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The influence of the atmospheric boundary layer on nocturnal layers of noctuids and other moths migrating over southern Britain

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Abstract. Insects migrating at high altitude over southern Britain have been continuously monitored by automatically-operating, vertical-looking radars over a period of several years. During some occasions in the summer months, the migrants were observed to form well-defined layer concentrations, typically at heights of 200-400 m, in the stable night-time atmosphere. Under these conditions, insects are likely to have control over their vertical movements and are selecting flight heights which are favourable for long-range migration. We therefore investigated the factors influencing the formation of these insect layers by comparing radar measurements of the vertical distribution of insect density with meteorological profiles generated by the UK Met. Office's Unified Model (UM). Radar-derived measurements of mass and displacement speed, along with data from Rothamsted Insect Survey light traps provided information on the identity of the migrants. We present here three case studies where noctuid and pyralid moths contributed substantially to the observed layers. The major meteorological factors influencing the layer concentrations appeared to be: (a) the altitude of the warmest air, (b) heights corresponding to temperature preferences or thresholds for sustained migration and (c), on nights when air temperatures are relatively high, wind-speed maxima associated with the nocturnal jet. Back-trajectories indicated that layer duration may have been determined by the distance to the coast. Overall, the unique combination of meteorological data from the UM and insect data from entomological radar described here show considerable promise for systematic studies of high-altitude insect layering.

Keywords

Insect layering - nocturnal boundary layer - temperature inversion - noctuid moths - entomological radar

Introduction

Many insect species have a migratory phase in their life-cycle, and migrations often begin when individuals ascend out of their ‘flight boundary layer’ (FBL¹: Taylor 1974), and climb to altitudes of several hundred metres above the ground. Here insects can utilize the typically stronger winds to travel much further during a night’s flight than would have been feasible at ground level (Johnson 1969; Drake and Gatehouse 1995; Pedgley et al. 1995; Gatehouse 1997). This behaviour leads to an enormous insect ‘bioflow’ through the atmosphere, particularly in the warmer regions of the world, and during summer at higher latitudes. Apart from its intrinsic interest, the atmospheric transport of insects is worthy of study because many migrant species are serious pests of agriculture and human health (Pedgley 1982, 1993; Irwin and Thresh 1988; Drake and Gatehouse 1995), while other insects are beneficial because they are important natural enemies of pest species (Farrow 1981; Riley et al. 1987; Chapman et al 2004). In both cases, knowledge of insect movement is necessary when formulating or improving management strategies for the species concerned.

The development and use of radar in insect migration studies has made it possible to make direct quantitative observations of the high-altitude movements *whilst they are in progress*, particularly for larger insect species such as Lepidoptera (e.g. Dickison et al. 1983; Riley et al. 1983; Drake and Farrow 1985; Chen et al. 1989; Wolf et al. 1990; Feng et al. 2003, 2004). These species have appreciable self-powered flight speeds and, if they cease wing-flapping, significant fall speeds: thus larger migrants cannot be regarded as completely passive tracers of the wind. Yet radar observations show that insects can be influenced by atmospheric processes, for example, gravity currents (Schaefer 1976; Greenbank et al. 1980; Pedgley et al.

¹ FBL: The layer of air next to the ground where wind speed is lower than an insect’s flight speed and hence where an insect has control over its velocity.

1982), atmospheric waves (Drake 1985) and cellular convection (Schaefer 1976; Reid et al. 1979). One commonly observed phenomenon is the accumulation of insects into layers of broad horizontal, but of relatively restricted (~ 50–200 m) vertical extent (Drake 1984; Drake and Farrow 1988; Drake and Rochester 1994). The layering phenomenon has been observed at various times of the day and night, at altitudes ranging from a few metres above the ground (as in the visual observations of Mel'nichenko (1936) and Larsen (1949) of moths flying in an early evening temperature inversion) up to ~ 2-3 km during the day (Campistron 1975; Drake and Farrow 1985). Sometimes several (up to five) layers are present simultaneously, one above the other (Drake and Farrow 1988). Despite the frequent observations of layers, the behavioural mechanisms causing them are by no means clear and more case studies are required.

Many high-altitude insect migrations begin around dusk and these continue for varying periods through the night (and sometimes all night: Drake et al. 1981; Feng et al. 2004). Nocturnal migrations have presumably evolved because the risk of bird predation is diminished (Drake and Farrow 1988) and thermal stress can be prevented (Rainey 1974). Moths, particularly the Noctuidae (which include important agricultural pests), are likely to be important constituents of the radar-detectable nocturnal fauna, particularly in temperate environments (Drake et al. 1981; Chen et al. 1989; Feng et al. 2003, 2004). Moreover, the dynamics and vertical structure of the nocturnal atmospheric boundary layer (NBL) itself are often optimal for long-range migrations in certain insect species – particularly in the stably stratified state found in flat inland areas during fine weather. In particular, radiative cooling from the earth's surface is much more rapid than cooling of the air itself, hence temperature inversions grow from the ground upwards near dusk (Mahrt 1979); many reports indicate that insects are associated with these surface radiative inversions. The resulting static stability increase causes a decoupling of flow above and below the inversion; above the inversion height

flow can accelerate into a nocturnal jet (Thorpe and Guymer 1977). There is evidence that migrants use these low-level jets to achieve rapid horizontal transport (Drake 1985; Drake and Farrow 1988).

Until recently, most observations of nocturnal layering have been made with azimuthally-scanning X-band radars, but because these systems are typically manually operated, field studies have generally been of short duration (~ 1-2 weeks). Entomological scanning radars are now tending to be replaced by autonomously-operating vertically-pointing systems; these can provide continuously updated profiles of insect vertical distribution over whole seasons, or even years (Smith et al. 2000; Drake 2002; Chapman et al. 2003).

The key objective of this paper is to present case studies where meteorology apparently has an effect on nocturnal insect layering in the UK and, by extension, in northern Europe – an area virtually unstudied by entomological radar techniques and previously suspected to be climatically marginal for night-time migrations. Additionally, this paper highlights that the new insect-monitoring radars, combined with outputs from the constantly-developing suites of computer models used in numerical weather prediction (in this case the UK Met. Office's Unified Model), hold out the prospect of employing these tools in *systematic* studies of insect layers in the atmosphere (Drake and Rochester 1994).

Materials and methods

Radar-derived insect data

Entomological radar is the only effective method of directly observing migrating insects at high altitude (see reviews in Vaughn 1985; Drake and Farrow 1988; Reynolds and Riley 1997; Smith et al. 2000; also The Radar Entomology Website: <http://www.ph.adfa.edu.au/a->

drake/trews/). In this study, we used the recently-developed vertical-looking insect-monitoring radar (VLR) technique (Chapman et al. 2002b, 2003; Drake 2002) which, at least for macro-insects, gives instantaneous vertical profiles of insect aerial density over virtually all migration altitudes to be expected over the UK.

The radar used in the present study was located at Malvern, Worcestershire – during the August 2000 case study at lat. 52° 7' 54"N, long. 2° 19' 55"W (86 m asl) and during the August 2003 studies at a nearby site 52° 06' 04" N, 2° 18' 38" W (59 m asl). The 3.2 cm wavelength (X-band) radar beam is circularly-symmetric and zenith-pointing, and the plane of linear polarisation is continuously rotated at about 5.8 Hz. In addition, the beam nutates due to a slight offset (0.1 beam widths) in the antenna feed, producing a narrow-angle conical scan. The 1.5 m diameter parabolic antenna gives a half-power beam width of 1.4°. The pulse duration is 100 nano-seconds, and the peak pulse power is 25 kW. Data are recorded during a 5-minute sampling period, repeated every 15 minutes, 24 hours a day. Return signals from individual insect targets flying through the radar beam are detected in fifteen range-gates (sampling volumes), each 45 m deep with 26 m non-sampled intervals to give coverage from 180-1218 m. The system routinely extracts the target's distance of closest approach to the beam's central axis, horizontal speed, displacement direction, body alignment, and three terms that describe the radar scattering properties of the target. Using laboratory measurements of radar cross-sections of insects, the back scattering terms can be routinely employed to estimate the target's mass and shape (Chapman et al. 2002b). All the radar-derived variables can then be used to infer the target's identity, and these inferences were further supported by ground trap data or, occasionally, by aerial sampling. The analysis program also routinely records the percentage of time the received signal power is above certain power (threshold) levels, particularly the -80 dBm (10^{-11} Watts) level which is ~ 10 dB over the noise floor of the radar receiver. These 'percentage above threshold' values provide a measure of the biomass of

insects flying: useful in situations where aerial-densities are too high for individual targets to be resolved by the radar. Further details of the radar system, its mode of operation and analysis protocols, including target identification procedures to deal with non-insect targets (such as precipitation, ‘chaff’, birds and bats) have been described elsewhere (Smith et al. 1993; Smith et al. 2000; Chapman et al. 2002b, 2003; Reynolds et al. 2005).

The radar database was scanned for evidence of layering using a ‘Visual Basic’ module which returned a “Layer Quality” code (a number from 0-7) indicating the layering status of each vertical profile (Reynolds et al. 2005), taking into account the numbers of all resolvable targets and the ‘percentage above threshold values’ (see above). If strong layers occurred in a succession of profiles (at 15-minute intervals) over a period of 2-3 hours during an evening, the relevant profiles were examined in more detail on a case study basis. Insect aerial densities (expressed here as the number of insects per 10^7 m^3) were calculated for targets which were well described by the underlying analysis model (Chapman et al. 2002b), and where estimated masses and other radar-derived variables were expected to be reliable.

Insect data from ground traps

Although the radars can provide information on several variables, which are useful for the identification of radar targets, it is rational to place this information in the context of the relative abundance and temporal occurrence of named species of insect. As we were particularly interested in moths, use of data from the Rothamsted Insect Survey’s (RIS) UK-wide network of light traps (Woiwod and Harrington 1994) complemented the radar-derived variables.

Meteorological data

The UK Met. Office's (UKMO) operational numerical weather prediction model – the Unified Model (UM) – is the source of meteorological data used here. The model assimilates real weather data, along with interpolative tools and physical equations, to provide meteorological output at various locations (grid-boxes) throughout the country. The UM (version 5 onwards) solves non-hydrostatic, deep-atmosphere dynamics using a semi-implicit, semi-Lagrangian numerical scheme (Cullen et al. 1997). The model includes a comprehensive set of parameterizations, including surface (Essery et al. 2001), boundary layer (Lock et al. 2000), mixed phase cloud microphysics (Wilson and Ballard 1999) and convection (Gregory and Rowntree 1990), with additional downdraft and momentum transport parameterizations. The model runs on a rotated latitude/longitude horizontal grid with Arakawa C staggering and a terrain-following hybrid-height vertical coordinate with Charney-Philips staggering. Operationally the UKMO run a 'mesoscale' domain with horizontal resolution of 0.11 degrees (approximately 12.5 km). The model runs with 38 levels spaced non-uniformly in the vertical range. Data are extracted every hour for the present study to provide vertical and temporal profiles of meteorological variables at the grid-box corresponding to the Malvern radar. Several meteorological variables were extracted: most relevant are wind speed, air temperature and relative humidity (RH).

Utilising these data is an improvement over the use of network radiosonde launches alone. Operational radiosonde ascents are available once per night and at locations not particularly close to the radar site (the most relevant upper-air stations currently in use are Camborne, Nottingham, Larkhill and Herstmonceux). Account was nevertheless taken of the radiosonde data in order to check the UM-derived profiles.

In order to estimate the take-off location of radar observed insects, back-trajectories were produced using the NAME trajectory model (<http://www.met->

office.gov.uk/research/nwp/publications/nwp_gazette/dec00/name.html) which uses UM analyses of wind evolution in space and time. This method does not take into account any local turbulence effects on insect flight and does not simulate self-powered insect flight speed. It does, however, assume that the insects maintain a constant height of 300 m above the ground.

All timings referred to in this work are in Coordinated Universal Time (UTC), which is one hour earlier than British Summer Time (BST).

Results

Three case studies of nocturnal layering events were selected from a substantial data-set (2000 onwards). Cases were chosen on the basis of the presence of well-defined and persistent insect layers, apparently consisting of rather similar species, which occurred during stable atmospheric conditions – specifically high atmospheric pressure with largely clear skies (which promote temperature inversions and nocturnal jets). It is worth noting that such meteorological conditions occur typically on 10-15 nights per summer month in southern Britain.

Case study A: 22-23 August 2000

The evolution of the vertical profile of insect aerial density throughout the evening and night of 22–23 August at Malvern (Fig. 1) shows a dusk take-off underway by about 19:30 with aerial-densities up to 100 insects per 10^7 m^3 . At 19:45 the decrease in the insect aerial density with height was semi-logarithmic (c.f. Drake 1984), up to approximately 650 m. A distinct layer in the vertical profile of insects had formed by 21:00 at about 400 m above ground level (agl).

The layer centre decreased in altitude by about 70 m from 23:00 to 01:00, after which the layer dissipated, and aerial densities then remained low for the rest of the night. Maximum densities recorded within the layer were ~ 100 insects per 10^7 m^3 , and layer depth (defined by the 25 insects per 10^7 m^3 contour) varied between 250 and 400 m.

It is highly likely that the insects forming the layer took off at dusk and that no further significant take-off of insects occurred later in the night. Therefore, the insects observed at Malvern at 01:00 had probably been flying for about 5 hours. Back-trajectory analysis revealed that a parcel of air at 300 m present over Malvern at 01:00 would have been located around 150 km to the east at 20:00 (Fig. 2, trajectory A). There is, however, good evidence of approximately downwind orientation by the insects forming the layer (see below), so an additional component of $\sim 4 \text{ m s}^{-1}$, representing the self-powered flight speed of the migrants, can be added to the wind speed. Thus the insects may have travelled a further 70 km between 20:00 and 01:00, giving a total estimated migration distance of around 220 km. This almost corresponds to the distance to the East Anglian coast, and the lack of further sources beyond this point may explain the sudden decline in insect numbers at 01:00. The maintenance of high insect numbers before 01:00 also suggests widespread source locations of the insect species concerned.

The meteorological data generated from the UM at Malvern showed that a surface temperature inversion was present from 21:00 to 04:00 on 22-23 August (Fig. 3). The top of the inversion (i.e. height of maximum temperature, depicted by the magenta line) was located near 300 m agl and showed a slight tendency to rise in altitude until 01:00. The temperature maximum decreased from 16 to 14°C between 21:00 and 01:00. For much of the night the centre of the insect layer was located ~ 50 m above the maximum temperature, implying that most insects were experiencing temperatures of at least 14°C . Furthermore, the upper boundary of the insect layer, located at around 500–550 m, coincided approximately with the 14°C

isotherm throughout (see dashed line denoting the 25 insects per 10^7 m^3 contour in Fig. 3). This may imply that the layer is a ‘ceiling layer’ with a threshold of minimum temperature for flight near 14°C . However, as the threshold for sustained flight in at least some British noctuids is lower than 14°C (e.g. 10.5°C in the mouse moth, *Amphipyra tragopoginis*: Taylor and Carter 1961), the observed upper extent of the layer may be indicative of a preferred temperature rather than an absolute threshold.

Both RH and wind shear featured high gradients near the altitude of 200 m throughout the migration period (not shown). These gradients are associated with the NBL top (e.g. Garratt 1994) which indicates that migration was occurring in the residual layer above, where turbulence is minimal. Further study revealed that the insect layer was located in a layer of less humid air ($< 65 \text{ \%RH}$), with higher values (up to 80 \%RH) above and below the layer.

UM data showed that a nocturnal jet formed from 20:00 to 00:00 (Fig. 4). The wind speed maximum occurred at about 300 m, and was most intense (14 m s^{-1}) between 21:00 and 22:00. Radio-soundings at Herstmonceux, East Sussex (00:00) and Larkhill, Wiltshire (06:00) also provided evidence for a jet, again with the strongest wind speeds (14 m s^{-1}) at about 300 m. The centre of the insect layer observed at Malvern apparently remained no more than 50 m above the jet centre (Fig. 4) and hence typical wind speeds experienced by insects were $12\text{--}13 \text{ m s}^{-1}$.

The displacement speeds (net speed relative to the ground) of insects flying in the layer and recorded by the radar are shown in Fig. 5. The highest displacement speeds ($18\text{--}20 \text{ m s}^{-1}$) occurred between 21:00 and 22:00, i.e. they corresponded approximately to the maximum speeds in the nocturnal wind jet. Scalar subtraction of UM wind speeds from the radar-derived insect displacement speeds gave an estimate of the insect flight speed. The majority of insects had estimated flight speeds of approximately 4 m s^{-1} , and this seems reasonable for the noctuid moths (Lewis and Taylor 1967; Lingren et al. 1995) contributing to the layer (see below).

Radar measurements of body alignment for insects in the layer showed evidence of the common orientation phenomenon (Reynolds and Riley 1986; Riley 1989) (see Fig. 6). There is a 180° ambiguity in the actual insect headings, but consideration of the insects' displacement vectors and flight speed make it clear that the mean orientation was towards the WSW (the mean orientation angle was 244.6° with a circular standard deviation (CSD) of 30.5°, 1062 targets). The mean displacement direction was towards 262.7° (CSD = 25.5°) so the insects were evidently orientating at an angle slightly anti-clockwise from the downwind direction.

The radar-estimated masses of insects in the layer provide an aid to identification (Fig. 7). The insect mass distribution remained almost unchanged throughout the duration of the layer: various sizes of insect were present, but a large peak (comprising 35% of the insects) occurred in the 80–160 mg group. Medium-sized noctuid moths are the most likely component of the nocturnally migrant insect fauna in this size range. Examination of the catch from the nearest RIS light trap to the Malvern radar (Bredon Hill, Worcestershire; 20 km east of the radar) showed that the most common species of noctuid moths caught on this night were: *Xestia c-nigrum* (setaceous hebrew character) *Diarsia rubi* (small square-spot), *Mythimna pallens* (common wainscot), *Ochropleura plecta* (flame shoulder), *Luperina testacea* (flounced rustic) and *Autographa gamma* (silver Y). There is some evidence for windborne migration in *X. c-nigrum* and *M. pallens* (Reynolds et al. 2005). *A. gamma* is a well-known migrant (Taylor et al. 1973) and indeed it has been caught by us in high-altitude aerial samples (see below): its mass (146 mg; n =11) places it within the 80–160 mg peak observed for the insects forming the layer.

Case study B: 14–15 August 2003

The evolution of the vertical profile of insect aerial density throughout the evening and night of 14–15 August 2003 at Malvern is shown as black contour lines on Fig. 8 – which also shows changes in the temperature profile. Dusk take-off lasted until 20:45, and a distinct layer had formed by 21:30 at about 200–450 m agl. The layer rose in altitude until 23:30 when it was located at 300–500 m, and then it decreased with altitude until 01:45, after which time the layer dissipated. Maximum densities of ~ 100 insects per 10^7 m^3 were recorded in the dusk emigration peak, but densities were always lower than 65 insects per 10^7 m^3 within the layer. Density values were thus less than in the layer described in case study A above. Layer depth – defined by the 25 insects per 10^7 m^3 contour – varied from < 70 m up to 200 m. The duration of migration is likely to have been about five hours (20:30 and 01:30). Back-trajectory analysis showed that a parcel of air moving at a height of 300 m and arriving at Malvern at 01:30, would have originated at around 130 km north of the radar site at 20:30 (Fig. 2, trajectory B). The back-trajectory has a curved shape due to a high pressure system centred west of the UK. The insects forming the layer again showed a tendency to orientate downwind – the mean orientation angle was 207.5° (CSD = 29.5° , 642 targets) and the mean displacement direction was towards 205.5° (CSD = 26.9°). Therefore, taking into consideration a component of 4 m s^{-1} for the insect flight speed (see case study A), an extra 70 km of flight distance may have been covered on this night, giving a total estimated migration distance of around 200 km. The fact that the source area of insect take-off is estimated to be near the coast is again a possible reason for the layer dissipation seen at Malvern after 01:45.

The evolution of the temperature profile on the night of 14–15 August is shown in Fig. 8. Before about 23:00, the warmest temperatures occur quite close to the surface (within the first 100 m), and sonde ascents also indicated that the top of the inversion was located at relatively low altitude on this night (e.g. 175 m at Nottingham at midnight). Therefore it seems clear that, in this case, the insects were *not* concentrated at the height of the warmest air. The

temperature profile suggests that up until 23:30 most insects flew in air of 14–16°C, with the upper layer boundary located in the 12–14°C region, suggesting a minimum temperature preference or threshold for flight near 12°C. The majority of insects flew in air of less than 75% RH and there was a general increase of RH with time at all altitudes (not shown). Fig. 9 shows that wind speeds of up to 10 m s⁻¹ were present in a nocturnal jet from 21:00 to 23:00 at 200–300 m, but insects still flying after 01:00 may have experienced winds as low as 5 m s⁻¹. The insect layer appeared to be centred slightly above the wind jet in a region of negative wind shear (i.e. wind speed decreased with altitude). This is similar to the observations in case study A.

The insect mass distribution again showed a peak in the 80–160 mg group (Fig. 7), and medium-sized noctuid moths are the most likely component of the nocturnally migrant insect fauna in this size range. Examination of the RIS catch from Bredon Hill showed that the most common species of noctuid moth caught on this night were: *X. c-nigrum*, *M. pallens*, *O. plecta* and *Thalpophila matura* (straw underwing). The highly migratory species, *A. gamma* (silver Y), was caught in the light-trap in the days before and after this layering event. Further evidence for mass migration of this species was the capture of specimens in a balloon-supported net at 200 m above Cardington airfield, Bedfordshire (52° 06' N, 0° 25' W) on the evenings of 19, 20 and 24 August 2003 (JW Chapman and DR Reynolds, unpublished). Although the noctuids may have constituted a large portion of the radar-detected insects, three migratory micro-moth species were also found in trap catches, namely: *Plutella xylostella* (diamondback moth) (Yponomeutidae), *Nomophila noctuella* (rush veneer) and *Udea ferrugalis* (rusty dot pearl) (both Pyralidae). *Plutella xylostella* is too small (1–4 mg) to have been easily detectable in the layer (Chapman et al. 2002a), but *N. noctuella* (16–25 mg) and *U. ferrugalis* (~10 mg) could have been among the smaller insects detected by the radar in the 10–40 mg size groups (Fig. 7). Altogether, these findings imply that good numbers of moths were

migrating, and indeed migrations in mid to late August are likely to involve southward return movements to over-wintering sites in several species. Certainly, the northerly winds recorded on this night would have aided such a migration strategy.

Case study C: 23–24 August 2003

The evolution of the vertical profile of insect aerial density throughout the evening and night of 23–24 August at Malvern is shown as black contour lines on Fig. 10 – in addition to the evolution of the temperature profile. Dusk take-off was seen to be ending by 20:00 judging by the reduced densities in the lowest range-gate. A distinct layer had formed by 20:30, in the region 200–600 m above ground. The lower edge of the layer was only just visible on the radar, but its centre apparently showed little tendency to change in altitude and its depth was never less than 300 m. The layer dissipated at 00:45, and aerial densities then remained low for the rest of the night. Maximum densities of ~ 90 insects per 10^7 m^3 were recorded near 20:30, and densities above 50 insects per 10^7 m^3 occurred within the layer for much of its duration. Migration duration appeared to be about 5 hours, based on a take-off just before 20:00 and layer dissipation just before 01:00. Back-trajectory analysis (Fig. 2, trajectory C) reveals that a parcel of air at 300 m over Malvern at 01:00 would have originated at around 80 km to the north at 20:00. The mean orientation angle of insects forming the layer was 208.7° (CSD = 39.8° , 1461 targets) and their mean displacement direction was towards 210.3° (CSD = 28.7°). Thus an extra 70 km can be added to take account of the insect flight speed (see case study A), and this suggests an estimated migration distance of around 150 km. Hence, the proposition that insects observed near the end of the layering event (01:00) took off near the coast cannot be discounted, particularly given the curvature of the horizontal wind field shown on back-trajectory C.

The evolution of the temperature profile (Fig. 10) shows that temperatures were distinctly warmer than in case studies A and B. Most insects flew above the height of the temperature maximum, in air of 18-21°C throughout the night. The upper layer boundary was located in the 17–18°C region; this is likely to be higher than the minimum temperature threshold for flight, given the likely species constituting the layer. This implies that temperature was not limiting insect flight altitude, as was suggested in cases A and B. It is possible that in this case maximum wind speed was the significant variable determining layer altitude. Fig. 11 shows that a nocturnal jet was present, with maximum speeds of 8 m s⁻¹ between 21:00 and 00:00 at height 200–300 m. The jet centre was located at the same altitude as the centre of the insect layer and most insects would have experienced wind speeds above 6 m s⁻¹.

The insect mass distribution again showed a peak in the 80–160 mg group (Fig. 7). The nearest working RIS light-trap was at Hereford (30 km west of the radar), and noctuids caught included *X. c-nigrum*, *M. pallens* and the rare UK migrant *Spodoptera exigua* (small mottled willow or beet armyworm). The northerly wind experienced in this migration event is consistent with southward return migrations to over-wintering sites. Catches of *S. exigua* in the RIS light-traps first appeared in June, probably indicating an early northward invasion of the species and subsequent return southwards (c.f. Johnson 1969, p. 516).

Discussion

Large insects flying in the stable atmospheric boundary layer at night would be expected to have more control over their altitude of migration than, say, small insects flying under convective conditions during the day (Gatehouse 1997). We envisage the migrant moths (which form the subject of the present study) climbing steeply after take-off in order to rise

above their FBL (Johnson 1969 p. 81; Lingren et al. 1995), and then ascending more gradually (at $\sim 0.5 \text{ m s}^{-1}$; Riley et al. 1983) until they reach altitudes of several hundred metres where conditions seem optimal for migratory flight. The migrants will then tend to accumulate at these altitudes, and if the resulting concentrations are relatively restricted in depth, they will be perceived on the radar as layers. Apart from the effects of atmospheric conditions on flight altitude (see below), there are presumably other limits on the vertical distance a large insect will climb before it levels out – these may be controlled by internal physiological restraints such as energy expended in climbing flight, or conceivably by optomotor reactions to ground patterns (of which little is known for high-flying insects: Riley 1989). After reaching their ‘cruising’ altitude, nocturnal migrants will maintain steady and continuous flight, often for a period of several hours, during which time they will be displaced considerable horizontal distances in an approximately downwind direction. In southern Britain, migrations are usually over by about midnight or 01:00 (present study; Reynolds et al. 2005). Flights of moths continuing through the whole night until dawn or beyond (which have been observed in other regions of the world; Drake et al. 1981; Drake 1985; Wolf et al. 1990; Beerwinkle et al. 1994; Feng et al. 2004) are apparently uncommon in the UK (Reynolds et al. unpublished data). Since Britain is an island, some of the more abrupt flight terminations may be due to a lack of source areas beyond the coasts – indeed this may have occurred in the current study – rather than because air temperatures have dropped below thresholds for sustained flight or because flight fuel reserves have been exhausted.

A key question is thus: which environmental factors present in, for example, the first kilometre of the nocturnal atmosphere will have most influence on the migration altitude of large insects? Temperature would be expected to be a primary influence, as this variable affects many other aspects of insect physiology and behaviour, and there are plenty of studies to support this view (Drake and Farrow 1988; Gatehouse 1997). The simplest case is where the

insects have selected the altitude of the warmest air, often at the top of a surface temperature inversion (Schaefer 1976; Drake 1984; Drake and Farrow 1988; Feng et al. 2003; Reynolds et al. 2005) or occasionally a higher-altitude temperature maximum, such as that due to a subsidence inversion (Reynolds et al. 2005). Selection of the warmest air by migrants appears to be most likely to occur in relatively cool conditions, and in taxa that have high optimum temperatures for migratory flight. For example, migratory acridoid insects (grasshoppers and locusts) have optimum temperature values for sustained flight of above 20°C (Clark 1969; Riley and Reynolds 1979), which are much higher than for instance the noctuid moths studied here (see also Taylor and Carter, 1961). On the other hand, there are many references in the literature in which insects, particularly moths, have ascended above the altitude of the temperature maximum. On some of these occasions insects may be forming ‘ceiling layers’, i.e., ascent has continued until insects reach an altitude corresponding to the lowest temperature at which they can sustain flight. A good example is the sharp upper boundary of layers of the brown planthopper, *Nilaparvata lugens* in China (Riley et al. 1991): these layers were well above the altitude of the temperature maximum, but the layer tops corresponded to known temperature thresholds (*ca.* 16 °C) for sustained flight in the planthoppers. Ceiling layers may also be implicated in cases of high-altitude layering where there is no obvious corresponding feature in the vertical profile of meteorological variables (Drake and Farrow (1985) observed one as high as 1900 m agl in eastern Australia), and in cases where maximum flight altitudes of certain taxa (grasshoppers, say) show a general decrease in line with seasonal air temperatures (Schaefer 1976; Reynolds and Riley 1997).

The present observations in southern UK were made in a cooler climate than most previous radar entomology studies, and it was to be expected that even noctuid moth migration would be strongly restrained by temperatures on many occasions. A good indication that temperatures were sub-optimal on many nights was the observation that when migratory

activity occurred at dusk, it frequently did not persist for long after dark (Wood, Reynolds et al. unpublished data). When night-time layering *did* develop, moths have been observed to fly at the altitude of the warmest air (Reynolds et al. 2005), but sometimes it may be difficult to distinguish (as in our case study A above) between this effect of temperature and the formation of a ‘ceiling’ layer. However, in our case study B, the observed insect layer was well above the height of the temperature maximum, and was most easily explained by a restriction on migratory flight due to the cooler air at higher altitudes.

Notwithstanding the above findings, there are many reports in the literature where the insect layers are closely associated with wind-related variables (i.e. wind velocity, shear zones, turbulence) and conspicuously *unrelated* to air temperature profiles (Wolf et al. 1986; Hobbs and Wolf 1989; Beerwinkle et al. 1994; Feng et al. 2004). A necessary condition in these cases is presumably that night-time air temperatures are significantly above flight thresholds for the taxa concerned, freeing the insects of the need to migrate at the warmest altitudes. Examples where moths contributed to wind-related layers include *Helicoverpa zea*, *Heliothis virescens*, *Peridroma saucia* and other species in the southern USA (Wolf et al. 1986; Beerwinkle et al. 1994), and *Loxostege sticticalis* and *Helicoverpa armigera* in north-eastern China (Feng et al. 2004). It seems clear that large insects – such as migratory noctuid moths – are able to detect zones of wind speed maxima, and to fly preferentially within them (Wolf et al. 1986), and this would appear to be an adaptive strategy for maximizing their displacement. Moreover, the migrants are often able either to align themselves in a downwind direction (as in the present study), or to orient at an angle to the wind (but generally one which avoids gross backwards (tail-first) displacement: Riley and Reynolds 1986). The mechanism(s) and adaptive significance of this orientation behaviour are still unclear. In some cases, orientation occurs under severely reduced illumination, which may suggest that insects are able to use non-visual

cues to detect wind speed and direction, such as anisotropies in turbulence due to Kelvin-Helmholtz waves (Riley 1989).

Because the boundary-layer wind speed maximum often occurs close to the top of the surface inversions, it can be difficult to distinguish the effects of wind speed from those of temperature. In our case study C, however, an association with the nocturnal jet seemed likely, as the migrants were evidently flying above the level of the warmest temperatures, and ‘ceiling layer’ effects seemed unlikely because layers of similar species have been observed to migrate at much lower temperatures on previous nights (c.f. case study B).

In summary, the results of the current study indicate that the altitude of layers of migrating moths in the UK may be constrained either by: the altitude of the warmest air (case study A); the altitudes with temperatures which may represent flight thresholds and/or preferenda (case studies A and B); or the altitude of regions of high wind speed when air temperatures are relatively high (case study C).

A case study approach is clearly useful for investigation of migration events involving a preponderance of particular species, as demonstrated by the recent studies of noctuid moths in the UK (the present paper; Reynolds et al. 2005) and elsewhere (Feng et al. 2003; 2004). However, the existence of continuous, high-resolution, co-located meteorological and entomological data-sets available, respectively, from the Unified Model and the vertical-looking radar, seem highly suitable for a systematic investigation of the meteorological mechanisms controlling insect layering: we are currently embarking on such a study (CR Wood, in preparation).

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Legends

Fig. 1 Evolution of insect aerial density (numbers per 10^7 m^3) with altitude and time for the night of 22-23 August 2000 at Malvern. The insect layer formed at 20:00UTC at 250 m and (rising in altitude) persisted until 01:00UTC. Only insects with mass 10 mg or greater have been included. Sunset was at 19:18, and the end of dusk is represented by a marker on the time axis at 19:48.

Fig. 2 Back trajectories calculated for an air parcel located near the insect layer height (300 m agl) and crossing the Malvern site at the time of layer termination. The trajectories show the presumed origin of the air parcel at dusk (5 hours previously). (A) 22-23 August 2000, (B) 14-15 August 2003, (C) 23-24 August 2003.

Fig. 3 Evolution of temperature ($^{\circ}\text{C}$) with altitude and time for the night of 22-23 August 2000 at Malvern. The magenta coloured line marks the top of the temperature inversion (altitude of warmest air). Superimposed are the insect aerial density contours from Fig. 1 for 25 (---) and 50 (—) insects per 10^7 m^3 , which indicate the development of the insect layer.

Fig. 4 Evolution of wind speed (m s^{-1}) with altitude and time for the night of 22-23 August 2000 at Malvern. Superimposed are the insect aerial density contours from Fig.1 for 25 (---) and 50 (—) insects per 10^7 m^3 , which indicate the development of the insect layer.

Fig. 5 Evolution of net insect displacement speed (ground speed) (m s^{-1}) with altitude and time, on the night of 22-23 August 2000 at Malvern. Sunset was at 19:36, and the end of dusk is represented by a marker on the time axis at 20:06.

Fig. 6 Equi-areal plot showing distribution of body alignments for insects flying in the layer in range-gates 3-5 at 290-480 m above ground from 21:00-01:00 UTC on 22-23 August 2000. 14.9% of the targets were aligned along $60 - 240^\circ$ axis.

Fig. 7 Mass distributions of insects forming nocturnal layers at Malvern. (A) 20:00-01:00 UTC on 22-23 August 2000 (B) 21:00-01:00 on 14-15 August 2003 and (C) 21:00-01:00 on 23-24 August 2003. Only insects of mass 10 mg or greater have been included.

Fig. 8 Evolution of temperature ($^\circ\text{C}$) with altitude and time, on the night of 14-15 August 2003 at Malvern. Superimposed are the insect aerial density contours of 25 (---) and 50 (—) insects per 10^7 m^3 as detected by the Malvern radar, which indicate the development of the insect layer. Sunset was at 19:36, and the end of dusk is represented by a marker on the time axis at 20:06.

Fig. 9 Evolution of wind speed (m s^{-1}) with altitude and time, on the night of 14-15 August 2003 at Malvern. Superimposed are the insect aerial density contours of 25 (---) and 50 (—) insects per 10^7 m^3 , which indicate the development of the insect layer.

Fig. 10 Evolution of temperature ($^\circ\text{C}$) with altitude and time, on the night of 23-24 August 2003 at Malvern. Superimposed are the insect aerial density contours of 25 (---) and 50 (—)

insects per 10^7 m^3 , which indicate the development of the insect layer. Sunset was at 19:18, and the end of dusk is represented by a marker on the time axis at 19:48.

Fig. 11 Evolution of wind speed (m s^{-1}) with altitude and time, on the night of 23-24 August 2003 at Malvern. Superimposed are the insect aerial density contours of 25 (---) and 50 (—) insects per 10^7 m^3 , which indicate the development of the insect layer.











