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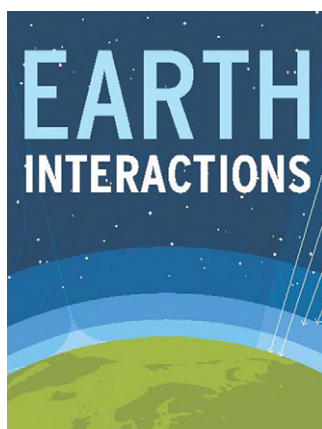
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The Role of CO₂ and Dynamic Vegetation on the Impact of Temperate Land-Use Change in the HadCM3 Coupled Climate Model

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ABSTRACT: Human-induced land-use change (LUC) alters the biogeophysical characteristics of the land surface influencing the surface energy balance. The level of atmospheric CO₂ is expected to increase in the coming century and beyond, modifying temperature and precipitation patterns and altering the distribution and physiology of natural vegetation. It is important to constrain how CO₂-induced climate and vegetation change may influence the regional extent to which LUC alters climate. This sensitivity study uses the HadCM3 coupled climate model under a range of equilibrium forcings to show that the impact of LUC declines under

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increasing atmospheric CO₂, specifically in temperate and boreal regions. A surface energy balance analysis is used to diagnose how these changes occur. In Northern Hemisphere winter this pattern is attributed in part to the decline in winter snow cover and in the summer due to a reduction in latent cooling with higher levels of CO₂. The CO₂-induced change in natural vegetation distribution is also shown to play a significant role. Simulations run at elevated CO₂, yet present-day vegetation show a significantly increased sensitivity to LUC, driven in part by an increase in latent cooling. This study shows that modeling the impact of LUC needs to accurately simulate CO₂-driven changes in precipitation and snowfall and incorporate accurate, dynamic vegetation distribution.

KEYWORDS: Geographic location/entity; Land surface; Physical meteorology and climatology; Albedo; Atmosphere–land interaction; Climate change; Models and modeling; Climate models

1. Introduction

Human-induced land-use change (LUC), such as the conversion of natural land cover to agriculture, transforms the land surface, altering its structure and influencing biogeophysical processes such as albedo, leaf area index (LAI), seasonality, surface roughness, and moisture fluxes. This has implications for the surface energy balance, altering shortwave radiation (SW) and the partitioning of latent and sensible heat (e.g., [Brovkin et al. 2009](#); [Bala et al. 2007](#); [Boisier et al. 2012](#); [Pielke et al. 2002](#)). Understanding the climatic impacts of LUC is crucial for improving climate prediction and potential mitigation strategies.

Observed and modeled data show that LUC influences regional and global climate including temperature, precipitation, and humidity (e.g., [Bonan 1997](#); [Bonan 2001](#); [Gameda et al. 2007](#); [Ge 2010](#); [Fall et al. 2010](#); [Beltrán-Przekurat et al. 2012](#); [Zhao et al. 2001](#); [Claussen et al. 2001](#); [Bathiany et al. 2010](#); [Chase et al. 2000](#); [Kalnay and Cai 2003](#); [Nuñez et al. 2008](#); [Ezber et al. 2007](#); [Zhao and Jackson 2014](#); [Kumar et al. 2013](#); [Pielke et al. 2007](#)). Studies have shown that the regional response depends on the latitude at which it occurs, which may either enhance or attenuate CO₂-induced warming. Tropical LUC in the form of deforestation acts to warm the climate by suppressing evapotranspiration ([Claussen et al. 2001](#); [Snyder et al. 2004](#); [Bathiany et al. 2010](#)), whereas temperate and boreal deforestation is dominated by an increase in albedo and snow persisting longer into the spring, which act to cool the climate ([Claussen et al. 2001](#); [Snyder et al. 2004](#); [Bathiany et al. 2010](#); [Davin et al. 2007](#); [Zhao and Jackson 2014](#); [Lee et al. 2011](#)). The opposing temperate and tropical impact of LUC may drive a small global effect; however, its regional-scale impact is likely to be significantly larger ([Lawrence and Chase 2010](#); [Brovkin et al. 2013](#); [de Noblet-Ducoudre et al. 2012](#)).

Despite the large number of studies, there remain disparities in modeling results as to the magnitude and sometimes direction of regional and global effects ([Pitman et al. 2009](#); [de Noblet-Ducoudre et al. 2012](#); [Brovkin et al. 2013](#)). These inconsistencies have been attributed to differences in how models parameterize albedo, how LUC is implemented, and how crop phenology and evapotranspiration is represented ([Pitman et al. 2009](#)).

With levels of atmospheric CO₂ expected to increase in the coming century and beyond, it is important to constrain how the biogeophysical impact of LUC may change under higher CO₂ forcing. A previous study by [Pitman et al. \(2011\)](#) showed that the biogeophysical impact of LUC depended on the background state of the climate. They

attributed a reduction in the winter impact of temperate LUC at higher concentrations of CO₂ to a reduced snow albedo effect. On the contrary, summertime impacts are shown to increase due to CO₂-induced increase in precipitation and latent cooling.

Higher CO₂ concentrations are expected to not only alter the background climate, including temperature and precipitation patterns, but also the distribution and physiology (i.e., stomatal conductance and LAI) of vegetation. We refer to this as land-cover change (LCC), which differs from LUC in that it is a shift from one natural vegetation type to another (e.g., needleleaf to broadleaf vegetation). Increasing CO₂ is expected to drive a northward migration of the tree line and a shift from one dominant vegetation type to another (e.g., O'ishi et al. 2009; Scholze et al. 2006). As with LUC, LCC has implications for the initial biogeophysical characteristics of the land surface including the albedo parameters, roughness length, canopy height, LAI, and rooting depth. This in turn has implications for evapotranspiration and soil moisture availability that can influence climate (e.g., Davies-Barnard et al. 2015; Niyogi and Xue 2006). Stomatal (or canopy) conductance has been shown in laboratory (Field et al. 1995; Brodribb et al. 2009), field (Hungate et al. 2002; Ainsworth and Rogers 2007), and modeling studies (Medlyn et al. 2001; Gopalakrishnan et al. 2011; Collatz et al. 1991; Boucher et al. 2009) to decrease with higher CO₂ concentrations. This effect has been labeled “CO₂ physiological forcing” (Betts et al. 2007a) and acts to decrease evapotranspiration, atmospheric water vapor, and latent cooling and increase surface temperatures (Boucher et al. 2009; Niyogi et al. 2002; Cao et al. 2010).

Many of the current generation of coupled climate models include a land surface component or dynamic global vegetation model (DGVM) that dynamically simulates LCC, including vegetation composition and physiology depending on climate (e.g., Arora 2002). In multimodel studies assessing the biogeophysical impact of LUC, such as Land-Use and Climate, Identification of Robust Impacts (LUCID), only three of the seven modeling groups used dynamic vegetation, and they all used a different initial land-cover distribution (Pitman et al. 2009; Brovkin et al. 2013). This was due to the difficulty in integration and calibration of a common land-cover map. Different vegetation distribution and physiology would be expected to influence the initial conditions of the land surface; however, their relative role in amplifying or attenuating the biogeophysical impact of LUC under higher CO₂ concentrations has not been investigated.

This study will investigate the biogeophysical impact of LUC in the Hadley Centre Coupled Model, version 3 (HadCM3), a coupled climate model with dynamic vegetation [Top-down Representation of Interactive Foliage and Flora Including Dynamics (TRIFFID)] and a land surface scheme [Met Office Surface Exchange Scheme (MOSES); see section 2]. There are two main aims: 1) to examine the impact of LUC under increasing CO₂ concentrations and 2) to assess the importance of dynamic vegetation and LCC simulated by the model and how this influences the extent to which LUC impacts climate. Section 2 gives a description of the HadCM3 climate model and the simulations used in this study. The results are outlined in section 3 followed by an energy balance analysis in section 4. A discussion and summary is presented in section 5.

2. Methods

HadCM3 is a coupled Earth system model comprising a 3D dynamical atmosphere and ocean components and includes a thermodynamic/free-drift sea ice model (Gordon

et al. 2000). The resolution of the atmospheric component is $3.75^\circ \times 2.5^\circ$ with 19 vertical levels, and the ocean resolution is $1.25^\circ \times 1.25^\circ$ with 20 vertical levels.

The model comprises an interactive dynamic vegetation model (TRIFFID; Cox 2001). TRIFFID simulates as a percentage of each grid box with five plant functional types (PFT)—C3/C4 grasses, shrubs, and broadleaf/needleleaf trees—and four nonvegetation types—bare soil, inland water, urban, and ice. Parameters such as albedo and LAI vary for the five PFTs (Table 1), which in turn influence land surface properties such as the surface energy balance. The PFTs vary according to a number of factors including atmospheric CO_2 concentration, temperature, soil carbon availability, and moisture. Soil carbon is calculated by litterfall and microbial respiration, the latter of which is controlled by temperature and soil moisture. Soil carbon and the configuration of PFTs in each grid box are updated every 10 days based on competition from other plants (i.e., trees top the hierarchy before shrubs and grasses) and the fluxes of carbon, which are calculated by the land surface exchange scheme MOSES, version 2.1 (Essery et al. 2003). The updated vegetation distribution is then reentered into MOSES to update land surface parameters such as albedo and surface roughness. This process maintains a consistent hydrological state between the vegetation and atmosphere (Cox 2001). The radiation scheme used in HadCM3 is that of Edwards and Slingo (1996).

MOSES models the physiological processes of transpiration, respiration, and photosynthesis. This links vegetation to CO_2 concentration and atmospheric conditions and impacts the partitioning of surface water. Higher concentrations of CO_2 act to generally decrease canopy conductance that in turn reduces evapotranspiration rate (Boucher et al. 2009). Conductance is also influenced by temperature, soil moisture, and humidity, with the overall effect scaled by LAI. Overall canopy conductance decreases with humidity and increases with soil moisture (Cox et al. 1999), photosynthetically active radiation (PAR) (Sellers et al. 1992), and LAI.

To simulate LUC and conversion to crop and pastureland, natural vegetation was replaced with C3 and C4 grasses (e.g., wheat/rice and sugarcane/maize, respectively) or bare soil using a mask originally derived from Betts et al. (2007b; Figure 1). The mask represents the global disturbed fraction as of the year 1990 constructed with combined pasture values from Goldewijk (2001) and crop fractions from Ramankutty and Foley (1999). The final LUC fraction for the model grid squares is the area designated as per Figure 1, minus that of the three remaining

Table 1. Albedo and LAI values for PFTs simulated by TRIFFID. Albedo values are split into three parameters; sfree is the maximum value for calculating snow-free albedo, and smin and smax are the respective minimum and maximum snow albedos used to calculate the cold deep snow albedo (see Essery et al. 2001). PFT values are based on the International Geosphere–Biosphere Programme (IGBP) classes (Jones 2004).

PFT	sfree	Albedo smin	smax	LAI
Broadleaf	0.10	0.15	0.30	9.0
Needleleaf	0.10	0.15	0.30	6.0
C3 grass	0.20	0.60	0.80	3.0
C4 grass	0.20	0.60	0.80	4.0
Shrub	0.20	0.40	0.80	3.0

nonvegetation types simulated by TRIFFID—that is, inland water, urban, and ice. The model does not include a harvesting or irrigation scheme.

The model was spun up for 1000 years and run at equilibrium for a further 300 years at four CO₂ concentrations: 350, 700, 1050, and 1400 ppm. For each CO₂ concentration, the model was run with and without LUC to isolate its effect. The long spinup period has resulted in the atmosphere reaching an equilibrium state after approximately 1100 years as shown by stabilized global-mean surface air temperatures (SATs; not shown). The LUC mask ([Figure 1](#)) is fixed throughout the run, whereas remaining natural vegetation can change. These simulations will hereon be labeled 1×, 2×, 3×, and 4×, respectively, and those with LUC will have the suffix LU. Analysis is performed on the final 50 years of each run, and all anomalies shown are 99% significant as calculated by a Student's *t* test.

To understand the impact of vegetation distribution, a further set of simulations was run at 4× CO₂ but with 1× vegetation distribution ([Figure 2a](#)), this is labeled 4×VEG1×. The model was run for 300 years with and without the LUC mask ([Figure 1](#)): the former has the suffix LU. The simulations were initialized from the 4× and 4×LU simulations; however, the background vegetation was replaced and initialized with that simulated by the 1× experiment, as shown in [Figure 2a](#). As a result of the initialization, the atmospheric component of the model reaches a relative state of equilibrium after 150 years, as indicated by stabilized global-mean surface air temperatures (not shown). The 4×VEG1× setup represents a simulation with dynamic vegetation switched off and no climatic influence on the distribution of LCC. The 4× CO₂ concentration was used with the 1× vegetation distribution as these represent the extreme end members of our experimental setup. A summary of experiments is shown in [Table 2](#).

3. Results

3.1. Impact of LUC at increasing CO₂ concentrations

Temperature and precipitation anomalies due to the biogeophysical effects of LUC for Northern Hemisphere summer [June–August (JJA)] are shown in [Figure 3](#).

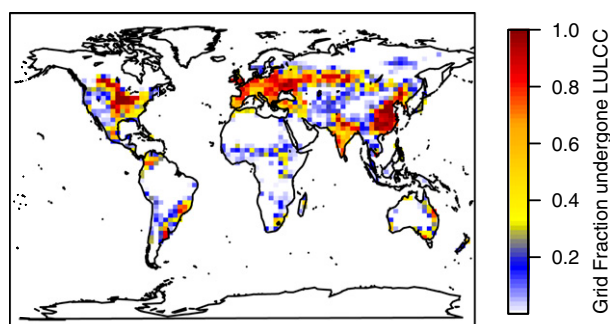


Figure 1. Map showing the proportion of each grid square in HadCM3 that has undergone LUC according to the land-cover mask of [Betts et al. \(2007b\)](#). Final disturbed fraction is calculated with this mask minus the non-vegetation types simulated by TRIFFID.

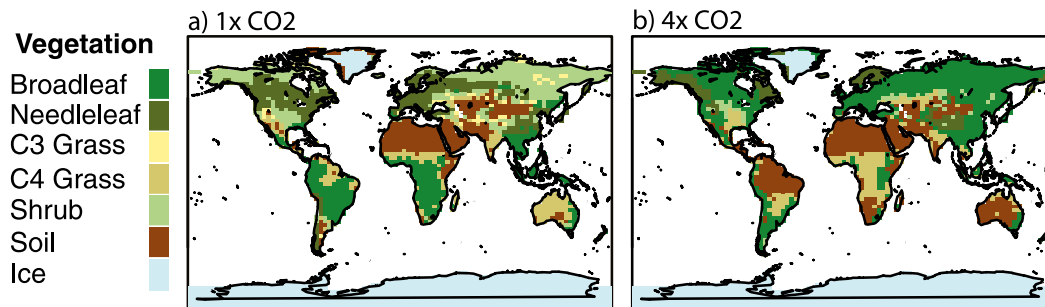


Figure 2. Distribution of natural vegetation simulated by the model occupying greatest proportion of grid square at (a) $1\times\text{CO}_2$ and (b) $4\times\text{CO}_2$.

Mean annual, JJA, and NH winter [December–February (DJF)] anomalies are shown for Europe, North America (NA), and globally for the different model simulations in Figure 4. Table 3 gives the global-mean annual, JJA, and DJF temperature and precipitation values and the anomaly due to LUC in brackets.

LUC acts to cool global SATs for all CO_2 concentrations. The impact at $1\times$ and $2\times$ is comparable but declines globally and regionally for the $3\times$ and $4\times$ simulations. Annual global cooling is in the region of -0.31°C for $1\times$, decreasing to -0.26°C for $4\times$ (Figures 3, 4). The crop mask has a significantly greater regional and seasonal impact. Cooling is heavily focused in the Northern Hemisphere (NH) midlatitudes peaking at around 50°N (Figure 5), correlating with areas of highest LUC fraction, that is, Europe, NA, and China. Cooling across this midlatitude band peaks in JJA due to higher levels of insolation that drive a greater anomaly in the net absorbed shortwave energy at the surface. European and NA JJA cooling is significantly greater than the global average, decreasing by -1.43° and -1.39°C for $1\times$ and -0.79° and -0.80°C for $4\times$, respectively.

The overall decrease in SATs due to LUC for all CO_2 concentrations is primarily a response to an increase in surface albedo and reduced SW energy at the surface. Replacement of natural vegetation with C3 and C4 grasses increases the maximum canopy/snow albedo parameters in the model (Cox 2001) and reduces LAI. The net surface shortwave energy anomaly peaks in the summer months during periods of highest insolation, whereas the surface albedo anomaly peaks in the winter months due to a significant increase in snow cover. The way in which CO_2 alters the surface energy balance is discussed in section 4.

Table 2. Summary of model simulations. Each simulation was run for 300 years with and without LUC (suffix: LU). Analysis is on the final 50 years of each run; all anomalies are 99% significant as calculated by a Student's t test.

Experiment	CO_2 (ppm)	Setup	
$1\times$	350	Equilibrium model run	Each experiment run with and without LUC. Those with LUC have suffix LU.
$2\times$	700	Equilibrium model run	
$3\times$	1050	Equilibrium model run	
$4\times$	1400	Equilibrium model run	
$4\times\text{VEG}$	1400	Vegetation fixed at $1\times\text{CO}_2$ distribution	

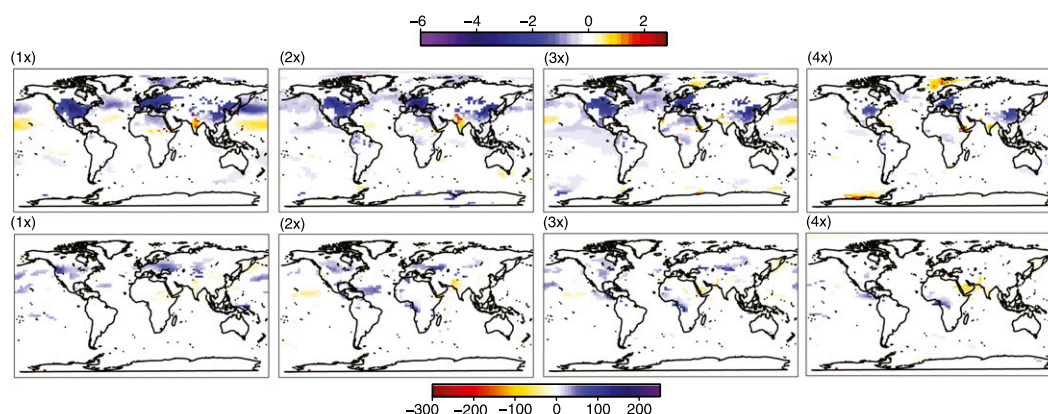


Figure 3. Mean (top) SAT (°C) and (bottom) precipitation (% change) anomalies due to LUC during Northern Hemisphere summer (JJA). Anomalies are 99% confident according to a Student's *t* test.

It is worth noting the contrasting effect seen in India (Figure 3) that experiences a JJA warming peaking at 2.3°C. This pattern was identified in Singarayer et al. (2009) and discussed in more detail by Singarayer and Davies-Barnard (2012), who attributed this to a negative feedback on the monsoon hydrologic cycle. The LUC-driven cooling causes a reduction in convective cloud cover, reducing atmospheric albedo and increasing incoming shortwave (InSW) energy at the surface by up to 3.6 W m⁻² at 1× CO₂×. This acts to counteract the initial cooling impact

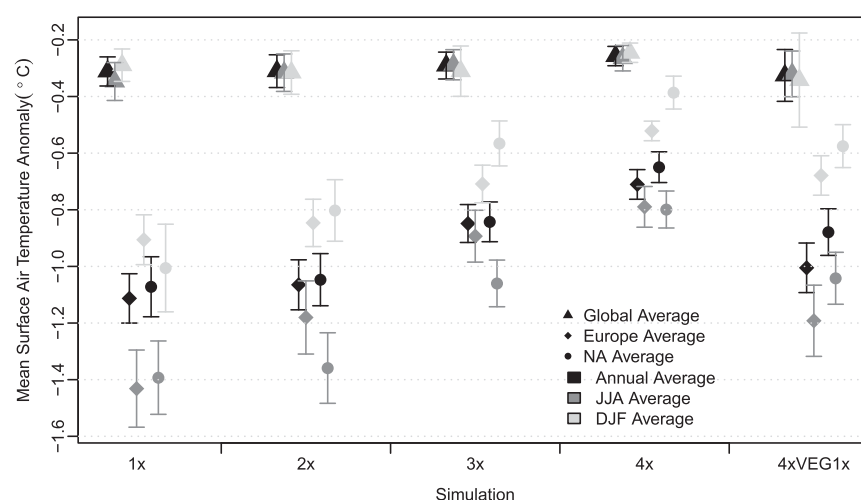


Figure 4. Mean global, European, and NA SAT anomalies due to LUC for the different model simulations (see Table 2 and text for details). Annual, NH, summer (JJA), and NH winter (DJF) are shown. The European region is defined as 35°–60°N and 52.5°E–11.25°W. The North American region is defined as 30°–55°N and 67.5°–123.75°W. Error bars represent the standard error.

Table 3. Mean annual/JJA/DJF SAT and precipitation for experiments. Anomalies due to LUC are shown in brackets.

Experiment	SAT (°C)			Precipitation (mm day ⁻¹)		
	Annual	JJA	DJF	Annual	JJA	DJF
1×	13.96 (−0.31)	15.63 (−0.27)	12.20 (−0.22)	2.90 (−0.0079)	2.93 (−0.0025)	2.85 (−0.0131)
2×	17.91 (−0.31)	19.69 (−0.25)	16.13 (−0.26)	3.04 (−0.0022)	3.07 (−0.0030)	3.00 (−0.0098)
3×	20.16 (−0.29)	22.00 (−0.23)	18.48 (−0.27)	3.10 (−0.0048)	3.12 (−0.0015)	3.08 (−0.0109)
4×	21.96 (−0.26)	23.85 (−0.21)	20.35 (−0.19)	3.16 (−0.0053)	3.16 (−0.0021)	3.15 (−0.0112)
4×	21.20 (−0.33)	23.06 (−0.32)	19.64 (−0.34)	3.08 (−0.0046)	3.09 (−0.0012)	3.08 (−0.0045)
VEG1×						

of LUC and also reduces summer precipitation by an average of $-0.42 \text{ mm day}^{-1}$ (26%). The potential for LUC to reduce Indian monsoon precipitation has also been shown in the observational study of [Niyogi et al. \(2010\)](#).

LUC influences the surface hydrology and the main flows and partitioning of water. Within HadCM3, precipitation is initially split into canopy interception and throughfall at the surface. Removal of forest and the consequent reduction in LAI decreases canopy interception resulting in a negative annual canopy evaporation anomaly in areas of significant LUC. The increase in throughfall drives a positive anomaly in surface runoff and soil evaporation. Soil moisture anomalies are negative in the top layers, shifting to positive with depth. This reflects the shift from the deep rooting depth associated with forest cover to shallow depths of grasses and therefore greater water extraction in the top layers.

LUC has a small impact on global precipitation, with levels decreasing by $-0.0079 \text{ mm day}^{-1}$ for 1×

The negligible change in precipitation due to LUC at 4× CO₂ is likely in part due to the reduced cooling impact of LUC. However, the influence of CO₂ on canopy conductance is also expected to play a role. Conductance decreases by -0.0016 ms^{-1} globally with a greater impact over the tropics. This decrease is enhanced in the northern temperate and boreal regions during Northern Hemisphere summer. This is likely to decrease the rate of evapotranspiration and therefore counteract any temperature–humidity-driven increases in precipitation.

3.2. Role of LCC on the impact of LUC

The shift in LCC for 1×

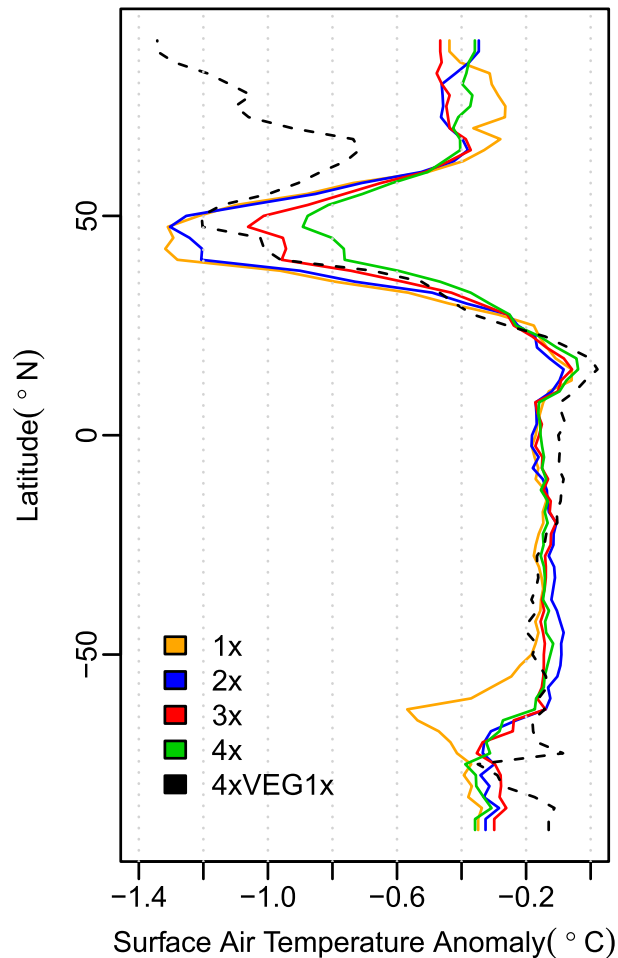


Figure 5. Zonal temperature anomaly during NH summer due to LUC for the five simulations. The dashed black line represents the 4×VEG1× anomaly.

modeling studies (Bala et al. 2006; Lucht et al. 2006; Port et al. 2012; O’ishi et al. 2009; Willeit et al. 2014). There is complete dieback of the Amazon region, a feature apparent in past studies using HadCM3 (Betts et al. 2004; Cox et al. 2000). Other notable impacts include dieback across much of Australia and the decline in the extent of grasses with the exception of sub-Saharan Africa.

The mean temperature and precipitation values for the 4×VEG1× simulation (without LUC) are shown in Table 3. Compared to the 4× simulation, the 4×VEG1× run has lower mean global temperatures on the order of -0.76°C annually. Precipitation levels show a small annual and JJA decline of -0.06 and $-0.08 \text{ mm day}^{-1}$, respectively. The cooler global temperatures for 4×Veg1× are primarily driven by higher surface albedo due to the replacement of predominantly broadleaf vegetation with shrubland and needleleaf. There is enhanced cooling impact in the northern boreal regions that may in part reflect an amplification of the positive sea ice albedo feedback. The reduced SATs show that dynamically

simulated vegetation that responds to CO₂ and climate change causes a decrease in surface albedo, acting as a positive feedback on warming.

In Europe and North America, altering vegetation distribution for 4×VEG1× results in a change from predominantly broadleaf vegetation to needleleaf. In Europe, this results in a cooling of -1.34°C ; however, the albedo parameters for these forest types are the same in HadCM3. Instead, this cooling may be driven by a reduction in LAI, which is also likely to contribute to a reduction in soil evaporation, transpiration, and consequently precipitation. This also acts to increase soil moisture in the deepest soil level (2 m), the rooting depth of trees.

The annual and seasonal anomalies for the 4×VEG1× simulation are shown in Figure 4, and the zonal impact of LUC for all simulations is shown in Figure 5. The cooling impact of LUC is significantly greater for 4xVEG1x relative to 4×, indicating that natural vegetation distribution simulated by the model has a big effect on the impact of LUC. When background vegetation is allowed to dynamically adjust to CO₂ concentration, the cooling impact of LUC is reduced. The global, JJA, and DJF annual cooling is -0.33° , -0.32° , and -0.34°C , respectively. The more significant winter anomaly is likely to be driven by an increase in high northern latitude sea ice cover, which responds to the greater cooling impact of the crop mask. There is a bigger regional impact, with annual European and North American SAT anomalies of -1.00° and -0.88°C , respectively, peaking at -1.19° and -1.04°C in the summer months.

4. Surface energy balance

To understand the mechanism by which the biogeophysical impact of LUC is altered by CO₂ concentration and the background vegetation distribution, the following analysis focuses on changes in the surface energy fluxes. We concentrate specifically on Europe to provide a more focused analysis; the following results are comparable to the North American region that is not shown here.

Figure 6 shows the JJA and DJF surface energy anomalies due to LUC for the 1×, 4×, and 4×VEG1× simulation. The fluxes of outgoing shortwave (OutSW) are increased due to LUC for all simulations due to an increase in surface albedo. LUC also drives an increase in summer latent energy flux due to increased evapotranspiration, although this effect is small at 4× CO₂. This acts to increase the Bowen ratio (i.e., ratio of latent to sensible heat fluxes) and reduce sensible energy at the surface. The conversion of forest to grassland has been shown in studies to actually decrease latent heat flux (Pitman et al. 2009); however, the increase in precipitation due to the albedo-driven cooling and subsequent increase in relative humidity counteracts this (see section 3.1). The 1× and 4×VEG1× anomalies also show a decrease in incoming SW (not shown) that indicates an increase in atmospheric albedo driven by greater cloud cover due to increased relative humidity, reducing SW reaching the surface and enhancing the cooling impact.

We can compare the JJA and DJF anomalies for the different simulations to indicate how CO₂ concentration and changing LCC influence the impact of LUC. There is a greater cooling effect due to LUC at 1× CO₂ compared to 4× CO₂ in both the summer and winter months; however, they appear to be driven by different factors. In summer, LUC significantly increases latent cooling at 1x CO₂ with only

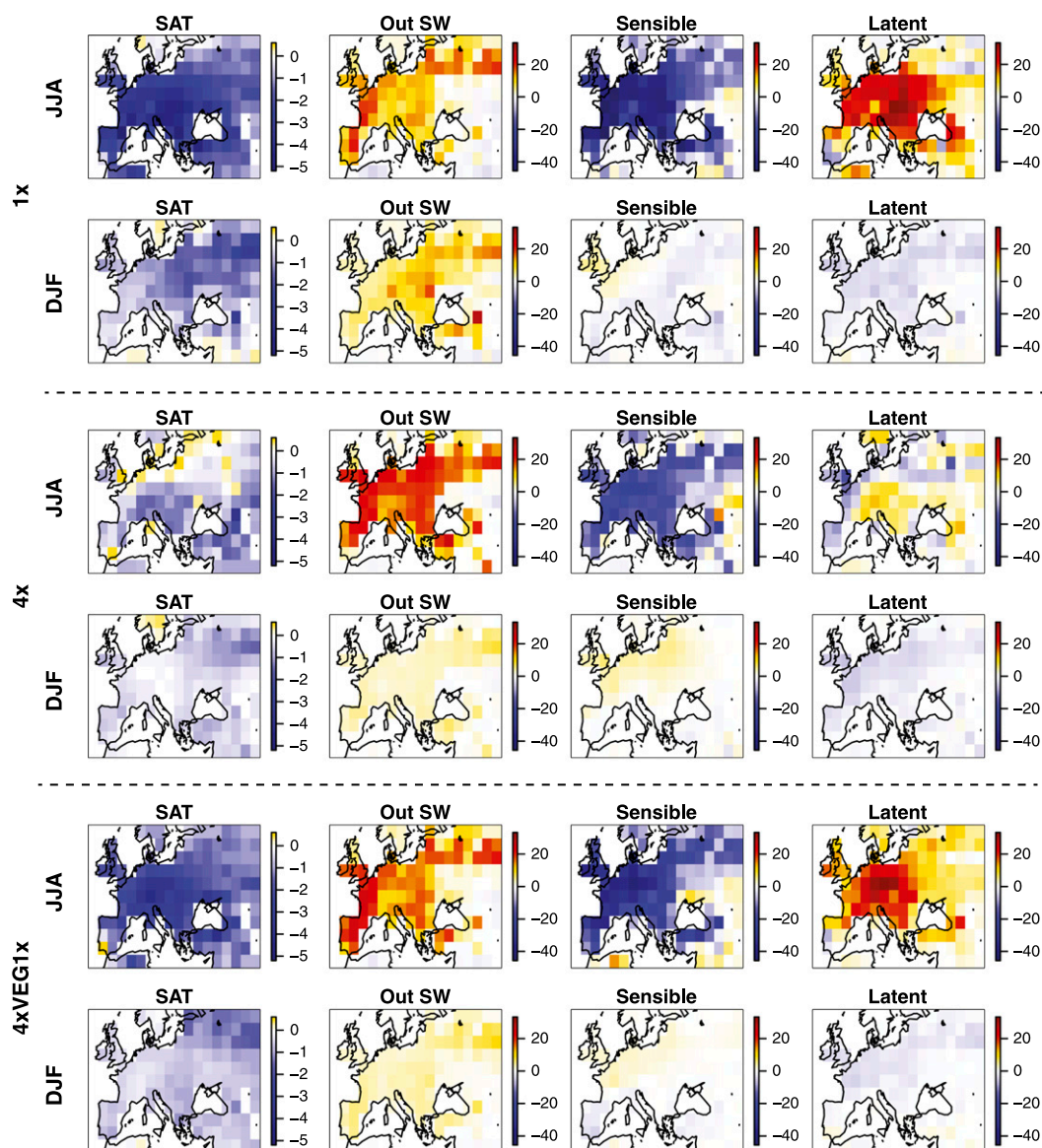


Figure 6. European SAT (°C) and surface energy anomalies (W m⁻²) due to LUC for Northern Hemisphere summer (JJA) and winter (DJF). (top) 1× CO₂, (middle) 4× CO₂, and (bottom) 4×VEG1× simulation (4× CO₂ but 1× vegetation distribution). OutSW represents outgoing shortwave anomaly. The European region is defined as 35°–60°N and 52.5°E–11.25°W.

a small impact at 4× CO₂. In the winter months, it is not a difference in latent cooling but a greater OutSW anomaly (i.e., albedo) at 1× CO₂ compared to 4× that drives enhanced cooling.

The lesser winter albedo anomaly at 4× CO₂ relative to 1× is likely driven by a weakened snow albedo feedback in a warmer world, an effect highlighted in other

climate modeling studies (e.g., Colman and McAvaney 2009). Figure 7 shows the seasonal cycle for European snow cover for the simulations and the anomaly due to LUC. Mean European DJF snow water equivalent decreases from 21.33 to 2.83 kg m⁻² for 1× and 4× CO₂ due to warmer temperatures. The impact of LUC is similarly reduced, with European mean DJF snow cover increasing by 1.99 kg m⁻² for 1×LU and 0.39 kg m⁻² at 4×LU. This contributes to a smaller increase in winter albedo and indicates a weakened winter cooling effect of LUC under 4× CO₂ concentrations.

In contrast to DJF, the JJA temperature discrepancy is driven not by interactions with snow albedo but with a reduction in latent cooling at higher CO₂. This may in part be driven by CO₂-induced reduction in JJA precipitation over Europe, which declines by 43.2% at 4× CO₂ relative to 1×. This decline may in part be driven by the reduction in canopy conductance due to higher CO₂ concentrations. Consequently, LUC occurs in a more moisture-limited environment at higher CO₂ concentrations. This reduces the rate of evapotranspiration, suppresses latent heat flux, and decreases the JJA cooling impact of LUC. The reduction in precipitation with CO₂ is also expected to decrease cloud cover, increasing InSW at the surface and further counteracting cooling.

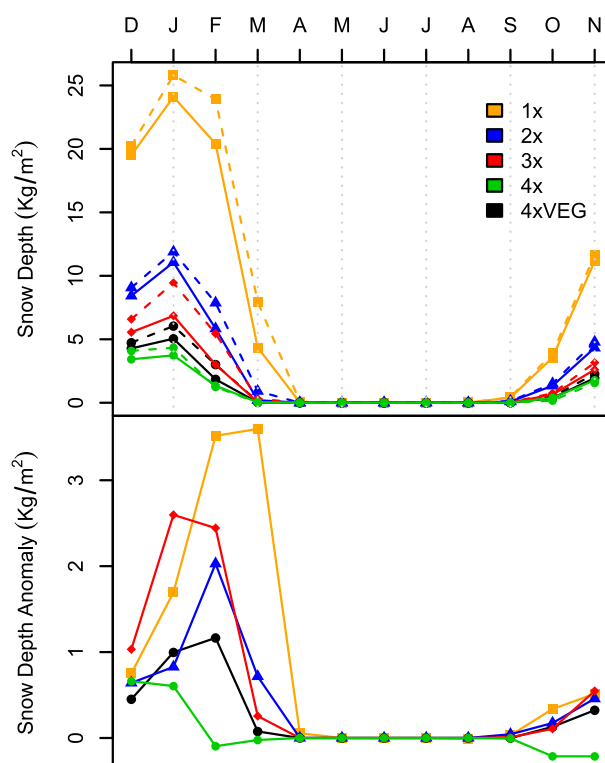


Figure 7. Seasonal cycle of European snow depth (kgm⁻²) for the four CO₂ concentrations and the 4×VEG1× simulation. Dashed lines in the upper chart are with LUC. The European region is defined as 35°–60°N and 52.5°E–11.25°W.

These results show that CO₂-induced changes in climate, that is, snow cover and precipitation, reduce the impact of LUC in the winter and summer months. To understand the influence of background vegetation, we can compare the 4×VEG1× and 4× anomalies. The JJA cooling impact of LUC is enhanced at 4×VEG1×, driven primarily by an increase in latent cooling. This shows that there is a greater anomaly in latent cooling when LUC is modeled with a 1× CO₂ background vegetation distribution instead of 4×, irrespective of the CO₂ concentration. The enhanced latent cooling is not due to an increase in overall precipitation as it is with 1× and 4× (there is a reduction of 4.3% for 4×VEG1×–4×). Instead, this is likely a response to the way in which water is partitioned at the surface by the vegetation. The 1× vegetation distribution is primarily needleleaf forest that has a lower LAI than broadleaf, influencing canopy height and increasing albedo; consequently, there is a reduction in soil evaporation, transpiration, and latent cooling that produces a positive soil moisture anomaly.

These results show that the decreased JJA impact of LUC under higher CO₂ concentrations is driven not only by the CO₂-induced reduction in overall precipitation but also by the CO₂-induced shift in vegetation and the way in which water is partitioned at the surface. This presents a strong case for the need for models to incorporate dynamic vegetation that is sensitive to climate to accurately simulate the impact of LUC.

It is worth noting that there is a negligible difference in the DJF anomalies between 4×VEG1× and 4×. This is likely due to only a small increase in snow depth and enhancement of the snow albedo effect for 4×VEG1× (Figure 7). This shows that in the winter months, LCC has minimal effect, and it is CO₂-driven changes in snow cover that dominate the overall impact of LUC.

5. Discussion and summary

We have used the HadCM3 coupled climate model to show that LUC cools global climate primarily due to an increase in albedo in temperate regions. The overall global SAT anomaly at a CO₂ concentration of 350 ppm is –0.31°C with a significantly enhanced regional effect in areas of high LUC fraction: Europe, North America, and China. This acts to increase regional precipitation due to an increase in relative humidity. The impact on SAT agrees in sign with a number of previous studies although varies in amplitude (e.g., Brovkin et al. 1999; Govindasamy et al. 2001; Matthews et al. 2004; Betts et al. 2007b).

A simplified flowchart outlining the key processes by which LUC cools climate in HadCM3 is shown in Figure 8a. There are two seasonally dependent positive feedback cycles that act to enhance the cooling impact of LUC. In the winter months, LUC-driven cooling further increases snow cover and surface albedo. In the summer months, cooling increases relative humidity and precipitation, which enhances latent cooling and amplifies cloud albedo.

The first part of this study suggests that the cooling impact of LUC will decrease with higher concentrations of CO₂ due to the breakdown of these positive feedback cycles. The mechanism by which this occurs is shown in Figure 8b. Higher concentrations of CO₂ and warmer conditions decrease winter snow cover (Figure 7) and reduce the degree to which it increases with LUC. The overall effect is to weaken the snow albedo feedback and reduce the cooling impact of LUC in the

winter. In the summer months, an increase in CO_2 is simulated to decrease overall precipitation and consequently reduce the evapotranspiration rate. This reduces the partitioning of latent to sensible energy flux and reduces the cooling effect of LUC. The model also simulates a CO_2 -induced decrease in cloud cover, increasing incoming SW energy at the surface and further counteracting cooling. This effect may in part be a response to a reduction in simulated canopy conductance under higher CO_2 concentrations.

These results show that the biogeophysical impact of LUC is dependent on CO_2 concentration and the background climate, an idea initially put forward by [Pitman et al. \(2011\)](#). Their study similarly concluded that the snow albedo mechanism dominates the impact of temperate and boreal LUC in the winter, which is set to diminish with warmer conditions. However, our results presented here disagree with the [Pitman et al. \(2011\)](#) study as to the sign of the summer change under higher CO_2 conditions. They found an increase in summer precipitation under higher CO_2 that drove an increase in latent cooling, the opposite effect to that simulated in HadCM3. Again, this highlights the uncertainty that remains in the modeled response of precipitation to climate change, specifically on a regional scale (see [Schaller et al. 2011](#)). However, the general conclusion that the impact of summer LUC is controlled by latent and sensible heat fluxes that are set to change with increasing CO_2 still holds.

The second part of this study investigates the importance of the background vegetation state in quantifying the impact of LUC. Under higher CO_2 concentrations, natural vegetation is expected to change with a general shift of the tree line and conversion of needleleaf and shrubland to broadleaf vegetation. This has implications for the biogeophysical, biogeochemical, and physiological characteristics of the land surface.

Our results show that within the HadCM3 climate model, the regional summer impact of LUC is strongly affected by LCC and the background vegetation type. Again the simplified mechanism for how this occurs is shown in [Figure 8b](#). The shift in natural vegetation with increasing CO_2 alters albedo, LAI, and canopy conductance. These consequently influence the partitioning of water at the surface, reducing soil moisture and transpiration, decreasing latent to sensible energy flux, and counteracting the cooling impact of LUC. When simulating the impact of LUC at $4\times \text{CO}_2$ but with a $1\times \text{CO}_2$ vegetation distribution, cooling is enhanced specifically in the northern temperate regions during summer. This is a response to an increase in latent cooling due to the replacement of predominantly broadleaf vegetation with needleleaf, a reduction in LAI, a suppressed transpiration rate, and higher soil moisture content, enhancing the positive feedback cycles shown in [Figure 8a](#). This shows that the reduced cooling impact of LUC with higher concentrations of CO_2 is driven not only by how the model simulates precipitation patterns, but also how the model dynamically simulates natural vegetation distribution and the way in which this partitions water at the surface.

There remains a high level of uncertainty associated with the impact of LUC in climate models that needs to be addressed in order to assess regional climate change ([Pielke et al. 2011](#)). We are aware that there are a number of limitations to our results, namely, that this is a single-model equilibrium experiment using a simplified representation of crops and a relatively simple DGVM. This is

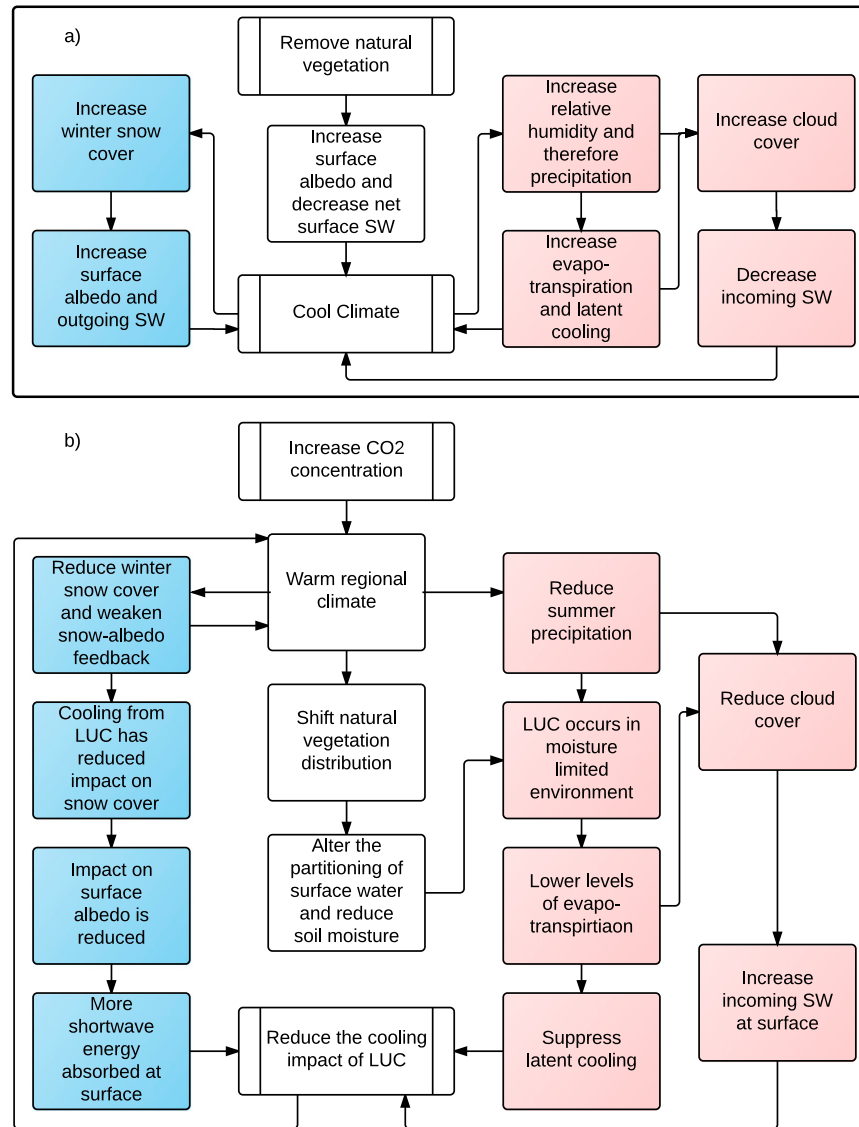


Figure 8. Simplified conceptual flowcharts outlining the key impacts of temperate LUC in HadCM3 and how these are altered by CO₂. (a) The general processes by which LUC causes temperate cooling, and (b) how this cooling impact is reduced under higher concentrations of CO₂ and the role of changing vegetation distribution. The red and blue boxes outline processes that are more prevalent in (although not exclusively) the summer and winter, respectively.

particularly problematic when interpreting the impact and role of precipitation, which remains highly uncertain, yet is crucial to the mechanisms presented in this paper. As a result, the extent and even the sign of changes are likely to be dependent on both the model used and the way in which climate and land-use change have been simulated. As with all modeling studies, a multimodel ensemble would be appropriate in order to validate these mechanisms with a range

of climate models and DGVMs. Also, a perturbed physics ensemble would help determine which specific variable(s) (soil moisture, LAI, albedo, etc.) drive the feedback mechanisms shown in Figure 8b. Building on this, a more comprehensive analysis should also include transient experiments incorporating a range of climate scenarios, such as the representative concentration pathways (RCPs) used in the IPCC report, with different LUC projections. This would not only account for the inconsistencies in models, but also with the uncertainty associated with LUC projections such as the total area and spatial pattern of change. Future projects that may help to address such issues include the Land-Use Model Intercomparison Project (LUMIP), part of the sixth phase of the Coupled Climate Intercomparison Project (CMIP6) with results planned for 2018 to 2019 (see <https://cmip.ucar.edu/lumip>). This will focus on outstanding uncertainty associated with LUC and may help to elucidate the above mechanisms.

Despite the uncertainty associated with specific mechanisms presented, this study again highlights the potential sensitive of LUC to CO₂ concentration and background climate. This is specifically associated with patterns of precipitation and snowfall and the way in which they may change under higher concentrations of CO₂. As a result, in order to accurately assess the impact of LUC in future projections, modelers need to be aware that the impact of LUC may be influenced by the amount and spatial distribution of these variables at increasing CO₂ concentrations. We also show the potentially important role of background vegetation in quantifying the impact of LUC, a variable that has been largely overlooked in previous studies. As discussed, only three of the seven models used in the LUCID experiment (HadGEM2-ES, MIROC-ESM, and MPI-ESM) incorporated a dynamic vegetation scheme (Pitman et al. 2009; Brovkin et al. 2013). These three models still demonstrate a varied response to LUC due to a range of other inconsistencies, including the way in which land surface processes (i.e., albedo and evapotranspiration) are parameterized and the way LUC is implemented (see Pitman et al. 2009). As such, the direct role of including a dynamic vegetation scheme in these models is not certain. The results of our study indicate that it may potentially be important, and it would be pragmatic to investigate its impact more specifically in future LUC modeling studies.

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