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The Hiccup: a dynamical coupling process during the autumn transition in the Northern Hemisphere – similarities and differences to sudden stratospheric warmings

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Abstract. Sudden stratospheric warmings (SSWs) are the most prominent vertical coupling process in the middle atmosphere, which occur during winter and are caused by the interaction of planetary waves (PWs) with the zonal mean flow. Vertical coupling has also been identified during the equinox transitions, and is similarly associated with PWs. We argue that there is a characteristic aspect of the autumn transition in northern high latitudes, which we call the “hiccup”, and which acts like a “mini SSW”, i.e. like a small minor warming. We study the average characteristics of the hiccup based on a superimposed epoch analysis using a nudged version of the Canadian Middle Atmosphere Model, representing 30 years of historical data. Hiccups can be identified in about half the years studied. The mesospheric zonal wind results are compared to radar observations over Andenes (69° N, 16° E) for the years 2000–2013. A comparison of the average characteristics of hiccups and SSWs shows both similarities and differences between the two vertical coupling processes.

Keywords. Meteorology and atmospheric dynamics (general circulation; middle atmosphere dynamics; waves and tides)

1 Introduction

Dynamical stratosphere–mesosphere coupling can be caused by gravity waves (GWs), tides and planetary waves (PWs), and strongly influences the dynamics of the middle atmosphere (e.g. Miyahara et al., 1993; Holton and Alexander,

2000). The most impressive and prominent vertical coupling process in the middle atmosphere is the sudden stratospheric warming (SSW), which occurs primarily in the Northern Hemisphere winter and is caused by the interaction of PWs with the mean flow (e.g. Matsuno, 1971; Andrews et al., 1987). SSWs are characterised by a rapid increase of the stratospheric polar temperature by up to 80 K with a simultaneous decrease of the mesospheric polar temperature by up to 30 K, along with a wind reversal/weakening of the zonal mean zonal wind in the middle atmosphere (e.g. Charlton and Polvani, 2007; Matthias et al., 2012). The mesospheric cooling is understood to be caused by anomalous upwelling driven by anomalous filtering of GW fluxes (e.g. Holton, 1983; Ren et al., 2008), although anomalous upwelling from the PW drag also contributes in the lower mesosphere (e.g. Ren et al., 2008).

Strong vertical coupling has also been identified during the equinox transitions between the easterly zonal flow, which characterises the summer half of the year, and the westerly flow, which characterises the winter half. Distinctive “sawtooth” mesospheric temperature and airglow enhancements, which pointed to circulation anomalies, were identified during the Northern Hemisphere spring transition by Stegman et al. (1992) and G. G. Shepherd et al. (1999). Their association with the final stratospheric warming was noted by M. G. Shepherd et al. (2002), pointing to the importance of vertical coupling. A similar (but opposite signed) “sawtooth” (or “V-shaped”) variability was identified in the autumn transition by Taylor et al. (2001) using local lidar measurements

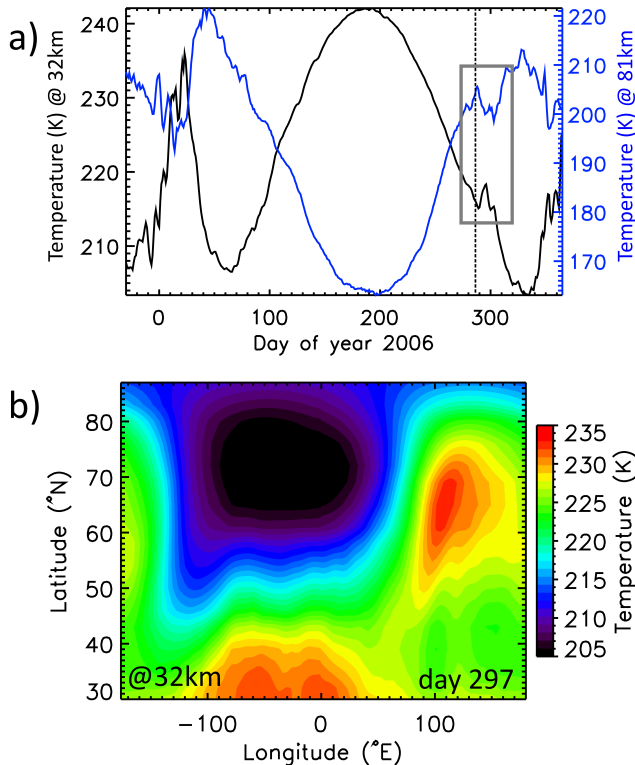


Figure 1. (a) Seasonal variation of the zonal mean temperature at 32 km (black) and 81 km (blue) averaged between 65 and 75° N over the year 2006. The vertical dashed line marks the hiccup onset day. (b) Temperature pattern as a function of longitude and latitude on day 297 of 2006, the local peak of the temperature at 32 km in (a). The data are derived from the CMAM30 data set (see text).

at mid latitudes, and by M. G. Shepherd et al. (2004) using global satellite observations from WINDII. G. G. Shepherd et al. (2004) suggested that the two equinox transitions should be treated as related phenomena.

Based on TIME-GCM simulations, Liu et al. (2001) argued that the variability in the autumn transition resulted from enhanced vertical propagation of stationary PWs (SPWs) into the mesosphere, expected during the weak (but rapidly changing) zonal winds characteristic of that time. The same mechanism would obviously apply to the spring transition. However, enhanced SPW amplitudes on their own would induce temperature perturbations of opposite sign at different longitudes, and M. G. Shepherd et al. (2004) showed that the zonal mean temperature also exhibited a “V-shaped” cooling during the autumn transition at 87 km. Since enhanced SPW propagation into the mesosphere would be expected to lead to enhanced polar descent and hence to warming, it seems likely that more than PW propagation is involved.

During the autumn transition, the zonal wind reversal from easterlies to westerlies allows the onset of SPW propagation into the middle atmosphere (Charney and Drazin, 1961). If

some of this SPW activity dissipates in the stratosphere, it would result in stratospheric polar downwelling and warming, which would weaken the zonal wind, somewhat like a mini SSW, i.e. like a small minor warming regarding the strength of the wind weakening. Thus, one might generally expect a sudden weakening of the zonal wind (and stratospheric polar warming) as SPW propagation switches on, before the wind resumes its strengthening as part of the overall autumn transition. We call such behaviour a “hiccup” in the autumn transition. Just as with SSWs, we may expect a concomitant, and largely synchronous, mesospheric zonal-mean “hiccup” of opposite sign in temperature, as a result of PW and GW induced polar upwelling (e.g. Siskind et al., 2005; Ren et al., 2008). The phenomenon is evident in the grey rectangle in Fig. 1a, which shows the seasonal variation of the zonal mean temperature around 70° N at 32 km (black) and 81 km (blue) for 2006. Figure 1b confirms that there is a prominent wave-1 feature at 32 km on day 297, when the peak 32-km temperature during the hiccup in Fig. 1a is observed.

In principle, the same process might be expected to be operative in the spring transition, where the sudden cessation of SPW propagation as the zonal winds weaken would turn off the stratospheric polar downwelling, allowing the zonal winds to suddenly strengthen before continuing their transition towards summertime easterlies. The stratospheric and mesospheric temperature anomalies in this case would be opposite in sign to those in the autumn “hiccup”, consistent with the mesospheric observations of M. G. Shepherd et al. (2002). However, the spring transition is much more variable than the autumn transition, because it occurs within an already disturbed environment, due to the high level of wave activity in the winter and spring (e.g. Wunch et al., 2005; Manson et al., 2002). These disturbances of the background wind make it hard to set the hiccup onset day. So the effect should be more easily detectable in the autumn transition where it is proceeding from the quiescent summertime state.

The goal of this study is to examine the morphology of the stratosphere–mesosphere coupling during the “hiccup” in the autumn transition using a new model data set, which provides an estimate of the stratospheric and mesospheric daily state over the last 30 years. We focus on the average characteristics of the hiccup over the polar cap (60–90° N), which is the latitude range where the hiccup appears to be strongest (cf. Fig. 1b). Although the hiccup is manifest in zonal-mean fields, we also want to see whether it might explain at least some of the behaviour seen at particular longitudes (e.g. Taylor et al., 2001). We use 30 years of historical data (1979–2010) of the nudged Canadian Middle Atmosphere Model CMAM30 and compare the wind results in the mesosphere with 14 years of local radar measurements (2000–2013) in Andenes (69° N, 16° E).

This paper is structured as follows. The next section provides an overview of the CMAM30 model simulation and a short specification of the radar measurements used in this

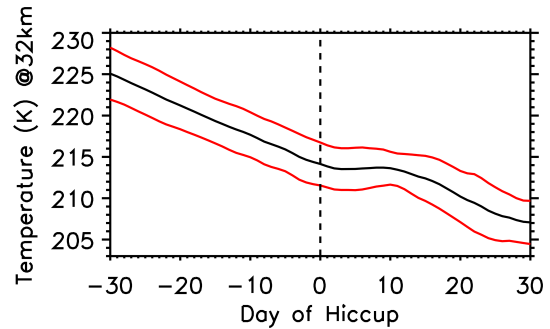
Table 1. Hiccup onset days of the hiccup-years, from 1979 to 2013.

Year	Hiccup onset day
1979	295
1980	284
1984	284
1985	294
1988	287
1990	288
1991	301
1992	305
1994	305
1996	296
1998	294
2003	275
2004	284
2005	285
2006	285
2008	299
2009	278
2010	297
2013	283

study, as well as a description of the analysis method. The results of the superimposed epoch analysis are shown in Sect. 3 and discussed in Sect. 4. Section 5 provides a summary of the results.

2 Data and analysis

To study the average characteristics of wind, temperature and PW 1 during the hiccup, the nudged version of the Canadian Middle Atmosphere Model (CMAM) is used (de Grandpré et al., 2000; Scinocca et al., 2008; McLandress et al., 2013). The model domain extends from 1000 to 0.008 hPa (0–100 km) with a vertical resolution of better than 1 km in the troposphere and 2.5 km in the mesosphere, and a horizontal resolution of 3.75° in latitude and longitude. However, the upper levels of the model will be contaminated by the sponge layer and the data above about 85 km are probably not reliable. The data analysed here have a temporal resolution of 6 h, and daily mean values are used for this study to avoid tidal impact. In this nudged configuration, the horizontal winds and temperatures are nudged to ERA-Interim below 1 hPa, and it is a free running model above. The resulting data set, which covers the 30-year period 1979–2010, is known as the CMAM30 data set. Constraining the dynamical state of the troposphere and stratosphere allows information to propagate up into the mesosphere through both PWs and parameterised GW fluxes, which strongly constrains the large-scale state of the mesosphere (Ren et al., 2008). McLandress et al. (2013) show that the CMAM30 data are in good agreement with MLS satellite observations in the mesosphere during the several extended winters studied. Thus, CMAM30 data pro-

**Figure 2.** Zonal mean temperature variation (black) and its standard deviation (red) at 32 km averaged between 65° and 75° N, averaged over all hiccups. The data are derived from CMAM30 for events occurring from 1979 to 2010, and from MERRA for events occurring from 2011 to 2013.

vide a good basis to study vertical coupling processes from the troposphere up to the mesosphere, and compare directly with observations over a long time period. The data are available from: <http://www.cccma.ec.gc.ca/data/cmam/cmam30/index.shtml>.

To corroborate the model results in the mesosphere, wind measurements of a mid-frequency (MF) radar in Andenes (69° N, 16° E) are used. The MF radar operates at 1.98 MHz with a peak power of 40 kW applying the spaced antenna technique (Singer et al., 1997). Horizontal winds have been continuously derived between 80 and 98 km since the year 2000 using the full correlation analysis method. Prevailing winds are obtained by subtracting terdiurnal, semidiurnal and diurnal tides using a least squares fit of hourly winds for a 4-day interval shifted by 1 day.

In order to apply the superimposed epoch analysis, we have to define what is meant by a hiccup, and an onset day. Motivated by Fig. 1a, the hiccup onset day is defined as the first day in each autumn where at 10 hPa and at 70° N:

1. the zonal mean zonal wind is larger than 15 m s^{-1} ,
2. $du/dt < 0$ for at least 5 days, and
3. the mean zonal wind subsequently increases above the value that was maximally reached before the hiccup.

The first condition is chosen because SPWs can only propagate into the middle atmosphere if the zonal wind is larger than about 15 m s^{-1} (Nash et al., 1996). The second condition ensures a sufficiently discernible hiccup. The third condition limits the events to the autumn transition. Based on these criteria, a hiccup was found to occur in 19 out of the last 35 years (see Table 1), i.e. it occurs on average every second year. Thus, the composites of the CMAM30 data include 18 events, while those of the radar data includes eight events. Note that from 2011 to 2013 MERRA data (e.g. Rienecker et al., 2011) were used to estimate the hiccup onset day, since the CMAM30 data only cover up to 2010. Figure 2 shows the

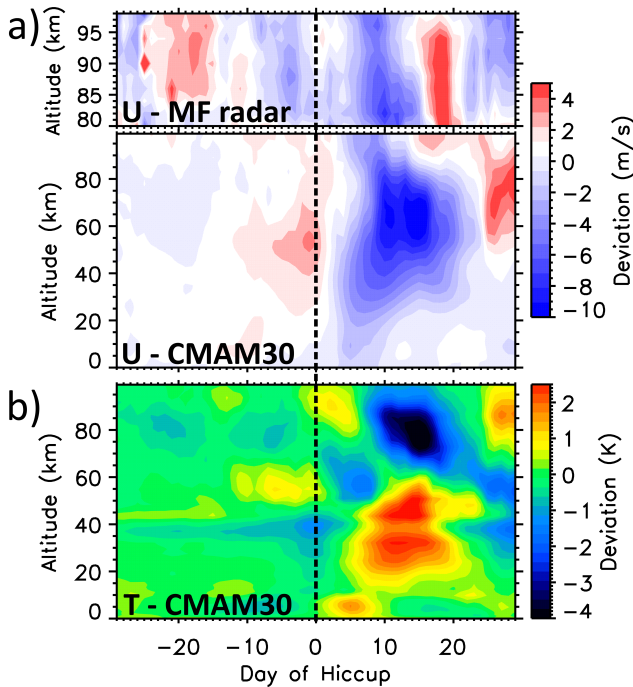


Figure 3. (a) Hiccup composite of the zonal wind anomaly at Andenes (69° N, 16° E) derived from MF radar data (top panel), and of the zonal mean zonal wind anomaly averaged between 60° and 90° N derived from CMAM30 data (bottom panel). (b) Hiccup composite of zonal mean temperature anomaly averaged between 60° and 90° N derived from CMAM30 data.

zonal mean temperature variation at 32 km centred around the hiccup onset, averaged over all 19 hiccups. This shows a similar behaviour to that seen in Fig. 1a for 2006, though of course smoothed out, with the temperature increasing after the hiccup onset day following a longer period of smoothly decreasing temperature in the stratosphere, before resuming its seasonal decrease after about 10 days.

Due to the small effects of a hiccup compared to those of a SSW, we cannot readily define anomalies from multi-year daily means since the latter are too noisy. Thus, we prepared the data as follows, before applying the superimposed epoch analysis:

1. a linear fit of atmospheric parameters over the autumn transition was prepared at each location;
2. the average linear fit was removed from each year;
3. the residual fields were averaged over non-hiccup years;
4. this average was then removed from each hiccup year to define the hiccup anomaly.

Non-hiccup years are those years between 1979 and 2013 that are not listed in Table 1. Please note that the resulting hiccup composites are statistically significant for values greater/lower than $\pm 1 \text{ m s}^{-1}$ or K.

3 Results

The hiccup composite of the zonal mean zonal wind deviation from the surface to 100 km and averaged from 60° to 90° N is shown in Fig. 3a, bottom panel. Although our selection criteria applies only at 32 km, the hiccup, i.e. the decrease of the zonal mean zonal wind, occurs from the troposphere right up to the mesosphere. It appears that the decrease of the zonal wind begins in the troposphere and propagates upwards into the mesosphere within a few days, which would be consistent with the role of upward propagating SPWs. Before the onset of the hiccup, there is an increase of the zonal wind in the stratosphere. It is presumably this feature which allows the enhanced SPW propagation that causes the hiccup. Comparison of these results to the average behaviour of major SSWs show that both coupling processes are characterised by a wind weakening/reversal in the middle atmosphere (e.g. Limpasuvan et al., 2004; Matthias et al., 2013). However, while the hiccup seems to propagate upward, studies of Matthias et al. (2012) show an ~ 4 days earlier onset of the wind reversal in the mesosphere than in the stratosphere in composite analysis of major SSWs implying a downward propagation.

Figure 3a top shows the composite of the zonal wind at Andenes (69° N, 16° E) in the mesosphere derived from local radar measurements. The decrease of the zonal wind is in good agreement with the mesospheric model results and thus corroborate them. However, there are differences between the observations and the model results. While the observations show a wave-like structure in the mesosphere especially before the onset of the hiccup, this structure does not occur in the CMAM30 data in the same way. These differences might occur because the model is free running above 1 hPa, because the composites include different years, or because the GW parametrisation scheme is not perfect. Note that we repeated the CMAM30 zonal mean wind analysis for an area around Andenes and got a similar behaviour of the hiccup as in the zonal mean (not shown). Thus, we omit the local version of Fig. 3a, bottom panel, and compare qualitatively the zonal mean model fields with the local observations.

The hiccup composite of the zonal mean temperature deviation from the surface to 100 km and averaged from 60° to 90° N is shown in Fig. 3b. The hiccup is characterised as a warming in the troposphere and stratosphere, and a cooling in the mesosphere. The temperature changes seem to lag the wind changes by a few days in the middle atmosphere. (They need not be synchronous, despite thermal-wind balance, because they are at the same latitude.) As was observed for the zonal wind, the temperature changes seem to propagate upward from the troposphere into the mesosphere within a few days, and there is a localised cooling in the stratosphere immediately before the hiccup (consistent with the strengthened zonal wind just before the onset day). Both the zonal wind and temperature show a downward-moving structure in the mesosphere from about day 10 onwards. Com-

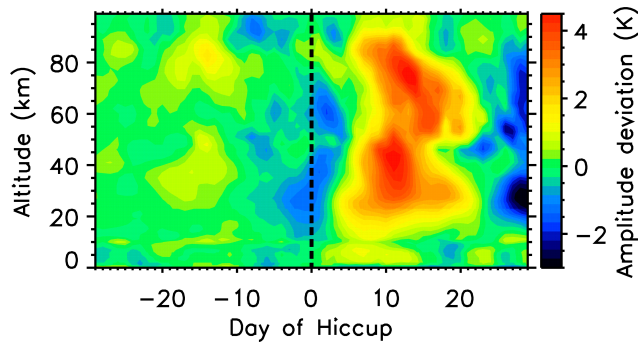


Figure 4. Hiccup composite of the PW 1 anomaly averaged from 60 to 90° N centred around the hiccup onset day and derived from CMAM30 data.

parison of these results to the average behaviour of SSWs shows that the temperature variations in the middle atmosphere are very similar in hiccups and in SSWs. Thus hiccups can be considered as “mini SSWs”. However, as was the case with the zonal wind, it seems that there is an upward propagation of the temperature changes during the hiccup while Limpasuvan et al. (2004) observed a downward propagation of the temperature anomalies during SSWs. These differences might arise because while SPW propagation is not suppressed during hiccups, it is suppressed during SSWs (Hitchcock et al., 2013), allowing downward-propagating wave-mean flow interaction to occur (Hitchcock and Shepherd, 2013).

Figure 4 shows the composite of PW 1 amplitude deviation centred around the hiccup onset day from the surface to 100 km and averaged from 60 to 90° N. During the hiccup the composite shows a strong increase in the PW 1 amplitude throughout the lower and middle atmosphere. The decrease in PW 1 amplitude in the stratosphere just before the hiccup onset day is presumably the precursor anomaly that leads to the positive zonal wind anomaly seen in Fig. 3a, bottom panel. Comparing these results to the average behaviour of PW 1 around SSWs (e.g. Pancheva et al., 2009; Matthias et al., 2012) we determine that during both events the amplitude of PW 1 is increased throughout the middle atmosphere. However, while the amplitude of PW 1 is high before a SSW (e.g. Matthias et al., 2012), it is very low before the hiccup. Part of this difference is caused by the difference in definition of the onset day – SSWs are defined after a period of zonal-wind decline, while we have defined the hiccup onset day to be before the period of zonal-wind decline – but part of the difference suggests that to initiate a hiccup, a precursor anomaly is required.

4 Discussion

Our study was inspired by the first report of a perturbation of the autumn transition by Taylor et al. (2001). They observed

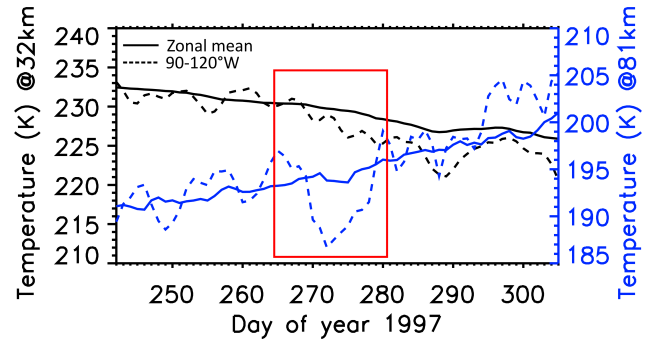


Figure 5. Temporal development of the temperature at 32 km (black) and 81 km (blue) averaged between 35 and 45° N of the autumn transition in 1997. The solid line represents the zonal mean and the dashed line the area around Ft. Collins (90–120° W). The data are derived from CMAM30.

a decrease of the mesospheric temperature of up to 23 K in Ft. Collins (41° N, 105° W) in the autumn transition of 1997 from day 260 to 280. However we did not identify a hiccup in 1997 (cf. Table 1). To examine whether this mid-latitude perturbation nevertheless acts like a hiccup, Fig. 5 shows the temporal development of the CMAM30 temperature averaged from 35 to 45° N at 32 km (black) and 81 km (blue), similarly as in Fig. 1a. The solid lines represent zonal mean values, and the dashed lines averages around Ft. Collins (90–120° W). The CMAM30 temperature shows a strong cooling in the mesosphere around Ft. Collins between day 260 and 280 (red rectangle), which is in good agreement with Taylor et al. (2001). There is a much smaller cooling in the zonal mean, suggesting that in the model the cooling over Ft. Collins is associated with a strong PW 1 magnitude but not with a zonal-mean circulation anomaly, as we have argued must be the case for a hiccup. The time period of occurrence of the small zonal-mean cooling is in agreement with the climatological zonal mean studies of WINDII temperatures between 35 and 45° N at 87 km by M. G. Shepherd et al. (2004), but we cannot use this version of the CMAM to examine behaviour at such high altitudes. In contrast to the high latitude hiccup seen in Fig. 1a, the stratospheric temperature of the mid latitude perturbation shows no warming during the time period of the mesospheric cooling, neither in the zonal mean nor in the regional temperature development shown in Fig. 5. Thus, the 1997 mid latitude perturbation seems to occur only at mesospheric heights, while the hiccup is a stratosphere–mesosphere coupling process.

Another indication that there are two distinct types of perturbations to the middle atmosphere during the autumn transition, differing in their latitudinal structure, comes from Manson et al. (2002) who studied the average seasonal variation of the zonal wind over Saskatoon (52° N, 107° W) from year 1991 to 1998 in the MLT. Manson et al. (2002) show a wind reversal to westward winds from day 260 to 290, which agrees with the observations of Taylor et al. (2001) and

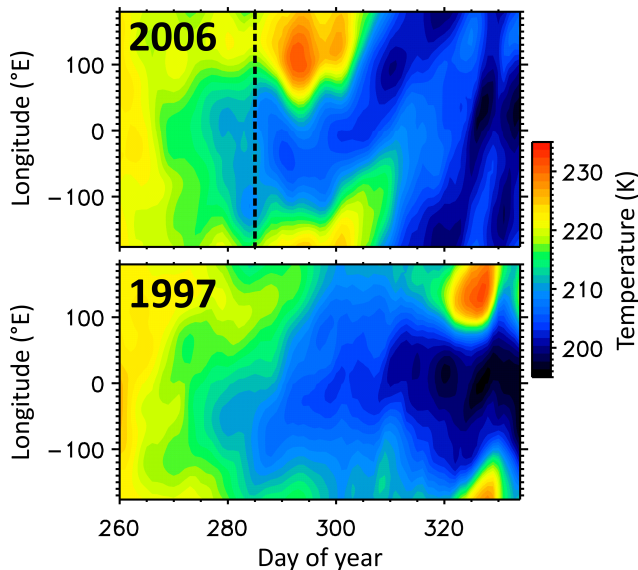


Figure 6. Temporal development of the longitudinal temperature pattern at 32 km in 2006 (top) and 1997 (bottom) averaged between 60 and 90° N. The vertical dashed line in 2006 marks the hiccup onset day. The data are derived from CMAM30.

M. G. Shepherd et al. (2004), and a second reversal starting at day 300, which is in agreement with our hiccup definition in this time period. The Saskatoon observations are located far enough north to see hiccups (cf. Fig. 1b), but not so far north to miss the mesospheric perturbations at mid latitudes. High latitude observations of Espy and Stegman (2002) at Stockholm (59° N, 18° E) from year 1991 to 1998 on average show a decrease of the mesospheric temperature from day 290 to 320, which is also in agreement with our observations of hiccups in this time period.

A possible cause for the mesospheric temperature perturbation of the autumn transition was first advanced by Liu et al. (2001), and more recently by Stray et al. (2014). Liu et al. (2001) argue that PW 1 transience, in both amplitude and phase, associated with the weak (but rapidly changing) zonal winds during the autumn transition can lead to large variability in local temperature measurements at mid and high latitudes. Liu et al. (2001) support their hypothesis with nudged TIME-GCM simulations of the 1997 autumn transition, which qualitatively (though not quantitatively) account for the mesospheric temperature perturbation observed by Taylor et al. (2001). In agreement with Liu et al. (2001), Stray et al. (2014) found a clear enhancement of the PW amplitude in the MLT in the composite (2000–2008) during autumn equinox using SuperDARN radars. As noted above, however, the 1997 event is not a hiccup according to our definition, and enhanced PW amplitudes at mesospheric heights appear to be a distinct phenomenon from the hiccup in the stratospheric zonal wind transition and its mesospheric echo.

Another attempt to explain perturbations to the autumn transition was presented by Espy and Stegman (2002). They assumed that the perturbation period coincides with the period when the atmosphere is virtually transparent to upward propagating GWs, and corroborated this hypothesis with a simple model of GW propagation with a constant flux of GWs throughout the year. The model results show that during the autumn transition, 75 to 90 % of the GWs reach the MLT. Thus they argue that the mesospheric temperature perturbations around the autumn equinox result from combined effects of upward propagating PWs and GWs. We would not disagree with that conclusion but emphasise the need for a hiccup in the stratospheric zonal wind transition, associated with the onset of SPW propagation into the middle atmosphere, in order to initiate the distinctive stratosphere–mesosphere coupling seen in Fig. 1a.

Therefore we regard the hiccup as a distinct phenomenon from that discussed by Liu et al. (2001). Nevertheless, enhanced PW 1 amplitudes are an essential part of hiccups (needed to initiate the weakening of the stratospheric high-latitude zonal wind). As a further means of distinguishing the two phenomena, we may consider the PW 1 phase. Liu et al. (2001) state that changes of the PW 1 phase also cause the variability of the mesospheric temperature, but clearly, changes in phase can have no effect on the zonal mean. Therefore we considered the temporal development of the longitudinal temperature pattern at 32 km for all 35 years. Figure 6 illustrates this pattern for the non-hiccup autumn transition in 1997 and for the hiccup transition in 2006. We find that the phase of the temperature wave 1 pattern is stable for about 10–15 days during most of the hiccup events, while it can change rapidly during autumn transitions without a hiccup. This makes physical sense since hiccups are essentially a delay, or hiatus, in the autumn transition, during which the zonal wind is comparatively stable. Thus, while the rapidly changing PW 1 phase can certainly induce dramatic mesospheric temperature perturbations over particular longitudes, as argued by Liu et al. (2001), these are to be distinguished from high-latitude hiccup events.

5 Conclusions

We have identified a novel high-latitude stratosphere–mesosphere coupling process, called a hiccup, which is responsible for some of the distinctive mesospheric temperature perturbations seen in observations during the autumn transition. The hiccup is fundamentally defined in the stratospheric high-latitude zonal wind, where it originates from the onset of SPW propagation into the middle atmosphere as the zonal winds become westerly during the autumn transition, leading to a sudden dynamical weakening of the stratospheric winds before they eventually resume their overall strengthening. Thus it is like a “mini SSW”, albeit with a distinctively smaller magnitude. The availability of the CMAM30 data

set, covering a 30-year period, provides an opportunity to characterise the average properties of hiccups throughout the middle atmosphere, which according to our definition occur on average every second year. As with SSWs, there is a coherent response through the depth of the middle atmosphere, with polar warming in the stratosphere and simultaneous polar cooling in the mesosphere, and a zonal mean zonal wind weakening in the stratosphere and mesosphere. However, the temperature and zonal wind changes associated with a hiccup seem to propagate upward, rather than downward as is observed for major SSWs by e.g. Matthias et al. (2012). This is argued to result from the fact the SPW propagation is not completely suppressed following hiccups, as it is for SSWs, and thus the middle atmosphere continues to be driven from the troposphere during hiccups.

Although enhanced PW 1 amplitudes are an inherent element of hiccups, hiccups are zonal-mean perturbations and are thus to be distinguished from local PW variations as discussed by Liu et al. (2001). Hiccups seem to dominate at high latitudes, and non-hiccup PW variations at mid latitudes. More study of these interesting events during the autumn transition would seem to be warranted, using global analysis data sets such as CMAM30 in conjunction with ground-based and satellite measurements.

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