

Attribution of climate extreme events

Article

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3	Attribution of climate extreme events
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5	Kevin E Trenberth ¹ , John T. Fasullo ¹
6	and Theodore G. Shepherd ²
7	
8	1. National Center for Atmospheric Research (NCAR)
9	P.O. Box 3000
10	Boulder CO 80307
11	
12	
13	Department of Meteorology, University of Reading,
14	Reading RG6 6BB, UK
15	
16	
17	Phone: (303) 497 1318
18	Fax: (303) 497 1333
19	email: <u>trenbert@ucar.edu</u>
20	
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37 There is a tremendous desire to attribute causes to weather and climate events 38 that is often challenging from a physical standpoint. Headlines attributing an event solely to either human-induced climate change or natural variability can be 39 40 misleading when both are invariably in play. The conventional attribution framework struggles with dynamically-driven extremes because of the small 41 42 signal-to-noise and often uncertain nature of the forced changes. Here, we 43 suggest that a different framing is desirable, which asks why such extremes 44 unfold the way they do. Specifically, that it is more useful to regard the extreme 45 circulation regime or weather event as being largely unaffected by climate 46 change, and question whether known changes in the climate system's 47 thermodynamic state affected the impact of the particular event. Some examples 48 briefly illustrated include "snowmaggedon" in February 2010, super storm Sandy 49 in October 2012, and super typhoon Haiyan in November 2013 and, in more detail, 50 the Boulder floods of September 2013, all of which were influenced by high sea 51 surface temperatures that had a discernible human component.

52

Weather and climate extremes happen all of the time, even in an unchanging climate. Yet there is a justifiably strong sense that some of these extremes are changing to be more frequent, and that the main reason is because of human-induced climate change. Indeed, the main way climate change is likely to be manifested on societies around the world is through changes in extremes. As a result, the scientific community faces an increasing demand for regularly updated appraisals of evolving climate conditions and extreme weather. Such information would be immensely beneficial for adaptation planning.

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61 The large-scale atmospheric circulation determines where it is dry, where it is wet, where 62 it is hot, and so on. A recent example is documented for the Pacific Northwest¹. Yet as 63 discussed below, in contrast to thermodynamic aspects of climate, forced circulation 64 changes in climate models can be very non-robust, and physical understanding of the 65 causes of these changes is generally lacking². Therefore, we suggest that separating out the 66 thermodynamic from dynamic effects may be a very fruitful way forward and result in a 67 different set of questions to be addressed. In turn, these provide a better basis for 68 communication of climate change to the public.

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70 Complexities associated with extreme event attribution

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The climate community has responded to the demand for timely information by attempting to perform attribution of climate extremes, both through the Intergovernmental Panel on Climate Change (IPCC)³ reports but most evident in closer to real time through an annual report, the most recent of which — "*Explaining extreme events of 2013 from a climate perspective*" — was published by the American Meteorological Society in September 2014⁴ and reported on several events from 2013. The question posed in each case was whether the likelihood or strength of the event was affected by anthropogenic climate change.

More generally, there are perhaps two main kinds of attribution performed. The first
relates the particular extreme event to the associated weather and weather patterns, and
this has been a useful and long-standing activity in climate science. We hear statements

like "the drought was caused by a blocking anticyclone"; or the "outbreak of tornadoes was
caused by a displaced and active storm track and jet stream"; "the flooding was caused by
El Niño", and so on. However, because these studies relate the event and phenomenon to
the weather situation or weather pattern, they are really a description of the event, not a
cause. As an explanation, the question should be, "why did that weather phenomenon
behave the way it did?" In particular, "what influences external to the atmosphere were
plaving a role and what climate factors were in play?"

90

91 The second kind of attribution relates especially to the objective of assessing the role of 92 human activities, and especially human-induced climate change in the event. Or perhaps

- 92 numan activities, and especially numan-induced climate change in the event. Or perhaps
 93 this might be generalized to a goal of assessing the role of external influences in the climate
 94 system on the event in question. So as well as human influences there may be influences
 95 from volcanic eruptions or the sun. However, results depend upon how questions are
 96 framed⁵⁻⁷.
- 90 97

98 The conventional approach to attribution of climate events is described by Stott et al.⁸. 99 Clearly, it is not possible to attribute a single climate extreme event, which by definition is 100 unique and which has a large element of chance in its occurrence, to a specific cause. Thus, 101 the approach is to characterize the event and ask (i) whether the likelihood or strength of 102 such events has changed in the observational record, and (ii) whether this change is

103 consistent with the anthropogenic influence as found in one or more climate models, and 104 thereby assess the "Fraction of Attributable Risk". The conventional approach has had

105 considerable success with extremes that are strongly governed by thermodynamic aspects

of climate change, especially those related to temperature. Since the attribution is
unconditional — the null hypothesis is that there is no effect of climate change — each

finding provides another independent line of evidence that anthropogenic climate changeis affecting climate extremes.

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111 However, the conventional approach is severely challenged when it comes to climate 112 extremes that are strongly governed by atmospheric circulation, including local aspects of 113 precipitation. The main reason is because changes in the atmospheric circulation related to 114 climate change are fairly small compared to natural variability, as has been shown 115 especially by several studies at NCAR using "large ensembles"^{9, 10} for 60 years into the 116 future. While large changes in atmospheric circulation can be readily apparent in a single 117 climate model run, they are not robust and change a lot in the next run or model. Indeed 118 what have often been interpreted as differences between models used in IPCC reports may 119 have arisen at least in part from natural variability. Hence forced circulation changes are 120 not well established and it is difficult to detect changes in circulation-related extremes in 121 observations because of small signal-to-noise. Thus, the anomalous weather pattern is 122 *always* the dominant influence for all short-lived events of a week or two, and this is generally true even for years or decades. Accordingly, the conventional approach to 123 124 extreme event attribution is rather ineffectual in cases that rely on the changed circulation, 125 with generally an inconclusive outcome. Even when a detectable anthropogenic influence is found in a model, the reliability of that finding cannot carry much weight. 126

128 Yet from a risk perspective, it is still important to assess the likely anthropogenic impact on

- such events. By starting from the null hypothesis of no climate change, the conventional
- 130 approach to extreme-event attribution has to re-establish an anthropogenic influence for
- each kind of event *ab initio*, which makes it inherently conservative and prone to "Type II
- 132 errors" which underestimate the true likelihood of the human influence^{5, 6}. Given that we
- have a large amount of confidence in many aspects of climate change³, it seems rather
- peculiar to ignore that prior knowledge in our assessment of climate events⁵.
- 135

136 As noted above the predictability of the dynamics is problematical² and it is mainly through 137 the thermodynamics that reliable statements can be made with confidence about the role 138 of climate change. In other words, changes in temperatures and temperature extremes 139 have a much more robust basis. That carries over to the atmospheric moisture amounts 140 through the Clausius-Clapeyron equation in which the water-holding capacity of the 141 atmosphere goes up exponentially at a rate of about 7% per degree Celsius. Indeed this is 142 what is observed in actual changes in moisture in the atmosphere over the oceans and 143 where surface moisture is not limited. But over land, especially in summer, water 144 availability is also a vital factor¹¹. Because the main rainfall almost always comes from 145 moisture convergence in the atmosphere, increased water vapor in the environment leads to more intense rains and a risk of flooding^{12, 13}, even if the total amount does not change 146 147 much. Moreover, it is in droughts that the extra heat from increased greenhouse gases 148 accumulates whereas more generally the presence of surface moisture adds an evaporative 149 cooling effect.

150

151 More fruitful scientific questions

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153 One can ask questions such as: What was the role of internal natural variability in setting up 154 the pattern? Studies to explore this question include those documenting the role of 155 anomalous SSTs and especially El Niño-Southern Oscillation (ENSO), the Pacific Decadal 156 Oscillation (PDO)¹⁴, the North Atlantic Oscillation (NAO), and other so-called modes of 157 variability. In general, it may not be possible to say much concrete about regional climate 158 change unless one can understand and predict such modes of variability and their statistics. 159 ENSO may be predictable for up to 18 months or so, some decadal predictability may exist 160 for decadal modes for several years¹⁵, but in general chaotic elements in the weather and 161 climate system preclude longer-term statements because these regional patterns are not 162 externally forced.

163

In most of these cases, the result of these studies will be a description of the large-scale
patterns, the anomalous SSTs, and the relationships between the atmospheric circulation,
storm tracks, blocking, temperatures and precipitation, and perhaps extremes. If

- 167 considered, this part will undoubtedly conclude that greenhouse gas forcing or aerosols
- played little or no role in the circulation changes, although sometimes claims otherwise aremade.
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- 171 Instead, with regard to climate change, the questions to be answered could be:
 - Given the weather pattern, how were the temperatures, precipitation, and associated impacts influenced by climate change?

174	• Given a drought, how was the drying (evapotranspiration) enhanced by climate				
175	change and how did that influence the moisture deficits and dryness of soils, and				
176	wild fire risk? Did it lead to a more intense and perhaps longer lasting drought, as				
177	is likely ^{16, 17} ?				
178	• Given a flood, where did the moisture come from? Was it enhanced by high ocean				
179	temperatures that may have had a climate change component?				
180	• Given a heat wave, how was that influenced by drought, changes in precipitation				
181	(absence of evaporative cooling from dry land), and extra heat from global				
182	warming?				
183	 Given extreme snow, where did the moisture come from? Was it related to higher 				
184	than normal SSTs off the coast or farther afield?				
185	• Given an extreme storm, how was it influenced by anomalous SSTs and ocean heat				
186	content, anomalous moisture transports into the storm, and associated rainfall and				
187	latent heating? Was the storm surge worse because of high sea levels?				
188	In other words, given the change in atmospheric circulation that brought about the event,				
189	how did climate change alter its impacts?				
190	now the chinate change after its impacts.				
190	To summarize therefore, at least in the present state of knowledge, in our view a more				
192	fruitful and robust approach to climate extreme-event attribution is to regard the				
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196	temperature, or atmospheric moisture content), concerning which there is a reasonably				
197	high level of confidence. Such questions immediately lead to a physically-based approach				
198	that is strongly linked to the event in question. Although such questions differ from the				
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201	questions are posed differently, their answers have a different meaning and focus more on				
202	impacts. That is still useful information.				
203	Letter the second se				
204	Frequently, the main influences identified external to the atmosphere are the changes in				
205	SSTs. Of course the SSTs at any time have a large natural variability component but the				
206	values are very often inflated over what they would have been without climate change. The				
200	latter accounts for order 0.6°C increase since the 1950s globally ³ and because this				
207	inherently occurs on multi-decadal timescales with a memory through the ocean heat				
209	content, it is always present. Small increases in atmospheric moisture associated with				
210	such an increase in SSTs are of order 5% (ref. 19). We also assume that the global sea level				
211	rise for the past century of 19 cm (1901-2010) ³ reflects the underlying human influence				
212	while regional values vary owing to atmospheric circulation variations.				
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044					

214 Some examples

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- 216 The following examples briefly discuss a few events whose attribution has not been
- addressed but which received an enormous amount of media attention.
- 218

Consider the widely publicized event called "snowmaggedon" in Washington, DC²⁰ in February 5-6, 2010. Key features included (1) it was winter and there was plenty of cold continental air; (2) there was a storm in the right place; and (3) the unusually high SSTs in the tropical Atlantic Ocean (1.5°C above normal) led to an exceptional amount of moisture flowing into the storm, which resulted in very large snow amounts. It is this last part that then relates in part to anomalous external influences on the atmosphere through the effects of climate change on SSTs. So while internal variability is also playing a role, the extremes

- 226 are magnified by climate change.
- 227

228 Super-storm Sandy caused tremendous damage when it made landfall on the New Jersey 229 coast and New York area on 30 October 2012. It began as a hurricane, with peak strengths 230 making it a category 3 storm, and it caused substantial impacts in the Caribbean²¹. 231 However, as it moved north it became a hybrid storm before making landfall, which greatly 232 increased its overall size. The worst problems on the Jersey Shore were caused by the 233 strong winds and the associated storm surge, leading to extensive flooding. Farther inland, 234 heavy precipitation was also a major problem. Widespread damage from flooding streets, 235 tunnels and subway lines and cutting power in and around New York City led to damages exceeding \$65 billion (2013 USD)²². Fortunately, the storm was well forecast a week in 236 237 advance in particular by European Centre for Medium-Range Weather Forecasts (ECMWF), 238 including the very unusual left hook turn in the track toward the coast. ECMWF has 239 performed a number of experiments on the performance and behavior of the storm using 240 an ensemble of forecasts five days in advance of landfall²³. Of note is that they swapped the 241 observed SSTs for climatological values that average 1 to 1.5°C cooler in a broad strip along 242 the coast. Only small changes occurred to the track of the storm but the observed SSTs led 243 to a bigger more intense storm, stronger winds and greater precipitation. The average 244 depth of the storm was increased by 7.6 hPa, the wind speeds were increased by 3.6 m s⁻¹ 245 and the precipitation increased by 35%. Moreover the storm was riding on sea levels that 246 were higher by about 19 cm due to global warming. Although perhaps only one-half to 247 one-third of the SST increase can be blamed on global warming from human activities, it is 248 readily apparent that the storm surge and associated damage was considerably influenced 249 by climate change. It is quite possible that the subways and tunnels may not have flooded 250 without the warming-induced increase in sea level and storm intensity and size, putting a 251 potential price tag of human climate change in this storm in the tens of billions of dollars.

252

253 Another example is super typhoon Haiyan²⁴ that devastated the Philippines in early 254 November 2013. It is among the largest and most intense typhoons on record with 255 estimated surface pressure in the eye down to 895 hPa and 1-minute sustained winds of 315 km/h (gusts to 378 km/h) on November 7, 2013 which makes it the strongest 256 257 recorded storm to ever strike land (Fig. 1). It has been called a "category 6" storm²⁵. The 258 ocean heat content (OHC) and sea level in the region had increased a great deal since 1993 259 and especially since 1998 in association with the PDO negative phase^{26, 14} so that since 260 1993 the linear trend in sea level was over 16 mm/yr as compared to the global trend of 261 3.3 mm/yr (Fig. 1; using data from AVISO²⁷). Consequently, as the typhoon approached the 262 Philippines, it was riding on very high SSTs with very deep support through the high OHC, 263 and the strong winds and ocean mixing did not cause as much cooling as would normally 264 be experienced, very likely helping the storm to maintain its tremendous strength²⁵.

- 265 Moreover, the storm surge was undoubtedly exacerbated considerably by the sea levels 266 that were some 30 cm above 1993 values. Although natural variability through the PDO 267 played a major role, there is also a global component through increased OHC from the Earth's energy imbalance²⁸.
- 268
- 269

270 A conclusion then, is that while climate change is no doubt altering the atmospheric 271 circulation, the change is relatively small and can only be discerned from a very large 272 ensemble of model runs. That sets the change in odds. But for any event, the particular 273 character of that storm or synoptic situation and natural variability rule, while 274 thermodynamic effects increase the impacts.

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- 276

A case in point: the Boulder floods of September 2013

- 277 278 One of the recent BAMS studies is of the major floods centered in Boulder in September 279
- 2013²⁹. The unfortunate headline of the news release which carried over in large letters to 280 the front page of the local newspaper in Boulder (*Daily Camera* 30 Sept 2014)³⁰ was:
- 281 "Climate change not to blame for 2013 Colorado floods". The paper summary was: "The
- 282 probability for an extreme five-day September rainfall event over northeast Colorado, as
- 283 was observed in early September 2013, has likely decreased due to climate change."
- 284
- 285 As noted above, in any weather event like this, the weather situation is always the main 286 player in the developments but that in and of itself says nothing about the role of climate 287 change. The study did note the importance of having abundant moisture in the region in 288 order to produce high rainfall amounts. But it did not include an assessment as to where 289 the moisture came from. In Denver the three highest total column water vapor amounts 290 ever recorded for September (since 1956) occurred on 12-13 September 2013 (as high as 291 34 mm). [This may not seem huge but recall Denver is a mile (>1600 m) above sea level].
- 292

293 It so happens that the SSTs off the West Coast of Mexico, south of Baja, west of Guadalajara, 294 were over 30°C and more than 1°C above normal in August 2013 (see Fig. 2), which made it 295 the hottest spot for the ocean in the western hemisphere. An incredible 75 mm of total 296 column water vapor was recorded in the atmosphere in that region by NASA satellites. The 297 high SSTs led to the large-scale convergence of moisture flowing into the region that was 298 siphoned north by a very unusual synoptic situation leading to a river of atmospheric 299 moisture flowing into Colorado (Figs. 3 and 4). The heaviest moisture convergence into 300 eastern Colorado coincided with the times of the plots in Fig. 4, while breaks in the rain 301 occurred in between times in some locations (Fig. 3). After that river shut off, twin tropical 302 storms formed both sides of Mexico: Manuel (to the west)³¹ and Ingrid (to the east)³² (Fig. 303 5) that formed a double whammy for Mexico and led to hundreds of deaths, tens of 304 thousands evacuated, tens of thousands of homes damaged and billions of dollars of 305 damage^{31, 32}.

306

307 Experience with simulations of atmospheric rivers at NCAR has shown that quite high

- 308 resolution of order one quarter the size of that used by Hoerling et al.²⁹ is required to
- 309 perform realistic simulations, and coarse resolution can lead to discrepancies in the
- 310 atmospheric moisture amounts simulated. Hoerling et al.²⁹ is likely correct in concluding

311 that the set up of weather systems was rare, and climate change is not a discernible factor 312 in that respect. However, one cannot realistically do attribution using a model that does 313 not have the requisite spatial resolution to represent the relevant dynamics and is unable 314 to replicate the event in question, and it is desirable to use more than one model. Moreover, 315 the extremely high SSTs and record water vapor amounts that accompanied the event and 316 were instrumental in its evolution likely would not have occurred without climate change. 317 Even a 10% increase in moisture in the atmosphere becomes concentrated when focused 318 by topography, and further amplified when on the ground as water drains into channels 319 and rivers.

320

321 SSTs have been as high in the past in this region west of Mexico (Fig. 2), but in previous 322 cases they were part of a much larger-scale pattern associated with El Niño events such as 323 in 1997-98, 2004-5, 2009-10. What seems to be unique in 2013 is that this was the 324 warmest spot in the western hemisphere and hence this was the preferred location for low 325 pressure to form and low-level wind convergence, which brought large amounts of 326 moisture into the region. Hence there was clearly an internal variability component to the 327 patterns of SST, but at the same time, the overall increase of global SST associated with 328 global warming that occurs on multi-decadal time scales was also a factor. In Fig. 2 this 329 warming is especially evident after 1997 when SSTs average >0.4°C above the mean.

330

331 Concluding remarks

332 333 We have suggested four climate events in which anomalously high SSTs played a key role in 334 feeding moisture into storms, helping to intensify the storm and causing heavy rains. Of 335 course the SSTs at any time have a large natural variability component on top of the 336 human-induced warming. Associated increases in atmospheric moisture are of order 5% 337 (ref. 19), which is magnified as moisture converges into a storm, further magnified as the 338 storm intensifies, and concentrated when it runs up against orography, and in streamflows. 339 Hence even though natural variability always dominates in such storms, the research task 340 is to refine the above estimates and properly account for the human component as at least 341 a partial explanation for why extremes are being magnified.

342

There are many other possible examples. For instance, conflicting results occurred for the Russian heat wave in 2010^{33, 34} that were to some extent reconciled³⁵ by recognizing that each study was about different aspects. One study³³ focused on dynamical aspects while the other³⁴ was much more about the record high temperatures and thermodynamic aspects. Indeed, the record high SSTs and associated nearby rains in several areas

- 348 (including the Pakistan floods) were factors external to the atmosphere³⁶ that altered
- 349 teleconnections into the Russian region.
- 350
- Another very recent example is the California drought beginning in 2012. While one
- 352 study³⁷ found no significant trends in winter precipitation in recent decades, another³⁸
- pointed out the critical role of the record high annual mean temperatures in combination
- with record low annual mean precipitation for 2013 which led to increased
- evapotranspiration and more intense drought that had impacts on water shortages,
- vegetation and agriculture, and increased wildfire risk. The odds of this combination have

increased with human-induced climate change and that anthropogenic warming has
increased drought risk³⁹. Again these two studies are consistent with the view that the
atmospheric circulation changes are not the dominant factor as far as the climate change
aspects are concerned.

361

362 In addition there were 22 articles in the special BAMS issue⁴. Over two-thirds focused on 363 aspects of atmospheric circulation and generally concluded there was little or no influence 364 of anthropogenic climate change. None of the eight studies that dealt with heavy rains or 365 snows performed an analysis of the moisture budget and where the abundant water came 366 from. Of the three studies that analyzed the California drought, none dealt with land 367 surface effects or changes in evapotranspiration associated with climate change. In 368 contrast, for the Australian heat wave, the role of decreased soil moisture and drought were noted for their effects on temperatures. The summary Table focused on columns 369 370 about anthropogenic influences vs no such influence detected.

371

There are many other examples of studies that might ask more targeted questions that better serve societal needs. Reframing the null hypothesis and asking different questions would help advance the understanding. Because global warming is real and present, it is not a question as to whether it is playing a role but what that role is. A Bayesian approach takes these "priors" appropriately into account. We can talk about these effects in terms of changing odds, as many have done. But we can also talk about them in physical terms.

378

379 To implement the approach suggested here, one needs to be able to simulate the event in 380 question (perhaps with short-term forecasts, as in the Sandy example, or with simulations 381 whose circulation is constrained in some manner), and then assess the impact of known 382 anthropogenic changes in the climate system's thermodynamic state, including changed 383 amounts of water vapor. Only in this way can the necessarily conditional nature of the 384 event attribution be implemented. Necessarily this addresses only a subset of the possible 385 influences and conditions. However, with the underlying assumption that the phenomena 386 are largely unchanged, this is perhaps as much as can be expected for event attribution. 387 After all a single extreme event is not a repeatable thing.

388

389 The climate is changing: we have a new normal! The environment in which all weather 390 events occur is different than it used to be. All storms, without exception, are different. 391 Even if most of them look just like the ones we used to have, they are not the same. 392 However, we cannot sort out these questions of degree without a large ensemble of model 393 simulations, particularly for events as rare as the Boulder floods, and the kinds of models 394 that can be run in such a way are often incapable of simulating the event in question and 395 thus lack physical credibility. We argue that under such conditions it is better for event 396 attribution to focus not on the synoptic event, but rather on the influences of the changed 397 large-scale thermodynamic environment on the extremes and temperatures and moisture 398 associated with the event.

400	References		
401			
402	1.	Johnstone, J. A., & Mantua, N. J. Atmospheric controls on northeast Pacific temperature	
403		variability and change, 1900–2012. <i>PNAS</i> , doi:10.1073/pnas.1318371111 (2014).	
404	2.		
405		projections. <i>Nature Geo.</i> , 7 , 703-708, doi: 10.1038/NGE02253 (2014).	
406	3.	IPCC (Intergovernmental Panel on Climate Change): <i>Climate change 2013: The Physical science</i>	
407	0.	<i>basis.</i> Fifth Assessment report of the IPCC [Stocker, T. F. et al. eds)]. Cambridge Univ. Press.	
408		1535 pp. (2013).	
409	4.	Herring, S. C., Hoerling, M. P., Peterson T. C. & Stott, P. A., Eds. Explaining extreme events of	
410	т.	2013 from a climate perspective. <i>Bull. Amer. Meteor. Soc.</i> , 95 , (9) S1-S96 (2014).	
411	5.	Trenberth, K. E. Attribution of climate variations and trends to human influences and natural	
412	5.	variability. Wiley Interdisciplinary Reviews (WIREs) Climate Change 2, 925-930, doi:	
412			
413	6	10.1002/wcc.142 (2011a).	
	6.	Trenberth, K. E. Framing the way to relate climate extremes to climate change. <i>Climatic Change</i> ,	
415	-	115 , 283-290, doi: 10.1007/s10584-012-0441-5 (2012).	
416	7.	Wallace, J. M. Weather- and climate-related extreme events: Teachable moments. <i>Eos</i> , 93 , 11,	
417	0	13 March, 120-121 (2012).	
418	8.	Stott P., et al. Attribution of weather and climate-related extreme events, Chapter 12 in " <i>Climate</i>	
419	0	Science Serving Society", Asrar, G. R. & Hurrell, J. W., Eds. Springer, 307-337 (2013).	
420	9.	Deser, C., Phillips, A., Bourdette, V. & Teng, H. Y. Uncertainty in climate change projections: the	
421		role of internal variability. <i>Clim. Dyn.</i> 38 , 527–546 (2012).	
422	10.	Deser, C., Phillips, A. S., Alexander, M. A. & Smoliak, B. V. Projecting North American climate	
423		over the next 50 years: uncertainty due to internal variability. <i>J. Clim.</i> , 27 , 2271–2296 (2014).	
424	11.	Trenberth, K. E. Changes in precipitation with climate change. <i>Climate Research</i> , 47 , 123-138,	
425		doi:10.3354/cr00953 (2011).	
426	12.	Trenberth, K. E. Atmospheric moisture residence times and cycling: Implications for rainfall	
427		rates with climate change. <i>Climatic Change</i> , 39 , 667–694 (1998).	
428	13.	Trenberth, K. E., Dai, A., Rasmussen, R. M. & Parsons, D. B. The changing character of	
429		precipitation. <i>Bull. Amer. Meteor. Soc.</i> , 84 , 1205–1217 (2003).	
430	14.	Trenberth, K. E., Fasullo, J. T., Branstator, G. & Phillips, A. S. Seasonal aspects of the recent pause	
431		in surface warming. <i>Nature Climate Change</i> , 4 , 911-916, doi:10.1038/NCLIMATE2341 (2014c).	
432	15.	Branstator, G. & Teng, H. Is AMOC more predictable than North Atlantic heat content? J. Clim.,	
433		27 , 3537-3550 (2014).	
434	16.	Seneviratne, S. I., et al. Investigating soil moisture-climate interactions in a changing climate: A	
435		review, <i>Earth Sci. Rev.</i> , 99 , 125–161 (2010).	
436	17.	Trenberth, K. E., Dai, A., van der Schrier, G., Jones, P. D., Barichivich, J., Briffa, K. R. & Sheffield, J.	
437		Global warming and changes in drought. <i>Nature Climate Change</i> , 4 , 17-22,	
438		doi:10.1038/NCLIMATE2067 (2014a).	
439	18.	Lyons, L., Discovery or fluke: Statistics in particle physics. <i>Physics Today</i> , 65 , (7), 45-51 (2012).	
440		Trenberth, K. E., Fasullo, J. & Smith, L. Trends and variability in column-integrated atmospheric	
441		water vapor. Clim. Dynam., 24, 741-758 (2005).	
442	20.	http://en.wikipedia.org/wiki/February_5%E2%80%936,_2010_North_American_blizzard	
443		Blake, E. S., Kimberlain, T. B., Berg, R. J., Cangialosi, J. P. & Beven II, J. L. Tropical cyclone report	
444		hurricane Sandy. Rep. AL182012, National Hurricane Center, 157 pp (2013).	
445	22.	http://en.wikipedia.org/wiki/Hurricane_Sandy	
446		Magnusson, L. Bidlot, J-R., Lang, S. T. K., Thorpe, A., Wedi, N. & Yamaguchi, M. Evaluation of	
447	_01	medium-range forecasts for hurricane Sandy. <i>Mon. Wea Rev.</i> 142 , 1962-1981 (2014).	
448	24	http://en.wikipedia.org/wiki/Typhoon_Haiyan	
449		Lin, I-I., Pun, I-F., & Lien, C-C. 'Category-6' Supertyphoon Haiyan in global warming hiatus:	

- 450 Contribution from subsurface ocean warming. *Geophys. Res. Lett.* **41**, 8547–8553,
- 451 doi: 10.1002/2014GL061281.
- 452 26. Trenberth, K. E. & Fasullo, J. T. An apparent hiatus in global warming? *Earth's Future*. 1, 19-32, doi: 10.002/2013EF000165 (2013).
- 454 27. <u>http://www.aviso.altimetry.fr/en/data/products/ocean-indicators-products/mean-sea-level.html</u>
- 456 28. Trenberth, K. E., Fasullo, J. T. & Balmaseda, M. A. Earth's energy imbalance. *J. Clim.*, 27, 3129457 3144, doi: 10.1175/JCLI-D-13-00294 (2014b).
- 458 29. Hoerling, M., et al. Northeast Colorado extreme rains interpreted in a climate change context [in
 459 "Explaining Extremes of 2013 from a Climate Perspective"]. Bull. Amer. Meteor. Soc., 95 (9), S15460 S18 (2014).
- 461 30. <u>http://www.dailycamera.com/news/boulder-flood/ci_26626817/boulder-researcher-2013s-flood-triggering-rains-not-caused</u>
- 463 31. <u>http://en.wikipedia.org/wiki/Hurricane_Manuel</u>
- 464 32. http://en.wikipedia.org/wiki/Hurricane_Ingrid
- 33. Dole, R., Hoerling, M., Perlwitz, J., Eischeid, J., Pegion, P., Zhang, T., Quan, X.-W., Xu, T. & Murray,
 D. Was there a basis for anticipating the 2010 Russian heat wave?, *Geophys. Res. Lett.*, 38,
 L06702, doi:10.1029/2010GL046582 (2011).
- 468 34. Rahmstorf, S., & Coumou, D. Increase of extreme events in a warming world, *Proc. Natl. Acad.* 469 *Sci.* U. S. A., **108**(44), 17905–17909 (2011).
- 35. Otto, F. E. L., Massey, N., van Oldenborgh, G. J., Jones, R. G. & Allen, M. R. Reconciling two
 approaches to attribution of the 2010 Russian heat wave, *Geophys. Res. Lett.*, 39, L04702,
 doi:10.1029/2011GL050422 (2012).
- 36. Trenberth, K. E., & Fasullo, J. T. Climate extremes and climate change: The Russian heat wave
 and other climate extremes of 2010, *J. Geophys. Res.*, **117**, D17103, doi:10.1029/2012JD018020
 (2012).
- 37. Seager, R, et al. Causes and Predictability of the 2011-14 California Drought. Available at
 cpo.noaa.gov/sites/cpo/MAPP/Task%20Forces/DTF/californiadrought/california_drought_re
 port.pdf (2014).
- 38. Diffenbaugh, N. S., Swain D. L. & Touma, D. Anthropogenic warming has increased drought risk
 in California. *PNAS*, **112**, 3931-3936, doi 10.1073/pnas.1422385112 (2015).
- 481
- 482 Correspondence and requests for materials should be addressed to K.E.T.
- 483

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489 Author contributions

- K.E.T led the writing of the paper and conceived of the paper and figures. J.T.F analyzed some dataand contributed to 2 figures. All authors contributed to writing the manuscript.
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- 499

Fig. 1. Haiyan and sea level: Linear sea level trends from August 1993 to July 2013 in mm/y,
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Fig. 2. August SSTs for 12-20°N 110-100°W, just west of Mexico. The mean value is
28.92°C for 1982-1999. The last value is for 2013.

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Fig. 3. Water vapor channel imagery for 18:45 GMT on 12 September 2013 showing the
extensive water vapor and associated activity both west of Mexico and in the Caribbean Sea
and the river of moisture from south of Baja, Mexico to eastern Colorado. Courtesy
weathertap.com, published with permission.

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513 Fig. 4. GOES 13 imagery from the 6.5 micron water vapor channel at 0545 GMT on 11 (top)

and 13 (bottom) of September 2013. The water vapor is in the mid to upper troposphere

and the brighter the imagery, the more saturated the air. Courtesy Axel Graumann,

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