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CO₂ emission estimation in the urban environment: Measurement of the CO₂ storage term

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HIGHLIGHTS

- Diurnal and seasonal CO₂ storage (ΔC₅) cycles in an urban street canyon reported.
- Effect of sensor position, sampling rate, and data processing on ΔC₅ assessed.
- Method of calculating ΔC₅ from measurements at one height shown to be invalid.
- Spectral analysis of ΔC₅ suggests an inverse power law relationship with frequency.
- Storage and vertical flux compared across a full seasonal cycle.

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ABSTRACT

Eddy covariance has been used in urban areas to evaluate the net exchange of CO₂ between the surface and the atmosphere. Typically, only the vertical flux is measured at a height 2–3 times that of the local roughness elements; however, under conditions of relatively low instability, CO₂ may accumulate in the airspace below the measurement height. This can result in inaccurate emissions estimates if the accumulated CO₂ drains away or is flushed upwards during thermal expansion of the boundary layer. Some studies apply a single height storage correction; however, this requires the assumption that the response of the CO₂ concentration profile to forcing is constant with height. Here a full seasonal cycle (7th June 2012 to 3rd June 2013) of single height CO₂ storage data calculated from concentrations measured at 10 Hz by open path gas analyser are compared to a data set calculated from a concurrent switched vertical profile measured (2 Hz, closed path gas analyser) at 10 heights within and above a street canyon in central London. The assumption required for the former storage determination is shown to be invalid. For approximately regular street canyons at least one other measurement is required. Continuous measurements at fewer locations are shown to be preferable to a spatially dense, switched profile, as temporal interpolation is ineffective. The majority of the spectral energy of the CO₂ storage time series was found to be between 0.001 and 0.2 Hz (500 and 5 s respectively); however, sampling frequencies of 2 Hz and below still result in significantly lower CO₂ storage values. An empirical method of correcting CO₂ storage values from under-sampled time series is proposed.

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1. Introduction

Human activity in urban areas is responsible for 30–40% of directly emitted anthropogenic greenhouse gases (Satterthwaite, 2008) and 70% of total emissions of CO₂ (Canadell et al., 2009). Several urban studies have assessed net emissions at the local scale (10²–10⁴ m) using the eddy covariance (EC) method (Griffith and Christen, 2012; Christen, 2014). Here net ecosystem exchange (NEE) is calculated as the sum of the CO₂ exchanges through the sides of a notional volume of air and the change in the amount of CO₂ stored (ΔC₅) within the volume (Aubinet et al., 2005). In practice, many monitoring sites (Griffith et al., 2002; Baldocchi, 2003; Velasco and Roth, 2010) approximate NEE as the net flux into or out of the top of the volume of interest, assuming horizontal homogeneity (horizontal advection into the volume is equal to...
transport out of the volume) and fully turbulent conditions with negligible storage. In urban environments during stable atmospheric conditions the $\Delta C$ term will be non-negligible (Helfter et al., 2011); for example, $\Delta C$ was found to be five times the magnitude of the turbulent vertical flux term ($FC_{CO2}$) close to dawn and dusk in suburban Vancouver, Canada (Crawford and Christen, 2014). Other urban studies found $\Delta C$ to be smaller, but still significant, with maximum $\Delta C$ values 11% and 22% of the magnitude of $FC_{CO2}$ in Edinburgh, Scotland (Nemitz et al. 2002) and Basel, Switzerland (Feigenwinter et al. 2012), respectively. Similarly, in rural environments horizontal advection may be non-negligible (e.g., Aubinet et al., 2003). This paper focuses on the methodological considerations when assessing $\Delta C$ from a vertical profile in an urban environment. Future manuscripts will address horizontal variation and advection.

Of three studies discussed above (Vancouver, Edinburgh, Basel) with reported urban CO2 storage values, only one (Vancouver, Crawford and Christen, 2014) presented values derived from a dataset of longer than one month; however, both these values and those reported for Edinburgh by Nemitz et al. (2002) were calculated based on the assumption of a constant relation between carbon dioxide concentration ($[CO2]$) measured above the blending height and the concentration in the street canyon. In contrast, for Basel, Feigenwinter et al. (2012) did not make this assumption and reported $\Delta C$ calculated from $[CO2]$ at ten levels; however, the results are only for one month (15th June to 15th July 2002). There is therefore scope to improve not only the understanding of the processes affecting CO2 storage over a greater range of meteorological and anthropogenic conditions, but also to develop recommendations for future measurement programmes.

The objective of this paper is to evaluate potential approaches for such studies, illustrated with examples from, and analysis of, high temporal resolution data collected at 10 locations from 6.5 to 46.4 m above ground level between 2011 and 2014 at King’s College London, in Central London, UK. The paper is organized as follows. In the rest of this Section 1 we provide a brief background of how CO2 storage is calculated and then (Section 2) a discussion of the methods used in this paper. This is followed by an exploration of the temporal variation of CO2 storage (Section 3) and the relation between measured CO2 storage and anthropogenic and natural factors in a highly urbanised environment (Section 4). The required number and placement of sample points for CO2 storage measurements in a deep urban street canyon is addressed (Section 5) and the effect of sensor response and sampling interval on calculated CO2 storage is tested (Section 6). At the processing stage, three different temporal and spatial interpolation methods are evaluated against measured data (Section 7). Finally, the impact of CO2 storage calculated by two different methods on the turbulent vertical CO2 flux is assessed. The Supplementary material (noted by S.1 to S.9) includes the notation with corresponding units (S.1) used in the text, further information on CO2 storage calculation (S.2), previous CO2 storage studies (S.3), equipment (S.4), meteorological characteristics (S.5), example time series (S.6), variations of CO2 with friction velocity, wind direction and height above ground level (S.7, S.8) and further references (S.9). The online version of this paper provides the figures in colour.

1.1. Calculation of CO2 storage

There are two main approaches to calculating CO2 storage flux density, i.e., the rate of change of CO2 per unit area below the Eddy Covariance (EC) measurement height. Here they are referred to as the ‘single height’ and the ‘profile’ approaches. For a brief discussion of the theory and some reported results, see Supplementary material S.2 and S.3, respectively. For a more in depth discussion of the theoretical considerations regarding CO2 storage the reader is referred to Finnigan (2006) and subsequent discussion (Kowalski, 2008; Finnigan, 2009).

The approach to calculate CO2 storage depends upon the number of vertical locations at which CO2 concentration ($[CO2]$), the symbol $f$ is used to indicate concentration) data are collected. In the first approach, ‘single height’ CO2 storage ($\Delta CS$) is calculated from $[CO2]$ data at one location, usually by eddy covariance equipment in the inertial sub-layer (Nemitz et al., 2002; Crawford and Christen, 2014). In the second approach, the ‘profile’ method, $\Delta C$ is calculated from data collected at multiple heights ($\Delta CPS$). The profile method uses a vertical $[CO2]$ profile at heights $z_i$, which is generally measured by cycling through all the sample locations within a set time period with a data-logger controlled valve array (Xu et al., 1999; Mölder et al., 2000; Vogt et al., 2006; Hutyra et al., 2008). This cycle period may not be the same as the averaging period used in the $\Delta CPS$ calculation. For example, measurements collected with a sampling interval, $t_i$, of 2 Hz for 75 s at 8 heights, giving a full profile cycle every 10 min, may be used to calculate $\Delta CPS$ with an averaging period ($T$) of 30 min. The storage is calculated as the sum of the changes in time averaged concentration ($FC_{CO2}$) between time $t_1$ and $t_2$ and $t_1$ and $t_2$ for each location ($i$) in the profile, weighted by the vertical span, $\Delta z_i$, over which each profile measurement is considered to be representative and divided by the averaging period ($T$), which can be expressed as (modified from Aubinet et al., 2005):

$$\Delta CPS = \frac{1}{T} \sum_{i} \left( \frac{[CO2]_{t_1} - [CO2]_{t_2}}{\Delta z_i} \right) \Delta z_i$$  \hspace{1cm} (1)

If the measurements at each height are not made concurrently, $[CO2]$ at each height may first be interpolated in time to generate instantaneous profiles from which $\Delta CPS$ can be calculated, though this is neglected in some cases (Iwata et al., 2005). The impact of interpolation on calculated $\Delta CPS$ is discussed further in Section 7.

The single height method is a simplification of the profile method to one height, which is usually the height of the eddy covariance equipment. The change in $[CO2]$ with time at $z_h$ ($\Delta [CO2]$/dt) is weighted by the vertical distance from the ground to the measurement point ($z_h$). The single height CO2 storage ($\Delta CSS$) is given by (modified from Nemitz et al., 2002):

$$\Delta CSS = \frac{\Delta [CO2]_{z_h}}{dt} \Delta z_h$$  \hspace{1cm} (2)

As the data are continuous, the change in the instantaneous CO2 concentration with time ($\Delta [CO2]$/dt) can be used instead of the change in the time averaged CO2 concentration with time ($FC_{CO2}$/dt), though it may still be advisable to average in time to reduce measurement noise (Finnigan, 2006).

The $\Delta CSS$ calculation assumes any change in $[CO2]$ below the measurement height results in a change of equivalent magnitude at the measurement height. This assumption appears not to be supported by any evidence in the literature; reported diurnal cycles of CO2 mixing ratios in the roughness sub-layer over both rural (e.g., Xu et al., 1999) and urban (e.g., Lietzke and Vogt, 2013) surfaces are known to vary with height. This problem is particularly acute during periods of low turbulence, such as at night or during cold weather, where measurements above the surface layer may become decoupled from processes near ground level (Helfter et al., 2011).

If temporal variability is large compared to the spatial variability, the single height method may provide a more accurate measure of storage than the profile method as the maximum data availability at each sample location for the latter may be $1/k$ of the total time.
series, where $k$ is the number of heights. Data availability may be improved by installation of multiple gas analysers (Simpson et al., 1998; Siebicke et al., 2010); however, this introduces the problem of ensuring that measurements are comparable between heights. If horizontal variability of $[\text{CO}_2]$ is high, measurements made at one particular height in multiple locations may be averaged prior to integration over the vertical extent of the volume of interest (e.g., Crawford and Christen, 2014). In this paper data from both continuous (two heights, $z_i = 20.5, 46.4$ m above the Strand street canyon (hereafter referred to as the Strand canyon) ground level (a.g.l.)) and switched (10 heights, $z_i = 6.5, 9.5, 12.5, 16.0, 19.5, 32.3, 33.7, 36.0, 39.3, 46.4$ m a.g.l.) $\text{CO}_2$ concentration profiles are reported and compared.

2. Methods

2.1. Site

The study area KS (red dot in Fig. 1a) is situated on the Strand, on the north bank of the River Thames in the centre of the Central Activity Zone in London, UK. The Central Activity Zone is the political, financial, retail and cultural ‘hub’ of London and, to a lesser extent, the UK. It contains nearly 30% of all of London’s jobs (GLA, 2008) and is heavily influenced by commuter traffic: 1.16 million people passed through the Central London cordon (Fig. 1a) during the 07:00 to 10:00 BST (British Summer Time, GMT +1, in effect from the last Sunday of March to the last Sunday of October) morning peak period on an average weekday in 2011, with 18% of the journeys made via a motorised vehicle (DfT, 2013). The relatively small nocturnal population results in lower demands for heating compared to residential areas (Velasco and Roth, 2010; Iamarino et al., 2012; Ward et al., 2015).

Sources of $\text{CO}_2$ in the vicinity of KS include rooftop emissions due to burning of natural gas for heating and venting of high-$\text{CO}_2$ air from air conditioning ducts (Fig. 1b, c; see also Kotthaus and Grimmond, 2012). Chamber measurements suggest the river (not shown) is not a significant source and there are no major power stations or other industrial activity nearby. Land cover characteristics for the EC source areas of the KSS tower (40 m to the east of KSSW) are (% plan area): roads (43%), buildings (38%), water (14%), vegetation, including trees, (5%), trees (2%) (Kotthaus and Grimmond, 2014b). The predominantly mature London Plane trees ($\text{Platanus acerifolia}$) line the Strand canyon and other major roads nearby. Whilst photosynthetic uptake of $\text{CO}_2$ does occur, it is likely to be an insignificant control on the vertical $\text{CO}_2$ flux or the $\text{CO}_2$ storage within the monitored street canyon. Ward et al. (2015) found the role of vegetation to be negligible at this site. The street canyon is symmetric and ~31 m deep with a smaller (~18 m)
obstacle (St Mary-le-Strand church, lower left, Fig. 1c) in the centre, which could be considered to form two asymmetric sub-canyons for part of the length of the Strand. The height:width ratio varies from 0.74 to 1.28 with a two-lane, one-way street running east-northeast to west-southwest carrying 16,000–19,000 vehicles (predominantly cars) per day (automated traffic count, Strand, 2013/189 to 196). Throughout this paper we use date format YYYY/DDDD = Year/Day of Year. Traffic speed is slow, ranging from a peak average speed of 45 km h⁻¹ between 03:00 and 05:00 BST to a minimum of 26 km h⁻¹ at the beginning of the evening rush hour peak in traffic volume (18:00 BST). Considerable congestion, particularly in the vicinity of two bus stops (located close to the measurement profile) is observed at this time.

2.2. Equipment and data set

Equipment installed on a guyed tower (13.31 m; T-35H, Alum Tower Inc., USA) is referred to here as the KSSW site (Fig. 1c). The second site, KSNW (Fig. 1c) consists of a pneumatic mast (<9 m, CQST9-6/HP, Clark Mast, Belgium) and air inlets for a switched /C0 profile on the exterior wall of a building extending down to 6.5 m above street level. Eddy covariance equipment, including fast response open-path infrared gas analysers (LI7500A, LI-COR Biosciences, USA), installed at KSSW and KSNW, have a horizontal separation of 70 m and are 46.4 and 20.5 m above ground level respectively. Note, as measurements are made within the street canyon, all instruments reported in this study are height above the Strand (14.5 m above sea level). Other papers presenting data from this site (e.g., Kotthaus and Grimmond, 2014b; Ward et al., 2015) reported elevations relative to the mean sea level height above sea level for the measurement footprint (10.6 m above sea level). In this paper, the former are used within the calculation of CO2 storage and the latter are used in stability and vertical CO2 flux.

As CO2 storage and vertical flux involve processes at the micro to local-scale, the air sampling heights (6.5 ≤ z ≤ 46.4 m a.g.l) were chosen to span from within the urban canyon to the top of the eddy covariance tower (z0, height A, Fig. 1d); in other words, within the roughness sublayer and inertial sublayer. According to a commonly used rule of thumb, the base of the inertial sublayer is 2–3 times the mean roughness element height (Roth, 2000; Grimmond et al., 2004). With a mean element height, z0, within 500 m of KSSW and KSNW of 21.7 m and 21.8 m respectively (Lindberg and Grimmond, 2011), the sampling heights span z/z0 = 0.30–2.14.

The 8-point vertical profile system (Fig. 1d) was installed at KSSW and KSNW to allow air to be sampled at a total of 16 locations (ten for a vertical profile, five for two horizontal profiles, one of which was designated an inter-comparison point (IC, Fig. 1c) with one air intake from each valve array). CO2 concentration is measured at each location at 2 Hz for 75 s by an infrared LI840 gas analyser (LI-COR Biosciences, USA), with each 8-point profile completing a cycle in 10 min (for further instrumentation details see Supplementary material S.4). The first 10 s of data for each run were removed to prevent contamination by the previous sample and the remaining concentrations averaged to give a ‘1 min value’. Data reported here were not gap filled. Reasons for data gaps include physical component failure (e.g., burnt out air pumps), theft or failure of on-site computer, electricity supply failure, extensive building works, instrument removal for calibration and software problems. Other measurements used in this paper include shortwave radiation measured at KSSW and KSNW by net radiometers (CNR8 and CNR1 respectively, Campbell Scientific, USA), wind speed and direction measured at KSSW by a 2D sonic anemometer (WX5750, Vaisala, Finland) and by 3D sonic anemometers (KSSW and KSNW: CSAT3, Campbell Scientific, USA; KSB: R3-50, Gill, UK). Calculated quantities used to classify CO2 storage values include z'/L, a measure of atmospheric stability, where z' is the height of the CSAT3 above ground less the zero plane displacement height, and L is the Obukhov length. The ratio z'/L was calculated at 30 min resolution using data from KSSW CSAT3 (Kotthaus and Grimmond, 2014a).

2.3. Climate

The conditions under which this study was undertaken include: prevailing winds from the southwest (for 2012 to 2014, mean wind speed 3.98 m s⁻¹, WXT520, KSSW, height A, Fig. 1d); and an air temperatures range of −2.1 to 31.3 °C (mean annual air temperatures of 11.2 and 12.7 °C in 2013 and 2014 respectively). For more details see Supplementary material S.5.

The observed atmospheric stability (see Section 2.2) was classified as: extremely unstable: −1 < z'/L ≤ −0.1, unstable: −0.1 < z'/L ≤ −0.05, unstable/near neutral: −0.05 < z'/L ≤ −0.01, neutral: −0.01 < z'/L ≤ 0.01, or stable: 0.01 < z'/L ≤ 1000. Kotthaus and Grimmond (2014a) showed that stable atmospheric stratification is rare for this central urban area. The infrequent stable conditions (3.2% of all periods June 2012 to Dec 2014) occurred predominantly at night (20:00 to 06:00 GMT) in autumn and winter, when vertical fluxes are low (Ward et al., 2015) and the relative importance of the storage term is accordingly expected to be enhanced. Although, nocturnal release of storage heat from the urban fabric (Grimmond and Oke, 1999) is usually sufficient to maintain fully turbulent conditions overnight, observations at this site found periods of stability throughout the year and at all hours of the day. For more details of the stability characteristics of the site, see Supplementary material S.5.

3. Diurnal, hebdomadal and seasonal cycles in CO2 storage

The forcings affecting ΔCSS and ΔCSP vary at daily, weekly and seasonal timescales. The general properties of ΔCSP are illustrated for a typical weekday and rest day (e.g., weekend, holiday) in June, 2013 (Fig. 2). These were dry, sunny days, with minimal CO2 emissions above roof height as there was low building occupancy (university summer vacation) and no need for building heating.

3.1. Example days

The 10 min ΔCSP data for 2013/153 to 154 tended (80% of values) to vary between −0.52 and 0.56 μmol m⁻² s⁻¹ (Fig. 2a) within an overall range of ±1.6 μmol m⁻² s⁻¹. Despite the difference in environment, this is comparable to data reported by Iwata et al. (2005), who calculated mean daily storage of 0.32 μmol m⁻² s⁻¹, with skew of −0.32, using a 30 min profile of 6 heights in the Brazilian rainforest. Larger values (−20 to 5 μmol m⁻² s⁻¹) were observed by Araujo et al. (2010), who made 5 s resolution measurements at 6 heights, cycling through the profile in 15 min. They noted a difference in the diurnal cycle of storage depending on whether the profile was situated on a plateau, slope, or valley bottom, with the latter having the most clearly defined cycle of accumulation overnight and dispersal during daylight hours. Similar cycles have been reported from a number of forested sites in Europe (Aubinet et al., 2005), albeit with a smaller range (−5 to 3 μmol m⁻² s⁻¹). Unlike the aforementioned rural studies, no clear diurnal or hebdomadal cycle in the instantaneous ΔCSP values was observed for these particular days (Fig. 2a) at the Strand. Calculated CO2 storage values were exceedingly noisy, with increases in ΔCt tending to be followed by decreases. There are some small visual and skew differences in the frequency-size distributions (Fig. 2b) between weekend (dashed blue line) and weekday (solid red line)
values (skew of 0.42 and −0.24 for Sunday 2013/153 and Monday 2013/154 respectively).

Changes in frequency (Fig. 2d) impact the cumulative CO2 storage (Fig. 2c). Weekend $\Delta C_{SP}$ accumulated nocturnally is lost during the afternoon. The weekend cycle is more muted than the weekday one. The latter has a larger peak but shorter nocturnal accumulation. From mid-rush “hour” there is less. The differences between the weekday and weekend $\Delta C_{SP}$ may relate to traffic volumes or traffic timing. However, comparison of the change in [CO2] with time at each profile height, $z_i$, and traffic volume (Automated traffic count, ATC, 250 m to the north east of KSSW) for 2012/151–2013/150 found no relation using a linear model fitted by least-squares between the two at any profile height ($R^2 < 0.01$, where $R^2$ is calculated as the residual sum of squares divided by the total sum of the squares and subtracted from one). The absolute value of the change in [CO2] with time appeared to increase with increasing traffic volume, with the strongest relation found at the lowest measurement heights. However, no $R^2$ for a linear regression of the change in [CO2] with time on traffic volume exceeded 0.05.

3.2. Seasonality

This section compares $\Delta C_{SP}$ calculated via the two approaches described in Section 1.1 as they are commonly implemented; in other words, $\Delta C_{SP}$ from [CO2] measurements made by a fast response infra red gas analyser at one point high above the height of the local roughness elements (e.g., Nemitz et al., 2002; Crawford and Christen, 2014), and $\Delta C_{SP}$ from [CO2] measurements made by a closed path infra red gas analyser, with air samples drawn from multiple locations in a vertical profile (e.g., Aubinet et al., 2005). The diurnal cycle (columns, Fig. 3) of $\Delta C_{SP}$ is not constant throughout the year (rows, Fig. 3) and varies with measurement method (Fig. 3a/b vs. c/d). The $\Delta C_{SP}$ values calculated from 2 Hz LI840 data averaged to 1 min values between 2012/160 and 2013/155 were averaged by hour of day and month of year. The majority (80%) of these hourly/monthly medians were within ±0.13 mmol m$^{-2}$ s$^{-1}$ (Fig. 3); minor in comparison to median turbulent vertical CO2 flux for the Strand of ~35 mmol m$^{-2}$ s$^{-1}$ (Kotthaus and Grimmond, 2012). For median $\Delta C_{SP}$ values calculated in the same fashion from 10 Hz LI7500 data over the same period, 80% of values were within 0.01–0.11 mmol m$^{-2}$ s$^{-1}$, three orders of magnitude larger and significantly less equally distributed about 0. The difference between $\Delta C_{SP}$ (LI840) and $\Delta C_{SP}$ (LI7500) is also present for the hourly mean values, with 80% of $\Delta C_{SP}$ values within ±0.16 mmol m$^{-2}$ s$^{-1}$ when calculated from LI840 data, and between −0.19 and 0.00 mmol m$^{-2}$ s$^{-1}$ for $\Delta C_{SP}$ LI7500 data for the same period. Not only does the $\Delta C_{SP}$ calculated from the LI7500 data have a greater range of values, it also has a greater (negative) skew than $\Delta C_{SP}$ calculated from the LI840 data. The impact of sampling interval and sensor response time are explored further in Section 6.

3.2.1. LI840 $\Delta C_{SP}$

Both hourly/monthly mean and median $\Delta C_{SP}$ calculated from LI840 data show a clear diurnal cycle, the magnitude and timing of which is seasonally variant (rows, Fig. 3a, b). Peak $\Delta C_{SP}$ is observed in the morning, followed by a midday trough and less clearly defined evening peak. Overnight storage values tend to be negative during the colder months (October to April) and weakly positive otherwise (May to September). The timing of the morning peak is closest to midday during January (~08:00 GMT) and furthest in June (~12:00 GMT), with a steady transition between the two that suggests that the dominant forcing is boundary layer stability and mixing, rather than anthropogenic behaviour. Comparison of hourly median CO2 storage for weekdays and weekends for June—August 2012 found no difference in the timing of the peak or trough CO2 storage values despite higher overnight and lower daytime traffic volume on weekends.

Negative overnight values during winter indicate loss of CO2 from the canyon airspace and coincide with minima in the observed diurnal cycle of $F_{CO2}$ (Kotthaus and Grimmond, 2012, Fig. 14a). This is unlikely to be due to photosynthetic uptake. An alternative explanation is drainage flow away from the canyon below $z_u$, potentially towards the River Thames; however, the EC measurements are predominantly unstable at KSSW even during overnight winter periods (Supplementary material S.5). Within canyon temperatures in winter (2012/335 to 2013/059) exceeded...
those measured at z0 by ~0.8 °C at all times of day on both weekdays and weekends and it is suggested that the losses are either due to sustained thermally induced turbulence as described by Velasco and Roth (2010), or the release of intermittent thermally induced ‘bursts’ of high CO2 air from the canyon (Supplementary material S.6) as described by Salmond et al. (2005) and Lietzke and Vogt (2013). The temporal resolution of the ΔC\textsubscript{CSS} measurements is not sufficient to distinguish the two and the question is revisited in the following section.

### 3.2.2. LI7500 ΔC\textsubscript{CSS}

The ΔC\textsubscript{CSS} calculated from LI7500 data shows a much weaker diurnal cycle relative to seasonality and a much greater difference between mean and median hourly values (rows Fig. 3c, d). Hourly means tended to be negative whilst hourly medians tended to be positive, with the difference between the two greatest for November to January. In other words, data during this period had a greater range of values and was more negatively skewed. The larger positive values could be due to greater emissions from space heating, whereas the intermittent but strongly negative values lend credence to the hypothesis of CO2 venting from the street canyon.

Previous work (Lietzke and Vogt, 2013) indicates that even for street canyons with height:width ratios as low as 0.34—0.70, strong wind shear across the top of a canyon can act as a ‘lid’, preventing turbulent mixing out of street level emissions. Intermittent air parcels break through this lid (Caton et al., 2003) and vent high CO2 air from the canyon into the airspace above. Periods with high shear (and high friction velocity) might therefore be expected to show negatively skewed CO2 storage within the canyon relative to periods with lower shear. The friction velocity calculated from data measured at KSSW was used to classify the change in [CO2] with time at three locations within and above the Strand canyon (Fig. 4). The change in CO2 concentration with time (\(\Delta[CO_2]/\Delta t\)) was chosen in preference to CO2 storage as the absolute value of the latter depends upon the vertical separation of the measurement points. Data collected at point J (Fig. 1d), the lowest measurement location, showed the least change in either magnitude or variance of \(\Delta[CO_2]/\Delta t\) with increasing friction velocity (Fig. 4, Fig. S.7 in Supplementary material S.7). \(\Delta[CO_2]/\Delta t\) measured above the canyon and at canyon half-height (A and G respectively, Fig. 1d) decreased from approximately 1.5 to 0.0 μmol m\(^{-2}\) s\(^{-1}\) with increased friction velocity (0.1—0.5 m s\(^{-1}\); S.7), however there was no consistent shift in skew or sign. Results indicate that increased friction velocity affects the mixing down to about half the canyon height whereas the lowest levels are effectively sheltered.

### 4. Built form and wind direction

CO2 storage depends not only on emissions but also on the efficacy of transport processes. In Section 3.2.2 the lower measurement levels were shown to be effectively sheltered from shear-based turbulence, even at high above-canyon wind speeds. In this section the effect of above-canyon wind speed and direction on within-canyon CO2 storage and CO2 transport is examined in more detail.

The asymmetric Strand canyon runs east-northeast to west-southwest (Fig. 5a), bifurcated opposite the western half of the Strand building by a church (St Mary le Strand). Airflow through the canyon within 20 m of the ground may, depending on the time of year, be deflected or attenuated by the trees lining the street and, depending on the time of day, be blocked or shifted by the large number of buses on the southern half of the roadway (i.e., the side of the measurements). Both sides of the street vary in height and the canyon is not of uniform width; due to the presence of the

![](image-reference)

Fig. 3. Diurnal (y-axis) and seasonal (x-axis) cycle of CO2 storage (key: far right) at KS, London, UK. (a) hourly/monthly mean CO2 storage calculated from profile measurements (ΔC\textsubscript{CSS}), (b) hourly/monthly median ΔC\textsubscript{CSS}, (c) hourly/monthly mean CO2 storage calculated from measurements at a single height (ΔC\textsubscript{CSP}), (d) hourly/monthly median ΔC\textsubscript{CSP}. ΔC\textsubscript{CSS} calculated from [CO2] measured at all heights (Fig. 1d) with a cycle time of 10 min ΔC\textsubscript{CSS} calculated from [CO2] measured at height A (Fig. 1d) by LI7500 at 10 Hz. All data measured 2012/159 to 2013/154 and aggregated by month of year (x-axis) and hour of day (y-axis). The ΔC\textsubscript{CSS} data are in units of μmol m\(^{-2}\) s\(^{-1}\) and the ΔC\textsubscript{CSP} data in mmol m\(^{-2}\) s\(^{-1}\).

![](image-reference)

Fig. 4. Change in CO2 concentration with time (\(\Delta[CO_2]/\Delta t\)) by above-canyon friction velocity for three different heights (A, G and J, see Fig. 1d) within the Strand canyon (London, UK, mean building height, z\textsubscript{0} ~ 21 m) calculated from 10 min [CO2] profiles collected 2012/159 to 2013/154. Bar: median; box: interquartile range; whiskers: 5th to 95th percentile.
church and the curvature of the road it varies from 13 to 46 m. Comparisons between wind direction and CO2 storage in this section are therefore limited to this case study and are unlikely to be widely applicable.

As wind speed and direction measurements at KSB (Fig. 5a) were not available throughout 2012/159 to 2013/154 (the focus of the previous section) a comparison (2013/100 to 2014/100) of data collected by the WXT520s at KSSW and KSNW for both time periods was undertaken. This found that neither the overall wind direction distribution nor the within canyon response differ significantly from year to year.

Wind speed and direction measurements taken at KSSW and KSB are assumed to be representative of the local and canyon air flow respectively, whilst measurements at KSNW represent either the intersection to the east or the canyon to the west. The prevailing wind was from the south-southwest (~210°, Fig. 5b). Although this direction had the highest mean above canyon wind speed, horizontal wind speeds measured at KSNW were at their lowest due to sheltering by the Strand building (Fig. 5d) and were scattered rather than channelled. Horizontal wind speed minima (KSB: 150°, KSNW: 240°) are measured at each site when in the lee of the Strand building and are associated with increasing [CO2] with time (Δ[CO2]/Δt) for heights G-J (Fig. 6).

Channelling of airflow does not occur for wind from the north east quadrant at either KSB or KSNW, potentially due to the more open canyon structure around the eastern intersection (Fig. 5a). This sector is associated with a horizontal wind speed minimum and a peak in vertical wind speed. The most negative (downward) vertical wind speeds were measured when above canyon wind at 290° was channelled to 240° (Fig. 5d), perhaps indicating the presence of a spiral vortex. Both the minimum (290°, KSNW) and the maxima (120°, KSB; 210°, KSNW) in vertical wind speed are associated with a reduction in Δ[CO2]/Δt at all z1 ≤ 16 m (heights G to J, Fig. 1d), suggesting that rather than just recirculating and trapping emissions, vortices act to reduce CO2 with time at both windward and leeward walls. For discussion of the change in variability of Δ[CO2]/Δt with height and above-canyon wind direction, see Supplementary material S.8.

5. Sample locations

One of the first decisions when preparing to make [CO2] profile measurements is the number and position of sample locations. We discuss in this section how these decisions were made. The aim, in general, is to accurately characterise the shape of the vertical [CO2] profile with as few sampling points as possible to minimise data loss due to switching, and installation and running costs. Ideally measurements would be more densely spaced where the concentration changes rapidly with height, and more sparse where [CO2] is more constant with height. It is commonly accepted that for rural profiles, measurement density should be greater closer to the ground as the major source of CO2 is sub-surface respiration and the atmospheric profile of the gas changes more quickly at lower elevations (Wofsy et al., 1993; Goulden et al., 1996; Jarvis et al., 1997; Molder et al., 2000; Iwata et al., 2005; Vogt et al., 2006; Araujo et al., 2010). Six-hourly, median vertical [CO2] profiles measured at the Strand during June and November/December 2012.
(Fig. 7a) suggests this may not be the case for high-density urban environments. The difference between median \([CO_2]\) observed at the top of the tower \((z_t = 46.4\, \text{m above Strand level})\) and the concentrations measured along the profile, with measurement height relative to the mean building height \((z_b)\), shows weekday concentrations in the canyon relative to those measured at \(z_t = 46.4\, \text{m}\) tend to be higher. This can be explained by the greater volume of traffic emissions as the difference is most pronounced in the daytime periods (06:00–12:00 GMT and 12:00–18:00 GMT) during weekdays – the time when traffic volumes are largest. The morning period reveals clearly enhanced concentrations during the summer period compared to the winter period in the lowest layers \((z/z_b < 1)\) for both weekends and weekdays, while in the afternoon, this seasonal shift is only evident for the weekdays. In winter within-canyon \([CO_2]\) for the weekend 12:00–18:00 GMT period exceed those of the weekday, however the difference is within the standard deviation of the data (Fig. 7b).

In Fig. 7a, rather than the smooth, exponentially decreasing concentration curve with height observed in Basel (10 locations spanning 0.17 to 2.3 times the mean building height \(z_b\), Lietzke and Vogt, 2013), median \([CO_2]\) shows a sharp transition at the Strand building roof level (Fig. 7a and 30.5 m, dot-dash horizontal red line), and potentially another close to the mean building height (Fig. 7a and 21 m, dotted horizontal blue line) during the day (06:00–18:00 GMT). The difference between observed profiles in Basel and London is likely to be due to the greater street canyon depth at KS; the height:width ratios for the street canyons in Basel and London are 0.34–0.70 and 0.74–1.28, respectively. Lietzke and Vogt (2013) characterised the flow regime in Basel as ‘wake interference’ (Oke, 1987), leading to greater mixing out of canyon emissions compared to the ‘skimming flow’ conditions likely to dominate at KS. The latter would result in a fast moving ‘lid’ of air over the roof of the Strand building during the prevailing southwestly winds and an active vortex within the canyon, leading to two fairly internally homogenous flow regimes, with some exchange at the interface at roof height. This is observed to some extent in Fig. 7a; however, the difference between above and within canyon \([CO_2]\) is more apparent in the standard deviation of the measured \([CO_2]\) for each time period and height (Fig. 7b). Measurements at or below the Strand building height have consistently higher variance than those above for both summer and winter, weekdays and weekends. This suggests that the \([CO_2]\) profile in an urban street canyon with near 1:1 height:width ratio could be adequately measured by two or three points, provided they were placed appropriately.

This ‘small sample’ hypothesis is tested by extending the analysis of Yang et al. (2007) to the urban \([CO_2]\) profile. The method consists of defining a \(CO_2\) storage time series calculated from a particular configuration of sample inlets as ‘best practice’ (\(\Delta C_{SP}\)) and assessing other configurations based on how closely the calculated \(\Delta C_{S}\) from each configuration agree with \(\Delta C_{SP}\). In this study, \(\Delta C_{SP}\) was defined as the \(CO_2\) storage calculated using data from all ten vertical profile locations. All configurations included \(z_b\), co-located with the eddy covariance system, as such a tower-top measurement is commonly present at micrometeorological sites observing turbulent fluxes. Configurations were grouped by number of sample locations included in the \(CO_2\) storage calculation, (e.g., all configurations with three sample locations fall in ‘Group 3’). Storage was calculated as the change in \([CO_2]\) weighted by the vertical span for each height and divided by the cycle length. The vertical span \((\Delta z_i)\) for a sample location \(z_i\) was defined as \((z_i + z_i)/2 + (z_i - z_i)/2\), except for the lowest height, where the span extended to ground level, and the uppermost height, where the span stopped at \(z_t = z_b = 46.4\, \text{m}\).

\(CO_2\) storage was found to be highly variable with configuration. The \(\Delta C_{S}\) time series for each configuration were compared with \(\Delta C_{SP}\). The configurations within each group with the lowest (Fig. 8a) and highest (Fig. 8b) root mean squared error (RMSE) were considered as the ‘best’ and ‘worst’ configurations respectively. Configurations were found to perform best if the inlet heights were evenly distributed and the vertical span for each available sample point was maximised. In other words if few inlet points are present, the configurations with the greatest distance between the inlets showed the best performance (Fig. 8a). This is in contrast to previous rural work (e.g., Iwata et al., 2005) which has tended to cluster sample points at the base of the profile. Configurations with
Bad performance (highest RMSE, Fig. 8b) have sample points clustered towards the top of the profile, adding little new information with each new location. The ‘value’ of each additional sample point can be assessed by comparing the coefficients of the regression of the DCS calculated from each group’s best configuration onto DCSBP (Fig. 9). Increasing the number of sample points from 1 to 2 increases the coefficient of determination ($R^2$) from 0.15 to 0.80 provided that the second sample point is placed appropriately. In this study this was found to be at just over half the height of the street canyon (height F: 19.5 m; canyon wall height: 34.5 m = Strand building height of 30.5 m + safety wall height of 4.0 m). This supports the ‘small sample’ hypothesis, namely that for street canyons with height:width ratios of approximately 1:1, the urban environment can be adequately measured, and perhaps, modelled, as a two-layer system of above and within-canyon air masses with a transition zone at roof level. It is not necessary to monitor the concentration all the way down to street level as the air within the canyon is well mixed and responds quickly to ground-level changes in $[CO_2]$. As expected, air above roof level is less responsive to street canyon processes and measurements made at 2–3 times mean building height should not be considered directly representative of processes within the street canyon below. The assumption required for the single height calculation of CO$_2$ storage, also known as the storage flux correction, is not valid — at least one other measurement point is required.

6. Sensor response and sampling interval

In this section, the measurement time resolution necessary to capture the majority of the CO$_2$ signal is assessed using the wavelet power spectrum. The effect of under sampling or inadequate sensor
response time is simulated by subsampling and progressively smoothing a $[CO_2]$ time series. The loss of calculated CO2 storage with decreasing time resolution is predictable and is shown to be correctable to an extent.

Continuous, high frequency (10 Hz) $[CO_2]$ measurements were made by LI7500 at heights A and F (Fig. 1d) from 2013/073 onwards. Days 2013/347—2013/365 were chosen for analysis as the campus was closed and rooftop emissions were likely to be minimal. Winter days were chosen despite the sensitivity of the open path sensor to inclement weather due to the greater probability of stable periods (Fig. S.2 in Supplementary material S.5) during which CO2 storage is expected to be significant relative to the vertical flux.

Data were divided into 30 min periods for wavelet analysis. Wavelet analysis has a number of advantages for atmospheric data relative to traditional Fourier analysis when constructing power spectra. Fourier analysis requires regular, complete data sets as any errors or missing values are delocalised throughout the entire spectrum (Farge, 1992). Fourier analysis also has difficulty resolving sudden transitions or aperiodic signals (Farge, 1992), both of which can occur frequently in an urban CO2 storage time series due to the prevalence of intermittent, strong CO2 sources. In contrast, wavelet analysis does not delocalise errors, can resolve sudden step changes or spikes (Salmond et al., 2005), and was found to be considerably quicker to compute than the gap-tolerant (Press and Rybicki, 1988) Lomb-Scargle Periodogram.

Wavelet analysis was conducted on the 10 Hz time series data following Torrance and Compo (1998), and used in-house software, written in R. If a data set of 30 min had fewer than 10 min of continuous data, the data set was rejected for analysis. Data were normalised by subtracting the mean and dividing by the standard deviation of the data set prior to processing. The Mexican Hat wavelet (negative of the second derivative of a Gaussian, appropriately normalized) was applied at 53 temporal scales ranging from 0.2 s to 27.3 min. The Mexican Hat wavelet was chosen as it has previously been used to successfully identify coherent structures in atmospheric turbulence over an urban area (Feigenwinter and Vogt, 2005). Wavelet powers were calculated as the square of the absolute value of the wavelet coefficients, then summed over time and weighted by the original series’ variance and number of data points to give the global wavelet spectrum. Global wavelet spectra were binned by frequency and atmospheric stability for each half hourly period ($z/L$, height A, KSSW, Section 2.2).

The wavelet power spectra for both $\Delta C_{SP}$ and $\Delta C_{SS}$ were found to increase as a power-law with frequency (i.e., a linear relationship on a log—log axis), with approximately a 1:1 relation between 0.005 and 0.1 Hz (200–10 s) and approximately 1:2 relation between 0.1 and 0.5 Hz (10–2 s) (Fig. 10a,b). When normalised by the natural frequency the peak spectral energy was found to be between 0.001 and 0.004 Hz (1000 and 500 s or 8–16 min), although this range has higher frequency during neutral and stable conditions for $\Delta C_{SP}$ (Fig. 10c, d). There was also a secondary peak at approximately 0.1 Hz or 10 s, above which spectral power declines as an inverse-power law with exponent of approximately $-2/3$ for all stability classes. It is therefore expected that measurements with a time resolution of better than 1 s should not produce significant improvement in measured CO2 storage. This was tested by altering the two main controls on the temporal resolution of measured $[CO_2]$: the sensor response time and the logging rate. A correction factor for cases where the required time resolution could not be
attained is proposed.

The basis of both the sensor response time and sensor logging rate analyses is data from two LI7500 installed at KSSW and KSNW during an intensive observation period between 2014/013 and 2014/043. During this time the sample rate was increased from 10 Hz to 20 Hz to ensure that the normal operating frequency (10 Hz) was resolvable. Different logging rates were simulated by sub-sampling this time series at intervals ranging from 10 Hz to 10 min. To check that the response was not sensor specific, collocated continuous LI840 data (2 Hz) time series were also sub-sampled at intervals up to 10 min. A similar technique was used by Heinesch et al. (2007) to estimate the uncertainty in CO2 concentration measurements; however, to enable analysis over a larger time range, data in this study were not resampled multiple times. Both LI7500 and LI840 data sets were also used to investigate the effect of sensor response time. This was simulated by smoothing the concentration time series with a modified, single-sided cosine function.

Due to the volume of data it was too computationally expensive to plot the instantaneous values. Instantaneous ΔC5 values at each time resolution were therefore summed over 30 min periods to give the total storage change within each period. The stability of each 30 min period was calculated as the effective height (z′) over the Obukhov length (L) (Section 2.2) from data measured at KSSW. There was some indication that CO2 storage values were lowest for neutral stability but this was not consistent across all time resolutions.

Slowing the sensor response time has the largest impact on calculated ΔC between 0.1 and 5 s (Fig. 11a), corresponding to a resolvable frequency of 5 to 0.1 Hz (as the highest resolvable frequency is half the sample rate). This frequency range corresponds to the aforementioned portion of the natural frequency normalised wavelet power spectrum as a function of frequency, best-fit by a power law relation with exponent –2/3. Improving the response time to below 0.1 s does not seem to affect calculated ΔC, suggesting that 0.1 s is sufficient. Below 1 s there is little change with increasing smoothing length.

In contrast, on a log–log plot, there is a linear decrease in median absolute half-hourly ΔC5 (|ΔC5|) with increasing sampling interval length, t1, (Fig. 11b) which can be written as:

$$\log_{10}\left(\frac{|\Delta C_5|}{|\Delta C_5|_{t_1}}\right) = a \log_{10}(t_1) + b$$

where a and b are empirically derived constants with a = −1.24 and b = 1.69 for the LI7500 data, and a = −1.14 and b = 1.28 for the LI840 data (all R² > 0.99). If a time series y1 has been under-sampled at time interval t1 by an instrument with constants a1 and b1, a time series y2 with desired time interval t2 can be measured by an instrument with constants a2 and b2, can be estimated from y1 using:

$$y_2 = y_1 \left(\frac{10^{b_2} t_2^{a_2}}{10^{b_1} t_1^{a_1}}\right)$$

If the instrument is the same for both time resolutions, the factor of 10b may be omitted as the instrument specific constant, b2 and b1, will be the same in both the nominator and denominator of (4). However, if converting between instruments, then the instrument specific constants for each series must be used. This correction is effective at reducing the underestimation of ΔC5 at a sampling interval of 10 min by five orders of magnitude (Figs. 11b and 12). Though simple to apply it is derived from statistical analysis of data only and does not reflect any atmospheric process.
7. Interpolation in time and space

Sections 5 and 6 considered the location and temporal resolution of CO₂ concentration measurements necessary to calculate CO₂ storage. In this section, the non-concurrent, spatially irregular CO₂ concentration time series is interpolated onto a regular time/space grid using three different interpolation methods. The interpolated concentration time series and derived CO₂ storage time series are compared to benchmarks and the necessity of interpolating CO₂ concentration time series prior to the CO₂ storage calculation is evaluated.

7.1. Theory

The exact form of the CO₂ storage calculation requires continuous measurements of the [CO₂] profile in time and space (Aubinet et al., 2005). As switched profile measurements are obtained by sampling from several heights sequentially (Mülder et al., 2000), the resulting data are discontinuous at each height. As calculation of the CO₂ storage requires a complete instantaneous profile, further processing can include interpolation of the ‘missing’ data to provide a continuous concentration time series from which complete profiles can be drawn. Alternatively, the average of the values observed throughout a cycle may be taken as representative of that cycle time period.

Calculation of CO₂ storage also requires integration of the change in [CO₂] over the vertical extent of the profile (Aubinet et al., 2005). This can either be accomplished by weighting the change in concentration at each height by the vertical span over which it is deemed relevant or interpolating to a metre grid and summing over the vertical extent without weighting. The former is mathematically equivalent to taking the average of adjacent heights (e.g., Yang et al., 2007).

7.2. Interpolation in time

The effectiveness of different methods of interpolating [CO₂] in time are evaluated using continuous 2 Hz LI840 [CO₂] measurements (KSSW, z₁ = 46.4 m). Data covering half a seasonal cycle (2013/160–365) were split into 75 s ‘runs’ and processed as described in Section 2.2 to provide a benchmark time series with 8 data points per 10 min. A switched profile time series was simulated by extracting every 8th data point (method illustrated in Fig. 13), which was used as input data for three different interpolation functions.

The first function, termed ‘constant’ interpolation, assumes all concentrations within 5 min either side of an extracted data point are equal to the extracted value (Fig. 13). This is equivalent to undertaking no interpolation in time but constructing [CO₂] profiles from the measurements closest to a particular time point. The other two functions, ‘linear’ and ‘cubic spline’ interpolation were performed using the approx and spline R functions (R Core Team, 2013) respectively (Fig. 13, ‘Linear’, ‘Cubic’). Leading or trailing missing values were not interpolated.

Interpolated [CO₂] time series regressed on to the benchmark had coefficients of determination R² > 0.98 (Table 1). The linear interpolation performed the best in terms of R² and RMSE, but the constant value method had values for the slope and intercept of the regression that more closely approached 1.0 and 0.0 respectively. The cubic interpolation was the least effective but the difference was small.
The high, positive correlation between interpolated and measured time series for [CO₂] does not hold as well for ΔC. None of the R² values exceeded 0.05. The ΔC calculated from the linearly interpolated time series was closest to the measured values; despite low R², the slope and intercept for the regression were 0.95 and 0.00 respectively and it had the lowest RMSE.

The lack of strong agreement between the ΔC time series calculated from interpolated data and the benchmark suggests that wherever possible measurements should be made continuously, rather than using a switched profile. Interpolation can adequately reproduce [CO₂]; however, as measured CO₂ storage values for a switched profile are typically on the order of 1 mmol m⁻² s⁻¹, the actual CO₂ storage signal is likely to be smaller than the interpolation error.

### 7.3. Interpolation in space

Given a ‘true profile’ with continuous measurements in space and time is unavailable, the full ten height profile was taken as a ‘benchmark’. Data from the four heights (A, E, F, J) identified (Section 5) as the ‘best’ configuration for that number of sample points were divided into 10 min profiles time-centred on the z̄ = 46.4 m a.g.l. (height A, KSSW) observation and were not interpolated in time. Each 4 height, 10 min vertical profile was interpolated to the remaining six sample points (B, C, D, G, H, I). The interpolation between observations included: linear, cubic and none (i.e., [CO₂] at the interpolated heights were assumed to be the same as [CO₂] at A, E, F or J).

Linear regression was performed between the interpolated and the measured time series for each interpolated height and interpolation method (Table 2). When summed over all interpolated heights for each interpolation method, the lowest RMSE (51.7, 58.3, 58.6 ppm) was when linear interpolation was used, followed by no interpolation and cubic interpolation, respectively (Table 2). The linear interpolation was poorest for the sample points C, D and G (R² = 0.86–0.87). Both the ‘none’ and cubic interpolation improve with decreasing height; however, this may be due to the choice of sample points. Due to the larger vertical separation between A and B–D than E and B–D, the latter had a much greater impact upon the ‘none’ (concentrations taken to be equal to the closest extracted sample point) and cubic interpolations at B–D than the former; despite concentrations measured at B–D often being more similar to A (Section 5) (Table 2). If B–D are considered to be ‘closer’ to A for the purposes of the ‘no’ interpolation, the RMSE for the profile drops to below that of the linear interpolation. Whilst the linear interpolation still performs best in the canyon, simple span weighting is more effective above. All methods perform acceptably (R² > 0.8) for all heights and could be applied without prior knowledge of the profile shape. If the typical profile shape is known and there are obvious physically induced ‘break points’ in the vertical profile, span weighting to the breakpoints and not to the midpoints between sample locations is the preferred method of spatial interpolation.

### 8. Contribution of ΔC to Net Ecosystem Exchange (NEE)

The observed vertical flux of CO₂ (FCO₂) at KSSW (Fig. 1c) varies both diurnally and seasonally (Kotthaus and Grimming, 2012; Ward et al., 2015). When aggregated by hour of day and month of year (Fig. 14a), mean FCO₂ is always positive. Minimum values (<10 mmol m⁻² s⁻¹) occur between 01:00 and 05:00 GMT in summer (June–August). Though fluxes remain high throughout the day, maximum FCO₂ values are observed just after midday (12:00–15:00 GMT) during all seasons, with those during the winter (December–February) 60–100 mmol m⁻² s⁻¹ higher than those in summer.

Addition of a storage term calculated from a switched vertical profile to the vertical flux (subset June 2012–May 2013, Fig. 14b) does not appear to alter the diurnal or seasonal pattern of emissions. Approximately 67% of hourly/monthly flux values were altered by less than 2.5% by the addition of switched ΔC_{\text{a}}. The greatest impact (flux values altered by >5%) was overnight (01:00–05:00 GMT) between July and November, 2012, when vertical fluxes were at their lowest; however, there is no overall inverse relation between size of the vertical flux and percentage change by ΔC_{\text{a}}.

Addition of a storage term calculated from measurements made at a single height (A, Fig. 1d) to the vertical flux (subset June 2013–June 2014, s) gives unrealistic estimates of the net emissions, particularly between 03:00–09:00 GMT December 2013 to March 2014 when the incidence of stable conditions was greatest (S5). The ΔC_{\text{a}} values were found to be 1 to 10 times the magnitude of FCO₂ and consistently opposite in sign (Figs. 3c and 14a). If accurate, this would suggest that either the urban environment is a substantial CO₂ sink or that the majority of urban CO₂ emissions are transported horizontally below the flux measurement height. Given the low vegetation fraction (Ward et al., 2015) and the lack of a horizontal CO₂ gradient within the canyon of comparable magnitude to the vertical CO₂ gradient (not shown), neither is likely. It is therefore suggested that in urban settings CO₂ storage values derived from measurements made at a single height above

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**Table 1**

Coefficient of determination (R²), Slope, Intercept, Root Mean Squared Error (RMSE) and Degrees of Freedom (DoF) for a linear regression of an interpolation in time of an extracted (Fig. 13) [CO₂] time series on to the benchmark time series (Fig. 13). CO₂ storage time series were calculated from interpolated [CO₂] time series. All values given to 3 s.f.

<table>
<thead>
<tr>
<th>Interpolation Method</th>
<th>R²</th>
<th>Slope</th>
<th>Intercept</th>
<th>RMSE</th>
<th>DoF x10⁸</th>
</tr>
</thead>
<tbody>
<tr>
<td>CO₂ Concentration (mmol m⁻²)</td>
<td>None</td>
<td>0.987</td>
<td>0.093</td>
<td>0.131</td>
<td>0.151</td>
</tr>
<tr>
<td></td>
<td>Linear</td>
<td>0.986</td>
<td>0.099</td>
<td>0.136</td>
<td>0.149</td>
</tr>
<tr>
<td></td>
<td>Cubic</td>
<td>0.987</td>
<td>0.097</td>
<td>0.136</td>
<td>0.151</td>
</tr>
<tr>
<td>CO₂ Storage (mmol m⁻² s⁻¹)</td>
<td>None</td>
<td>0.000</td>
<td>0.000</td>
<td>0.022</td>
<td>7.73</td>
</tr>
<tr>
<td></td>
<td>Linear</td>
<td>0.030</td>
<td>0.945</td>
<td>−0.002</td>
<td>5.34</td>
</tr>
<tr>
<td></td>
<td>Cubic</td>
<td>0.011</td>
<td>0.601</td>
<td>−0.001</td>
<td>5.39</td>
</tr>
</tbody>
</table>
the canopy cannot be used to correct the vertical flux term for CO₂ stored below the flux measurement height. Despite the relative ease of measurement and calculation of a storage correction term from measurements made at a single height, the data presented here and the analysis of Section 5 discourage its use.

9. Conclusions

CO₂ storage calculated from a switched, time averaged vertical profile (ΔCSS) typically takes values that are smaller than 100th of the magnitude of the vertical flux term and therefore may be considered negligible during well mixed conditions. CO₂ storage calculated from 10 Hz measurements taken at the same height as the flux measurements (ΔCSS) have approximately the same magnitude as the flux measurements. Both are commonly used approaches to evaluating stored carbon dioxide (e.g., Aubinet et al. 2005; Nemitz et al. 2002 for ΔCSS and ΔCSS respectively). Neither fulfill the requirements described in Supplementary material S.2 of continuous CO₂ concentration measurements in time and space, the latter over the vertical extent of the volume of interest. Therefore although both are often assumed to evaluate the same variable (ΔCSS), neither can strictly be said to do so. Whilst this paper does not explicitly discuss which of the two (ΔCSS and ΔCSS) is closer to the ‘true’ ΔCSS signal, it does evaluate the impact of undersampling in time (relevant to ΔCSS) and space (relevant to ΔCSS). Suggestions for spatial and temporal sampling density of the CO₂ concentration measurements required to calculate CO₂ stored within an urban street canyon, with an empirical correction if the latter cannot be attained, are provided.

Results from London, UK have been reported for a full seasonal cycle and show CO₂ storage to be diurnally and seasonally variable. Approximately 80% of values calculated from a switched profile were within 0 ± 0.13 μmol m⁻² s⁻¹, of similar magnitude to those reported by studies in tropical and temperate forests. Higher emissions due to traffic and combustion for space heating do not result in larger or more positive CO₂ storage values. The timing of peak ΔCSS is controlled by boundary layer expansion and contraction, and the change in CO₂ concentration with time at each height within the canyon varies more consistently with above canyon wind direction than traffic volume, even at the lowest level. The response to synoptic wind direction suggests the presence of a spiral vortex within the canyon which acts to reduce stored CO₂ on the canyon. The magnitude of this effect was one thousand times larger than ΔCSS for the same period; 80% of values were within 0.01–0.11 mmol m⁻² s⁻¹ (i.e., of similar or greater magnitude to observed vertical CO₂ fluxes).

Air within the canyon is well mixed; the vertical CO₂ concentration profile is better approximated by a set of internally homogenous ‘zones’, rather than a logarithmic decrease with increasing height. For an approximately regular, non-pedestrianised canyon the assumption required for the single-height calculation of CO₂ storage is not valid. However, two heights are sufficient to characterise the profile if placed appropriately with an even spatial distribution. Spatial interpolation of the CO₂ concentration profile was found to be ineffective and unnecessary. Temporally interpolated under-sampled CO₂

<table>
<thead>
<tr>
<th>Interpolation Method</th>
<th>Sample Location</th>
<th>R²</th>
<th>Intercept (ppm)</th>
<th>Slope (ppm)</th>
<th>RMSE (ppm)</th>
<th>DoF x10³</th>
</tr>
</thead>
<tbody>
<tr>
<td>None</td>
<td>B (E)</td>
<td>0.82</td>
<td>64.3</td>
<td>0.84</td>
<td>10.4</td>
<td>41</td>
</tr>
<tr>
<td></td>
<td>C (E)</td>
<td>0.82</td>
<td>50.9</td>
<td>0.87</td>
<td>10.9</td>
<td>41</td>
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<td></td>
<td>D (E)</td>
<td>0.85</td>
<td>50.3</td>
<td>0.87</td>
<td>10.0</td>
<td>41</td>
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<tr>
<td></td>
<td>G (F)</td>
<td>0.86</td>
<td>3.2</td>
<td>1.00</td>
<td>10.7</td>
<td>38</td>
</tr>
<tr>
<td></td>
<td>H (J)</td>
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<td>28.9</td>
<td>0.93</td>
<td>7.9</td>
<td>38</td>
</tr>
<tr>
<td></td>
<td>I (J)</td>
<td>0.90</td>
<td>30.0</td>
<td>0.93</td>
<td>8.4</td>
<td>38</td>
</tr>
<tr>
<td>Linear</td>
<td>B (A)</td>
<td>0.90</td>
<td>24.0</td>
<td>0.95</td>
<td>7.7</td>
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Table 2
As Table 1, for [CO₂] time series measured at four heights and interpolated in space. Sample locations (A, E, F, J, Fig. 1d) in brackets indicate the data source for the ‘none’ interpolation method. The R² values are shaded red (poor fit, lower values) to green (better fit, higher values). The RMSE values are shaded green (lower values) to red (higher values). Both R² and slope given to 2 d.p., RMSE and intercept to 1 d.p. Also given are the degrees of freedom (DoF).
concentration time series showed very low agreement with a benchmark CO₂ storage time series. It is preferable to have fewer, continuous measurements than a spatially dense switched profile. Under-sampling of the CO₂ concentration time series can lead to an underestimation of CO₂ storage by several orders of magnitude, however this is correctable. CO₂ storage underestimation due to slow sensor response time is not as simple to rectify as the decrease is non-linear. To achieve the recommended response time of 0.1 s or better, fast response/high frequency gas analysers would need to be used.

Previously, both ΔCSS and ΔCSP have been assumed to represent (with some degree of uncertainty) the true ΔC signal. The uncertainty in the two measurement methods is in part a function of their underlying assumptions. For example, ΔCSS assumes constant change of CO₂ concentration with time in height such that Δ[CO₂]/Δt is the same at all heights. During periods with strong emissions at ground level, such as rush hour, or periods with weak vertical mixing, the uncertainty in the ΔCSS varies increases. Additional uncertainty arises due to insufficient measurement density in space or time, although in this paper the latter (a notable problem with the profile method) has been found to be partially correctable. It is therefore suggested that where possible flux measurements should be corrected by a CO₂ storage term calculated from a vertical profile. If this is not possible then no storage correction should be applied to the vertical flux.

**Acknowledgements**

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**Appendix A. Supplementary material**

Supplementary material related to this article can be found at http://dx.doi.org/10.1016/j.atmosenv.2015.10.012.

**References**


Microsoft, Simmons, 2011. Strand Campus, King’s College London, UK. S1.511946. -0.116481. (Online). https://www.bing.com/maps/#Y3A9NTIuMjE3NDk5fjAuMTcxNzAvjnx2k0D2jno5e1yjnnVvFtAncT1XQdjStnwMkxt (accessed 14/09/15).


