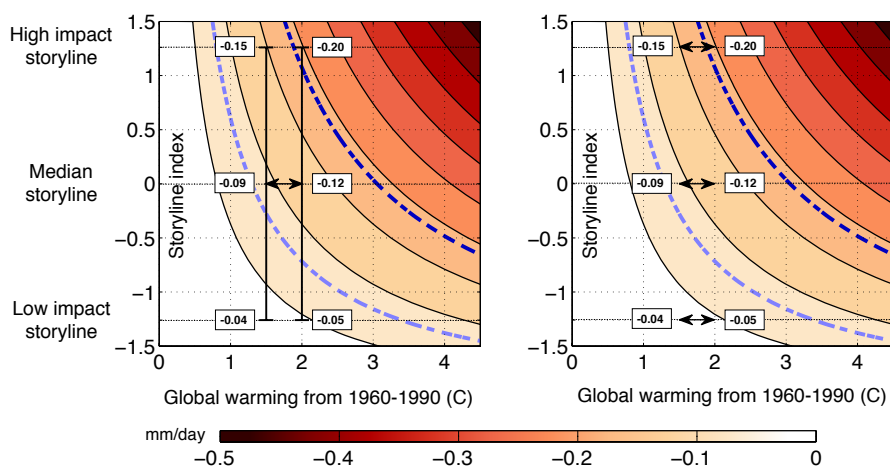


60. Pfahl S, O’Gorman PA, Fischer EM 2017 Understanding the regional pattern of projected future changes in extreme precipitation. *Nature Clim. Change* **7**, 423–427. (doi:10.1038/nclimate3287)
61. van den Hurk B, et al. 2014 KNMI’14: Climate change scenarios for the 21st century—a Netherlands perspective. Scientific Report WR2014-01, KNMI, De Bilt, the Netherlands. (<http://www.climatescenarios.nl/>)
62. Kretschmer M, Coumou D, Donges JF, Runge J. 2016 Using causal effect networks to analyze different Arctic drivers of midlatitude winter circulation. *J. Clim.* **29**, 4069–4081. (doi: 10.1175/JCLI-D-15-0654.1)
63. Funtowicz SO, Ravetz JR. 1993 Science for the post-normal age. *Futures* **25**, 739–755. (doi:10.1016/0016-3287(93)90022-L)
64. French S, Argyris N. 2018 Decision analysis and political processes. *Decision Analysis* **15**, 208–222. (doi:10.1287/deca.2018.0374)
65. van der Sluijs JP, Craye M, Funtowicz S, Klopogge P, Ravetz J, Risbey J. 2005 Combining quantitative and qualitative measures of uncertainty in model-based environmental assessment: The NUSAP system. *Risk Analysis* **25**, 481–492. (doi:10.1111/j.1539-6924.2005.00604.x)
66. Nuzzo R. 2014 Statistical errors. *Nature* **506**, 150–152.
67. Gigerenzer G. 2004 Mindless statistics. *J. Socio-Economics* **33**, 587–606. (doi:10.1016/j.socec.2004.09.033)
68. Adams P, Eitland E, Hewitson B, Vaughan C, Wilby R, Zebiak S. 2015 *Toward an ethical framework for climate services*. A White Paper of the Climate Services Partnership Working Group on Climate Services Ethics. Available from [www.climate-services.org](http://www.climate-services.org)
69. O’Neill O. 2002 *A Question of Trust*. The BBC Reith Lectures 2002. Cambridge, UK: Cambridge University Press.
70. Nissan H, Goddard L, Coughlan de Perez E, Furlow J, Baethgen W, Thomson MC, Mason SJ. 2019 On the use and misuse of climate change projections in international development. *WIREs Clim. Change* **10**, e579. (doi:10.1002/wcc.579)
71. Rutter H, et al. 2017 The need for a complex systems model of evidence for public health. *Lancet* **390**, 2602–2604. (doi:10.1016/S0140-6736(17)31267-9)

## Change in precipitation

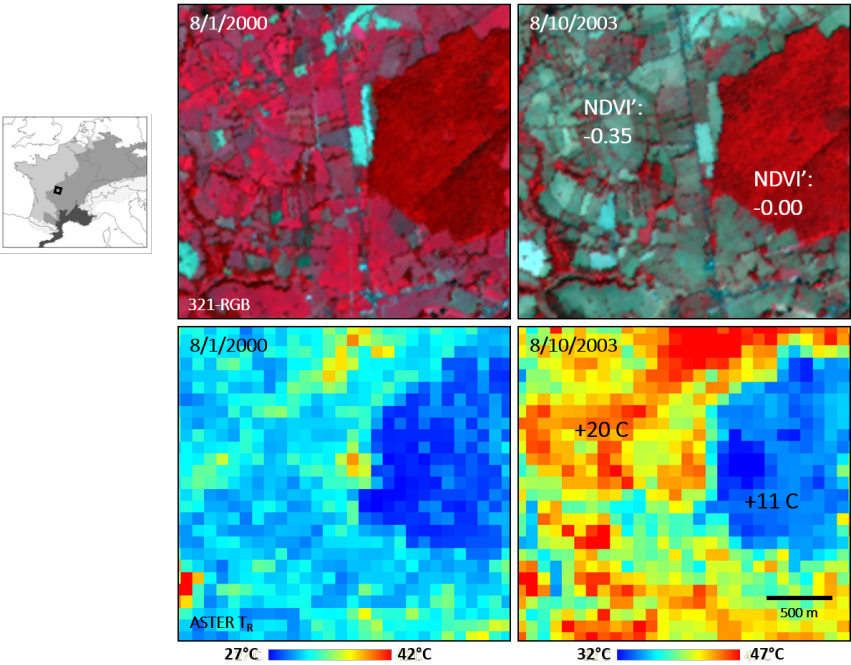


**Figure 1.** Projected changes in precipitation (in %) over the 21<sup>st</sup> century under a high climate forcing scenario (RCP8.5). Stippling indicates where the multi-model mean change is large compared with natural internal variability in 20-year means (greater than two standard deviations) and where at least 90% of models agree on the sign of change. Hatching indicates where the multi-model mean change is small compared with internal variability (less than one standard deviation), but this does not mean that individual model changes are small. From the Summary for Policymakers of [2].

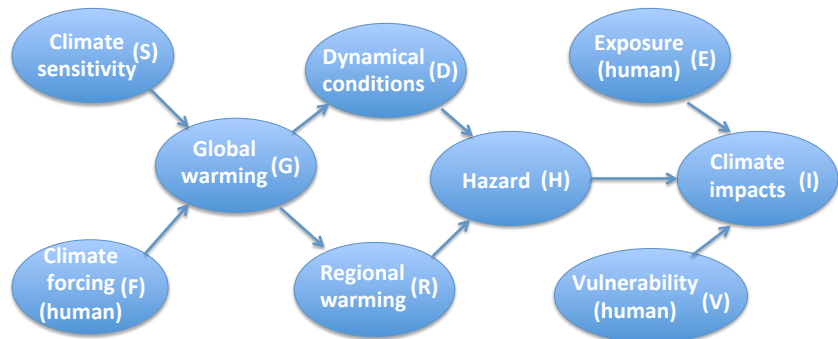


**Figure 2.** Projected average wintertime precipitation change (in mm/day) over the Mediterranean basin plotted as a function of global warming level (in C) and a 'storyline index' that represents the uncertainty in the pattern of circulation change in the region. The high impact storyline corresponds to the combination of strong tropical upper tropospheric amplification of surface warming and a strengthening of the stratospheric polar vortex, and the low impact storyline to weak tropical upper tropospheric amplification of surface warming and a weakening of the polar vortex. The light blue dashed line represents a magnitude of change that is statistically detectable, and the dark blue dashed line to one standard deviation of the interannual variability. In

the left panel, the standard representation of the difference between global warming levels of 1.5 C and 2.0 C is shown, taking the low and high impact storylines as spanning a range of uncertainty. In the right panel, differences are shown conditioned on different storylines. Adapted from [16].



**Figure 3.** Surface conditions derived from infrared remote sensing for a small region in central France, for 1 August 2000 (left panels) and 10 August 2003 (right panels). The top panels show the normalized difference vegetation index (NDVI), with the red colours indicative of vegetation. The lower panels show the radiometric temperature, with the colour scale at the bottom. The distance scale is shown in the lower-right panel, and the values given in the right panels indicate the average differences in those parts of the scene between the left and right panels. Adapted from [45].



**Figure 4.** A causal network describing regional climate risk. The arrows indicate the directions of causal influence. See text for details.