

nonlinMIP contribution to CMIP6: model intercomparison project for non-linear mechanisms: physical basis, experimental design and analysis principles (v1.0)

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Good, P., Andrews, T., Chadwick, R., Dufresne, J.-L., Gregory, J. M. ORCID: <https://orcid.org/0000-0003-1296-8644>, Lowe, J. A., Schaller, N. and Shiogama, H. (2016) nonlinMIP contribution to CMIP6: model intercomparison project for non-linear mechanisms: physical basis, experimental design and analysis principles (v1.0). *Geoscientific Model Development*, 9 (11). pp. 4019-4028. ISSN 1991-9603 doi: <https://doi.org/10.5194/gmd-9-4019-2016> Available at <https://centaur.reading.ac.uk/68392/>

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Published version at: <http://dx.doi.org/10.5194/gmd-9-4019-2016>

To link to this article DOI: <http://dx.doi.org/10.5194/gmd-9-4019-2016>

Publisher: European Geosciences Union

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nonlinMIP contribution to CMIP6: model intercomparison project for non-linear mechanisms: physical basis, experimental design and analysis principles (v1.0)

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Received: 8 March 2016 – Published in Geosci. Model Dev. Discuss.: 25 April 2016

Revised: 28 September 2016 – Accepted: 13 October 2016 – Published: 9 November 2016

Abstract. nonlinMIP provides experiments that account for state-dependent regional and global climate responses. The experiments have two main applications: (1) to focus understanding of responses to CO₂ forcing on states relevant to specific policy or scientific questions (e.g. change under low-forcing scenarios, the benefits of mitigation, or from past cold climates to the present day), or (2) to understand the state dependence (non-linearity) of climate change – i.e. why doubling the forcing may not double the response. State dependence (non-linearity) of responses can be large at regional scales, with important implications for understanding mechanisms and for general circulation model (GCM) emulation techniques (e.g. energy balance models and pattern-scaling methods). However, these processes are hard to explore using traditional experiments, which explains why they have had so little attention in previous studies. Some single model studies have established novel analysis principles and some physical mechanisms. There is now a need to explore robustness and uncertainty in such mechanisms across a range of models (point 2 above), and, more broadly, to focus work on understanding the response to CO₂ on climate states relevant to specific policy/science questions (point 1).

nonlinMIP addresses this using a simple, small set of CO₂-forced experiments that are able to separate linear and non-linear mechanisms cleanly, with a good signal-to-noise ra-

tio – while being demonstrably traceable to realistic transient scenarios. The design builds on the CMIP5 (Coupled Model Intercomparison Project Phase 5) and CMIP6 DECK (Diagnostic, Evaluation and Characterization of Klima) protocols, and is centred around a suite of instantaneous atmospheric CO₂ change experiments, with a ramp-up–ramp-down experiment to test traceability to gradual forcing scenarios. In all cases the models are intended to be used with CO₂ concentrations rather than CO₂ emissions as the input. The understanding gained will help interpret the spread in policy-relevant scenario projections.

Here we outline the basic physical principles behind nonlinMIP, and the method of establishing traceability from abruptCO₂ to gradual forcing experiments, before detailing the experimental design, and finally some analysis principles. The test of traceability from abruptCO₂ to transient experiments is recommended as a standard analysis within the CMIP5 and CMIP6 DECK protocols.

1 Introduction

Robust climate impact assessments require, at regional scales, an understanding of physical mechanisms of climate change in general circulation model (GCM) projections. A

further, pragmatic requirement for impact assessments is the ability to emulate (using fast but simplified climate models) GCM behaviour for a much larger range of policy-relevant scenarios than may be evaluated using GCMs directly. These two requirements may be combined into a single question: what is the simplest conceptual framework, for a given well-defined model application, that has quantitative predictive power and captures the key mechanisms behind GCM scenario projections?

Often, one choice has been to assume some form of linearity. In studies of the global energy balance, linearity is often assumed in the form of a constant climate feedback parameter. This parameter may be used to quantify feedbacks in different models (e.g. Zelinka et al., 2013) or, in emulation methods, to parameterize global energy balance models (e.g. Huntingford and Cox, 2000). In understanding or emulating regional patterns of climate change, it is often assumed explicitly that regional climate change is roughly proportional to global-mean warming. In emulation work, this is termed “pattern scaling” (Santer et al., 1990; Mitchell, 2003; Ishizaki et al., 2012; Tebaldi and Arblaster, 2014), but this assumption may also be applied implicitly in understanding mechanisms. Often, physical mechanisms are studied for a single period of a single forcing scenario or in a single high-forcing experiment such as abrupt4xCO₂ (implicitly assuming that the understanding is relevant for other periods or scenarios). The use of pattern scaling is prevalent in studies of climate impacts.

While these approximations appear to work well under many circumstances, significant limitations are increasingly being revealed in such assumptions. The following are of two types: different timescales of response, and non-linear responses. In discussing this, a complication arises in that different linearity assumptions exist. Henceforth, we define “linear” as meaning “consistent with linear systems theory”; i.e. responses that are linear in model forcing (i.e. where doubling the forcing doubles the response). This is different from assuming that regional climate change is proportional to global-mean warming – as in pattern scaling.

Even in a linear system (where responses are linear in forcing), the relationship between two system outputs (e.g. between global-mean temperature and regional sea surface temperature – SST) will in general not be linear. This is due to different timescales of response in different locations and/or variables (Sect. 3.1). Examples include lagged surface ocean warming due to a connection with the deeper ocean (Manabe et al., 1990; Williams et al., 2008; Held et al., 2010; Chadwick et al., 2013a; Andrews and Ringer, 2014) or the direct response of precipitation to forcings (Mitchell et al., 1987; Allen and Ingram, 2002; Andrews et al., 2010; Bala et al., 2010; Bony et al., 2014). One (generally false, but potentially acceptable) assumption of pattern scaling is that regional climate responds over the same timescale as global-mean temperature. Different timescales of response are especially important in understanding and predicting behaviour

under mitigation and geo-engineering scenarios (or over very long timescales).

Non-linear system responses (e.g. Schaller et al., 2013) are more complex to quantify, understand and predict than those of linear systems (Sect. 3.2). Some examples have been known for some time, such as changing feedbacks through retreating snow/sea ice or increasing water vapour (Hansen et al., 2005; Colman and McAvaney, 2009; Jonko et al., 2013; Meraner et al., 2013). Some palaeoclimate evidence supports the idea that climate sensitivity increases with warming (Caballero and Huber, 2013; Shaffer et al., 2016), which is important for the risk of high-end global warming (Bloch-Johnson et al., 2015). The non-linear behaviour of the Atlantic Meridional Overturning Circulation is another example (Hofmann and Rahmstorf, 2009; Ishizaki et al., 2012). More recently, substantial non-linear precipitation responses have been demonstrated in spatial patterns of regional precipitation change in two Hadley Centre climate models with different atmospheric formulations (Good et al., 2012; Chadwick and Good, 2013). This is largely due to simultaneous changes in pairs of known robust pseudo-linear mechanisms (Chadwick and Good, 2013). Regional warming has been shown to be different for a first and second CO₂ doubling, with implications primarily for impact assessment models or studies combining linear energy balance models with pattern scaling (Good et al., 2015). Non-linearity has also been demonstrated in the response under idealized geo-engineering scenarios, of ocean heat uptake, sea-level rise, and regional climate patterns, with different behaviour found when forcings are decreasing than when they are increasing (Bouttes et al., 2013, 2015; Schaller et al., 2014).

Investigation of these mechanisms at regional scales has been constrained by the type of GCM experiment typically analysed. Most previous analyses (e.g. Solomon et al., 2007) have used results from transient-forcing experiments, where forcing changes steadily through the experiment. There are three main problems with this approach. First, information about different timescales of response is masked. This is because the GCM response at any given time in a transient-forcing experiment is a mixture of different timescales of response (Li and Jarvis, 2009; Held et al., 2010; Good et al., 2013), including short-timescale responses (e.g. ocean mixed-layer response from forcing change over the previous few years) through long-timescale behaviour (including deeper ocean responses from forcing changes multiple decades to centuries earlier). Second, in transient-forcing experiments, non-linear behaviour is hard to separate from linear mechanisms. For example, in an experiment where CO₂ is increased by 1 % per year for 140 years (1pctCO₂), we might find different spatial patterns at year 70 (at 2xCO₂) than at year 140 (at 4xCO₂). This could be due to non-linear mechanisms (due to the different forcing level and associated different climate state). However, it could also be due to linear mechanisms: year 140 follows 140 years of forcing increase, and therefore includes responses over longer response

timescales than at year 70 (only 70 years of forcing increase). Third, signal-to-noise ratios of regional climate change can be relatively poor in such experiments.

These three issues may be addressed by the use of idealized abruptCO₂ GCM experiments: an experiment where CO₂ forcing is instantaneously changed, then held constant. The simplified forcing in such experiments simplifies the understanding of physical mechanisms of response. In these abrupt CO₂ experiments, responses over different timescales (fast and slow responses) are separated from each other. Further, responses at different forcing levels may be directly compared, e.g. by comparing the response in abrupt2xCO₂ and abrupt4xCO₂ experiments over the same timescale – both have identical forcing time histories, apart from the larger forcing magnitude in abrupt4xCO₂. Finally, high signal-to-noise is possible: averages may be taken over periods of 100 years or more (after the initial ocean mixed-layer adjustment, change is gradual in such experiments). Recent work (Good et al., 2012, 2013; Zelinka et al., 2013; Bouttes et al., 2015; Good et al., 2015) has established that these experiments contain global- and regional-scale information quantitatively traceable to more policy-relevant transient experiments – and, equivalently, that they form the basis for fast simple climate model projections traceable to the GCMs. In other studies (e.g. Frolicher et al., 2014), pulse experiments have been used to separate different timescales of response (where forcing is abruptly increased, then abruptly returned to the control state). We use abruptCO₂ experiments because they offer greater signal-to-noise in the change signal (important for regional-scale studies) as well as also for consistency with the CMIP6 (Coupled Model Intercomparison Project Phase 6) DECK (Diagnostic, Evaluation and Characterization of Klima) abrupt4xCO₂ experiment.

The CMIP5 abrupt4xCO₂ experiments have thus been used widely: including quantifying GCM forcing and feedback behaviour (Gregory et al., 2004; Zelinka et al., 2013), and for traceable emulation of GCM projections of global-mean temperature and heat uptake (Good et al., 2013; Stott et al., 2013). Abrupt4xCO₂ is also part of the CMIP6 DECK protocol (Meehl et al., 2014).

nonlinMIP builds on the CMIP5 and CMIP6 DECK designs to explore non-linear responses (via additional abruptCO₂ experiments at different forcing levels). It also explores responses over slightly longer timescales – extending the CMIP5 abrupt4xCO₂ experiment by 100 years.

2 Relating abruptCO₂ to gradual forcing scenarios: the step-response model

In using the highly idealized abruptCO₂ experiments, it is essential that their physical relevance (traceability) to more realistic gradual forcing experiments is determined. We cannot a priori reject the possibility that some GCMs could respond unrealistically to the abrupt forcing change. A key tool here

is the step-response model (described below). This (Hasselmann et al., 1993) is a response-function method, which aims to predict the GCM response to any given transient-forcing experiment, using the GCM response to an abruptCO₂ experiment. Such a prediction may be compared with the GCM transient-forcing simulation, as part of a traceability assessment (discussed in detail in Sect. 5).

Once some confidence is established in traceability of the abruptCO₂ experiments to transient-forcing scenarios, the step-response model has other roles: to explore the implications (for different forcing scenarios) of physical understanding gleaned from abruptCO₂ experiments, to help separate linear and non-linear mechanisms (Sect. 5), and potentially as a basis for GCM emulation. The method description below also serves to illustrate the assumptions of linear system theory.

The step-response model represents the evolution of radiative forcing in a scenario experiment by a series of step changes in radiative forcing (with one step taken at the beginning of each year). The method makes two linear assumptions. First, the response to each annual forcing step is estimated by linearly scaling the response in a CO₂ step experiment according to the magnitude of radiative forcing change. Second, the response y_i at year i of a scenario experiment is estimated as a sum of responses to all previous annual forcing changes (see Fig. 1 of Good et al., 2013, for an illustration):

$$y_i = \sum_{j=0}^i w_{i-j} x_j, \quad (1)$$

where x_j is the response of the same variable in year j of the CO₂ step experiment. w_{i-j} scales down the response from the step experiment (x_j) to match the annual change in radiative forcing during year $i-j$ of the scenario (denoted ΔF_{i-j}):

$$w_{i-j} = \frac{\Delta F_{i-j}}{\Delta F_s}, \quad (2)$$

where ΔF_s is the radiative forcing change in the CO₂ step experiment. All quantities are expressed as anomalies with respect to a constant-forcing control experiment.

This approach can in principle be applied at any spatial scale for any variable for which the assumptions are plausible (e.g. Chadwick et al., 2013a).

3 Linear and non-linear mechanisms, and the relevance of abruptCO₂ experiments

Here we discuss further, with examples, the distinction between linear and non-linear mechanisms, when they are important, and the relevance of abruptCO₂ experiments.

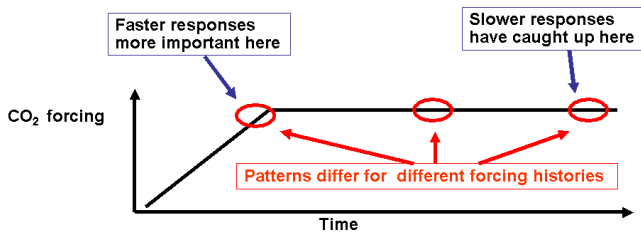


Figure 1. Schematic illustrating a situation where linear mechanisms can cause climate patterns to evolve. This represents a scenario where global-mean radiative forcing (black line) is ramped up, then stabilized. At the time indicated by the left red oval, responses with shorter timescales are relatively important, due to the recent increase in forcing. At the time marked by the right-hand oval, forcing has been stabilized for an extended period, so the responses with longer timescales (such as sea-level rise) have had more time to respond to the initial forcing increase.

3.1 Linear mechanisms: different timescales of response

Even in a linear system, regional climate change per kelvin of global warming will evolve during a scenario simulation. This happens because different parts of the climate system have different timescales of response to forcing change.

This may be due to different effective heat capacities. For example, the ocean mixed layer responds much faster than the deeper ocean, simply due to a thinner column of water (Li and Jarvis, 2009). However, some areas of the ocean surface (e.g. the Southern Ocean and south-east subtropical Pacific) show lagged warming, due to a greater connection (via upwelling or mixing) with the deeper ocean (e.g. Manabe et al., 1990; Williams et al., 2008). The dynamics of the ocean circulation and vegetation may also have their own inherent timescales (e.g. vegetation change may lag global warming by years to hundreds of years; Jones et al., 2009). At the other extreme, some responses to CO₂ forcing are much faster than global warming: such as the direct response of global-mean precipitation to forcings (Mitchell et al., 1987; Allen and Ingram, 2002; Andrews et al., 2010) and the physiological response of vegetation to CO₂ (Field et al., 1995).

In a linear system, patterns of change per kelvin of global warming are sensitive to the forcing history. For example in Fig. 1, a scenario is illustrated where forcing is ramped up, then stabilized. Three periods are highlighted, which may have different patterns of change per kelvin of global warming, due to different forcing histories: at the leftmost point, faster responses will be relatively more important, whereas at the right the slower responses have had some time to catch up. A key example is the different responses of global-mean warming and global-mean sea-level rise under Representative Concentration Pathway 2.6 (RCP2.6), as shown in Figs. SPM.7 and SPM.9 of the IPCC Fifth Assessment Report (IPCC, 2013). Under RCP2.6, global-mean warming

ceases after 2050, when radiative forcing is approximately stabilized (corresponding qualitatively to the period when the black line is horizontal in Fig. 1). In contrast, sea-level rise continues at roughly the same rate throughout the century. Therefore, in RCP2.6, the sea-level rise per kelvin of global warming increases after 2050. This is largely because the timescale of deep ocean heat uptake is much longer than that of ocean mixed-layer warming.

By design, abruptCO₂ experiments separate GCM responses with different timescales (i.e. separating faster responses from slower responses): the response of a given variable in year Y of the experiment corresponds to the response of that variable over the timescale Y . This is used, for example, (Gregory et al., 2004) to estimate radiative forcing and feedback parameters for GCMs: plotting radiative flux anomalies against global-mean warming can separate “fast” and “slow” responses. For example, the top-of-atmosphere outgoing shortwave flux shows a rapid initial change before the global-mean temperature has had time to respond.

3.2 Non-linear responses

Non-linear mechanisms arise for a variety of reasons. Often, however, it is useful to describe them as state-dependent feedbacks. For example, the snow–albedo and sea-ice–albedo feedbacks become small at high or low snow depth (Hall, 2004; Eisenman, 2012). Soil moisture–temperature feedbacks can also be state dependent (Seneviratne et al., 2006, 2010): feedback is small when soil moisture is saturated, or so low that moisture is tightly bound to the soil (in both regimes, evaporation is insensitive to change in soil moisture). Sometimes, non-linear mechanisms may be better viewed as simultaneous changes in pairs of properties. For example, convective precipitation is broadly a product of moisture content and dynamics (Chadwick et al., 2013b; Chadwick and Good, 2013; Bony et al., 2014; Oueslati et al., 2016). Both moisture content and atmospheric dynamics respond to CO₂ forcing, so in general we might expect convective precipitation to have a non-linear response to CO₂ forcing. In addition, the Clausius–Clapeyron equation introduces some non-linearity in the increase of specific humidity with warming. Of course, more complex non-linear responses exist, such as for the Atlantic Meridional Overturning Circulation.

In contrast to linear mechanisms, non-linear mechanisms are sensitive to the magnitude of forcing. For example, the two points highlighted in Fig. 2 may have different patterns of change per kelvin of global warming, due to non-linear mechanisms. In contrast, linear mechanisms would cause no difference in the patterns of change per kelvin of global warming between the two points in Fig. 2, because the two scenarios have the same forcing history apart from a constant scaling factor.

An example is the snow–ice albedo feedback, which tends to change in magnitude with increased global temperature,

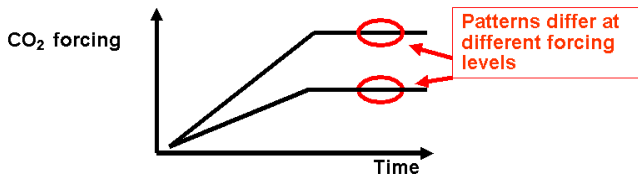


Figure 2. Schematic illustrating the point that non-linear mechanisms can cause climate patterns to differ at different forcing (and hence global temperature) levels. This represents two different scenarios, whose forcing time series is identical apart from a constant scale factor (the higher forcing scenario has about twice the forcing of the lower scenario).

due to declining snow and ice cover, and the remaining snow and ice being in areas of lower solar insolation (Colman and McAvaney, 2009).

AbruptCO₂ experiments may be used to separate non-linear from linear mechanisms. This can be done by comparing the responses at the same timescale in different abruptCO₂ experiments. Figure 3 compares abrupt2xCO₂ and abrupt4xCO₂ experiments over years 50–149. A “doubling difference” is defined (Good et al., 2015), measuring the difference in response to the first and second CO₂ doublings. In most current simple climate models (e.g. Meinshausen et al., 2011), the radiative forcing from each successive CO₂ doubling is assumed identical (because forcing is approximately linear in log[CO₂]; Myhre et al., 1998). With this assumption, a linear system would have zero doubling difference everywhere. Therefore, the doubling difference is used as a measure of non-linearity. The question of which abruptCO₂ experiments to compare, and over which timescale, is discussed in Sect. 5.

In some GCMs, the forcing per CO₂ doubling has been shown to vary with CO₂ (Colman and McAvaney, 2009; Jonko et al., 2013). However, this variation depends on the specific definition of forcing used (Jonko et al., 2013). Currently, this is folded into our definition of non-linearity. If a robust definition of this forcing variation becomes available in the future, it could be used to scale out any difference in forcing between pairs of abruptCO₂ experiments, to calculate an “adjusted doubling difference”.

4 Experimental design

nonlinMIP is composed of a set of abruptCO₂ experiments (the primary tools), plus a CO₂-forced transient experiment. AbruptCO₂ experiments are driven by changes in atmospheric CO₂ concentration: CO₂ is abruptly changed, then held constant. These build on the CMIP5 and CMIP6 DECK protocols (the required runs from these are detailed in Table 1). The additional nonlinMIP runs (Table 2) are assigned three priority levels. The three options for participation are (1) only the “essential” simulation, (2) all “high priority” plus

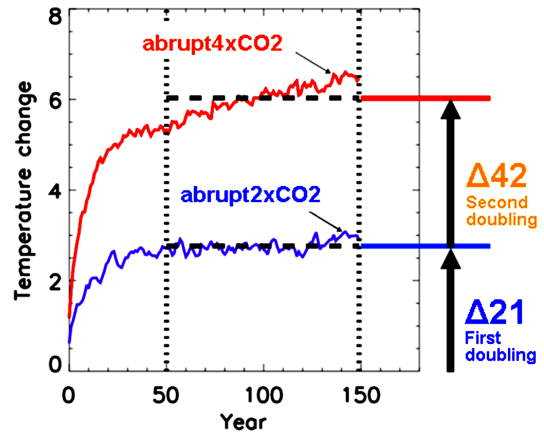


Figure 3. Defining the “doubling difference”. The red and blue lines show illustrative time series of a variable (in this example, global-mean temperature from HadGEM2-ES) from the abrupt4xCO₂ and abrupt2xCO₂ experiments. Doubling difference = $\Delta 42 - \Delta 21$ (the difference in response between the first and second CO₂ doublings). This is defined for a specific timescale after the abrupt CO₂ change – in this example, it is for means over years 50–149.

the “essential” simulations, or, preferably, (3) all simulations. The experiments in Table 1 are required in all cases. All experiments must be initialized from the same year of a pre-industrial control experiment, except for abrupt4xto1x (see Table 2). A typical analysis procedure is outlined in Sect. 5.

The nonlinMIP design is presently limited to CO₂ forcing, although the same principles could be applied to other forcings.

5 Basic analysis principles

This section outlines the applications and general principles behind analysis of nonlinMIP results. First, some general applications are introduced, before giving more detail on how one particular application (quantifying and understanding non-linear change) may be analysed. The addition of the abrupt2xCO₂ experiment to the standard DECK abrupt4xCO₂ permits quantifying and understanding climate change due to CO₂ for three main applications:

1. under global warming approximately comparable to that envisaged by the Paris agreement (quantified by abrupt2xCO₂ – pre-industrial control)
2. climate change approximately comparable to that avoided by mitigation (quantified by abrupt4xCO₂ – abrupt2xCO₂)
3. non-linear change (the difference between 2 and 1).

Applications 1 and 2 are expected to be of the widest interest to the community, as they could be analysed using the same methods as have already been used extensively to study

Table 1. List of CMIP5/CMIP6 DECK experiments required by nonlinMIP.

Experiment	Description	Role
piControl	Pre-industrial control experiment	
Abrupt4xCO2	CO ₂ abruptly quadrupled, then held constant for 150 years.	Separate different timescales of response.
1pctCO2	CO ₂ increased at 1% per year for 140 years (i.e. as CMIP5 1pctCO2 experiment), then decreased by 1% per year for 140 years (i.e. returning to pre-industrial conditions).	To test traceability of the abruptCO ₂ experiments to more realistic transient-forcing conditions. Adding the ramp-down phase explores physics relevant to mitigation and geo-engineering scenarios.

the response in the CMIP5 abrupt4xCO₂ experiment, but for climate states more relevant to the policy questions outlined in (1) and (2). Useful signal-to-noise ratios should be possible because ~ 100 -year means may be analysed (e.g. over years 50–149, where climate is relatively stable as it follows the initial ocean mixed-layer warming). Application 3 is more specialized, and is discussed in more detail below.

The abrupt0.5xCO₂ experiment permits analogous work, extending the relevance to colder past climates, and exploring one aspect of how past change may differ from future change. It also allows non-linear mechanisms to be studied with greater signal-to-noise ratio:

- change under past cold climates (abrupt0.5xCO₂ – piControl)
- non-linear change: like 3, but with larger signal-to-noise ratio ($[\text{abrupt4xco2} - \text{abrupt2xco2}] - [\text{piControl} - \text{abrupt0.5xCO2}]$).

In quantifying non-linear change (applications 3 or 5 above), the primary idea is to find where the step-response model (Sect. 2) breaks: since the step-response model is based on a linear assumption, this amounts to detecting non-linear responses.

The aim is to focus subsequent analysis. If non-linearities in a quantity of interest are found to be small, then analysis may focus on understanding different timescales of response from a single abruptCO₂ experiment: linearity means that the physical response (over a useful range of CO₂ concentrations) is captured by a single abruptCO₂ experiment. This represents a considerable simplification. If, on the other hand, non-linearities are found to be important, the focus shifts to understanding the different responses in different abruptCO₂ experiments. The choice of which abruptCO₂ experiments to focus on, and over which timescales, is discussed below.

5.1 First step: check basic traceability of abrupt4xCO₂ to the transient-forced response near 4xCO₂

The test described here is recommended as a routine analysis of the CMIP6 DECK experiments (even if nonlinMIP experiments are not performed). The aim is to confirm whether the

abruptCO₂ experiments contain realistic physical responses in the variables of interest, as previously done for global-mean temperature and heat uptake for a range of CMIP5 models (Good et al., 2013), for regional-scale warming and ocean heat uptake (Bouttes et al., 2015; Good et al., 2015), and for other global-mean quantities for HadCM3 (Good et al., 2011). This also, rules out the most pathological non-linearities (e.g. if the response to an abrupt CO₂ change in a given GCM was unrealistic). Although this test has been done for a range of models and variables, traceability cannot be assumed to hold for all models and variables.

The linear step-response model should first be used with the abrupt4xCO₂ response, to predict the response near year 140 of the 1pctCO₂ experiment (i.e. near 4xCO₂). This prediction is then compared with the actual GCM 1pctCO₂ result. This should first be done for global-mean temperature: this assessment has previously been performed for a range of CMIP5 models (Good et al., 2013), giving an idea of the level of accuracy expected. If the abruptCO₂ response is fundamentally unrealistic, it is likely to show up in the global temperature change. This approach may then be repeated for spatial patterns of warming, and then for the quantities of interest. Abrupt4xCO₂ is used here as it has larger signal-to-noise than abrupt2xCO₂, yet is representative of forcing levels in a business-as-usual scenario by 2100. However, the tests may also be repeated using abrupt2xCO₂ – but compared with year 70 of the 1pctCO₂ experiment (i.e. at 2xCO₂).

The step-response model emulation under these conditions should perform well for most cases: the state at year 140 of the 1pctCO₂ experiment is very similar to that of abrupt4xCO₂ (same forcing, similar global-mean temperature), so errors from non-linear mechanisms should be minimal. If large errors are found, this may imply caution about the use of abruptCO₂ experiments for these variables, or perhaps point to novel non-linear mechanisms that may be understood by further analysis.

5.2 Second step: characterising non-linear responses

Having established some level of confidence in the abruptCO₂ physical response, the second step is to look for non-linear responses. This first involves repeating the tests

Table 2. nonlinMIP experimental design. Three options are only the “essential” simulation, all “high priority” plus the “essential” simulations, or all simulations. The experiments in Table 1 are required in all cases.

Experiment (priority)	Description	Role
Abrupt2xCO ₂ (essential)	as abrupt4xCO ₂ (see Table 1), but at double pre-industrial CO ₂ concentration.	to diagnose non-linear responses (in combination with abrupt4xCO ₂); assess climate response and (if appropriate) make climate projections with the step-response model at forcing levels more relevant to mid- or low-forcing scenarios.
Abrupt0.5xCO ₂ (essential)	as abrupt4xCO ₂ (see Table 1), but at half pre-industrial CO ₂ concentration	to diagnose non-linear responses (in combination with abrupt4xCO ₂ and abrupt2xCO ₂); offers greater signal-to-noise ratios for regional precipitation change than if just abrupt2xCO ₂ was used; also relevant to palaeoclimate studies.
Extend both abrupt2xCO ₂ and abrupt4xCO ₂ by 100 years (high priority)		permit improved signal-to-noise ratio in diagnosing some regional-scale non-linear responses; explore longer timescale responses than in CMIP5 experiment; permit step-response model scenario simulations from 1850 to 2100; allow traceability tests (via the step-response model) against most of the 1pctCO ₂ ramp-up–ramp-down experiment; provide a baseline control for the abrupt4xto1x experiment.
1pctCO ₂ ramp-down (medium priority)	initialized from the end of 1pctCO ₂ ; CO ₂ is decreased by 1% per year for 140 years (i.e. returning to pre-industrial conditions).	to test traceability of the abruptCO ₂ experiments to more realistic transient-forcing conditions; adding the ramp-down phase explores a much wider range of physical responses, providing a sterner test of traceability; relevant also to mitigation and geo-engineering scenarios, and offers a sterner test of.
Abrupt4xto1x (medium priority)	initialized from year 100 of abrupt4xCO ₂ , CO ₂ is abruptly returned to pre-industrial levels, then held constant for 150 years.	quantify non-linearities over a larger range of CO ₂ (quantifies responses at 1xCO ₂); assess non-linearities that may be associated with the direction of forcing change.
Abrupt8xCO ₂ (medium priority)	as abrupt4xCO ₂ , but at 8× pre-industrial CO ₂ concentration; only 150 years required here.	quantify non-linearities over a larger range of CO ₂ .

from step 1 above, but for different parts of the 1pctCO₂ and 1pctCO₂ ramp-down experiments, and using different abruptCO₂ experiments for the step-response model.

An example is given in Fig. 4 (but for different transient-forcing experiments). This shows results for global-mean precipitation in the HadCM3 GCM (Good et al., 2012), under an idealized simulation where forcing is ramped up at a constant rate for 70 years, then ramped down at the same rate for 70 years. Here, the step-response model prediction using abrupt4xCO₂ (red curves) is only close to the actual GCM simulation (black), where the transient-forced simulation is near to 4xCO₂ (i.e. near year 70). Similarly, the prediction using abrupt2xCO₂ (blue curves) works only near 2xCO₂ (near years 35 or 105). Otherwise, quite large errors are seen, and the predictions with abrupt2xCO₂ and abrupt4xCO₂ are quite different from each other. This implies that there are large non-linearities in the global-mean precipitation response in this GCM, and that they may be

studied by comparing the responses in the abrupt2xCO₂ and abrupt4xCO₂ experiments.

Having identified some non-linear response, and highlighted two or more abruptCO₂ experiments to compare (in the previous example abrupt2xCO₂ and abrupt4xCO₂), the non-linear mechanisms may be studied in detail by comparing the responses in the different abruptCO₂ experiments over the same timescale (e.g. via the doubling difference, as in Fig. 3). This allows for (Good et al., 2012; Chadwick and Good, 2013; Good et al., 2015) non-linear mechanisms to be separated from linear mechanisms (not possible in a transient-forcing experiment). It is expected that analysis will focus on the 100-year period over years 40–139 of the experiments (the relatively stable period after the initial ocean mixed-layer warming).

In the same spirit as other CMIP5 and CMIP6 idealized experiments, nonlinMIP will help understand non-linear mechanisms by isolating the signal of non-linear mechanisms

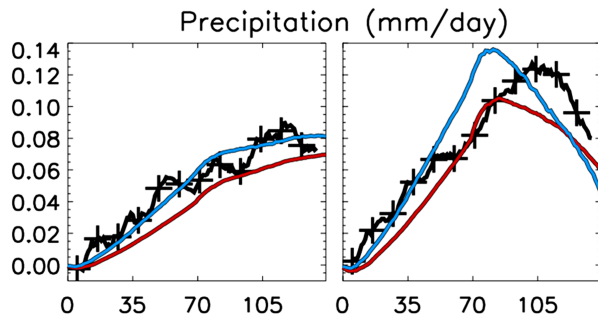


Figure 4. Finding non-linear responses in transient-forcing experiments (figure from Good et al., 2012). Time series of global-mean precipitation change under two experiments. Left: where CO₂ is increased by 1 % per year, then stabilized at 2× pre-industrial levels. Right: where CO₂ is increased by 2 % per year for 70 years, then decreased by 2 % per year for 70 years. Black: GCM. Red: step-response model using the abrupt4xCO₂ response. Blue: the abrupt2xCO₂ response.

more effectively. This occurs in two ways: first, by using simplified forcing compared to the time-dependent, RCP projections (the latter feature multiple forcings of evolving strength). The simplified forcing means that alternative mechanisms (from different forcing agents or linear mechanisms) may be ruled out by design. Secondly, contamination of the signal from internal variability may be reduced, as averages of around 100 years are possible.

The magnitude of internal variability may also be estimated at the different levels of CO₂ forcing. This could be used to help explore changes in variability with warming (Seneviratne et al., 2006; Screen, 2014), and to assess significance of any signal of non-linear change in the time mean climate. Internal variability could be estimated from years 40 to 139 of the experiments (after the initial warming of the ocean mixed layer), after removing a fitted linear trend.

6 Conclusions

These experiments can help improve climate science and consequent policy advice in a number of ways. The focus is on understanding mechanisms (given the idealized nature of the experiments). A further application, however, is that energy balance models could be tuned to the different experiments, to explore the importance, for projections, of state dependence of feedback parameters (Hansen et al., 2005; Colman and McAvaney, 2009; Caballero and Huber, 2013). Also, if certain regions are found to show strongly non-linear behaviour in these experiments, this could help focus assessment of impact tools like pattern scaling or time shifting (e.g. Herger et al., 2015).

Probably of widest interest is the fact that the additional experiments will allow understanding work to focus on climate states more directly relevant to discrete policy/science

questions (the benefits of mitigation; impacts of scenarios consistent with the Paris agreement; or understanding past cold climates; see start of Sect. 5). These questions may show important differences, due to state dependence (non-linearity) of mechanisms, but for many cases the nature of the non-linearity may not need to be assessed. A classical example is the snow–albedo feedback: the strength of this would be different in a warm vs. a cold world (due to different baseline snow cover), but if the focus is on understanding the warm world, the first priority is to study experiments representative of the warm world (with the correct climate state).

There is also a need to quantify and understand, at regional scales, non-linear mechanisms of climate change; that is, do the above science/policy questions give significantly different answers (e.g. different patterns of rainfall change), and why? This is difficult to do using transient model experiments alone, for two reasons: contamination due to different timescales of response, and noise from internal variability.

This paper outlines the basic physical principles behind the nonlinMIP design, and the method of establishing traceability from abruptCO₂ to gradual forcing experiments, before detailing the experimental design and finally some general analysis principles that should apply to most studies based on this dataset.

7 Data availability

Results will be made available as part of the CFMIP project, within the sixth model intercomparison project, CMIP6.

Acknowledgements. This work was supported by the Joint UK DECC/Defra Met Office Hadley Centre Climate Programme (GA01101). Jonathan Gregory received funding from the European Research Council under the European Community’s Seventh Framework Programme (FP7/2007–2013), ERC grant agreement 247220, project “Seachange”. Hideo Shiogama was supported by the SOUSEI program from the Ministry of Education, Culture, Sports, Science and Technology of Japan and the Environment Research and Technology Development Fund (S-10) of the Ministry of the Environment of Japan. Nathalie Schaller was supported by the Swiss National Science Foundation.

Edited by: D. Roche

Reviewed by: two anonymous referees

References

- Allen, M. R. and Ingram, W. J.: Constraints on future changes in climate and the hydrologic cycle, *Nature*, 419, 224–232, doi:10.1038/nature01092, 2002.
- Andrews, T. and Ringer, M. A.: Cloud feedbacks, rapid adjustments, and the forcing–response relationship in a transient CO₂ reversibility scenario, *J. Climate*, 27, 1799–1818, doi:10.1175/Jcli-D-13-00421.1, 2014.

- Andrews, T., Forster, P. M., Boucher, O., Bellouin, N., and Jones, A.: Precipitation, radiative forcing and global temperature change, *Geophys. Res. Lett.*, 37, L14701, doi:10.1029/2010gl043991, 2010.
- Bala, G., Caldeira, K., and Nemani, R.: Fast versus slow response in climate change: Implications for the global hydrological cycle, *Clim. Dynam.*, 35, 423–434, doi:10.1007/s00382-009-0583-y, 2010.
- Bloch-Johnson, J., Pierrehumbert, R. T., and Abbot, D. S.: Feedback temperature dependence determines the risk of high warming, *Geophys. Res. Lett.*, 42, 4973–4980, doi:10.1002/2015GL064240, 2015.
- Bony, S., Bellon, G., Klocke, D., Sherwood, S., Fermepin, S., and Denvil, S.: Robust direct effect of carbon dioxide on tropical circulation and regional precipitation (vol 4, pg 447, 2013), *Nat. Geosci.*, 7, 547–547, doi:10.1038/Ngeo2192, 2014.
- Bouttes, N., Gregory, J. M., and Lowe, J. A.: The reversibility of sea level rise, *J. Climate*, 26, 2502–2513, doi:10.1175/Jcli-D-12-00285.1, 2013.
- Bouttes, N., Good, P., Gregory, J. M., and Lowe, J. A.: Nonlinearity of ocean heat uptake during warming and cooling in the famous climate model, *Geophys. Res. Lett.*, 42, 2409–2416, doi:10.1002/2014GL062807, 2015.
- Caballero, R. and Huber, M.: State-dependent climate sensitivity in past warm climates and its implications for future climate projections, *P. Natl. Acad. Sci. USA*, 110, 14162–14167, doi:10.1073/pnas.1303365110, 2013.
- Chadwick, R. and Good, P.: Understanding non-linear tropical precipitation responses to CO₂ forcing, *Geophys. Res. Lett.*, 40, 4911–4915, doi:10.1002/grl.50932, 2013.
- Chadwick, R., Wu, P. L., Good, P., and Andrews, T.: Asymmetries in tropical rainfall and circulation patterns in idealised CO₂ removal experiments, *Clim. Dynam.*, 40, 295–316, doi:10.1007/s00382-012-1287-2, 2013a.
- Chadwick, R., Boutle, I., and Martin, G.: Spatial patterns of precipitation change in CMIP5: Why the rich do not get richer in the tropics, *J. Climate*, 26, 3803–3822, 2013b.
- Colman, R. and McAvaney, B.: Climate feedbacks under a very broad range of forcing, *Geophys. Res. Lett.*, 36, L01702, doi:10.1029/2008gl036268, 2009.
- Eisenman, I.: Factors controlling the bifurcation structure of sea ice retreat, *J. Geophys. Res.-Atmos.*, 117, D01111, doi:10.1029/2011jd016164, 2012.
- Field, C. B., Jackson, R. B., and Mooney, H. A.: Stomatal responses to increased CO₂ – implications from the plant to the global-scale, *Plant Cell. Environ.*, 18, 1214–1225, doi:10.1111/j.1365-3040.1995.tb00630.x, 1995.
- Frolicher, T. L., Winton, M., and Sarmiento, J. L.: Continued global warming after CO₂ emissions stoppage, *Nat. Clim. Change*, 4, 40–44, doi:10.1038/Nclimate2060, 2014.
- Good, P., Gregory, J. M., and Lowe, J. A.: A step-response simple climate model to reconstruct and interpret aogcm projections, *Geophys. Res. Lett.*, 38, L01703, doi:10.1029/2010gl045208, 2011.
- Good, P., Ingram, W., Lambert, F. H., Lowe, J. A., Gregory, J. M., Webb, M. J., Ringer, M. A., and Wu, P. L.: A step-response approach for predicting and understanding non-linear precipitation changes, *Clim. Dynam.*, 39, 2789–2803, doi:10.1007/s00382-012-1571-1, 2012.
- Good, P., Gregory, J. M., Lowe, J. A., and Andrews, T.: Abrupt CO₂ experiments as tools for predicting and understanding cmip5 representative concentration pathway projections, *Clim. Dynam.*, 40, 1041–1053, doi:10.1007/s00382-012-1410-4, 2013.
- Good, P., Lowe, J. A., Andrews, T., Wiltshire, A., Chadwick, R., Riddle, J. K., Menary, M. B., Bouttes, N., Dufresne, J. L., Gregory, J. M., Schaller, N., and Shiogama, H.: Nonlinear regional warming with increasing CO₂ concentrations, *Nat. Clim. Change*, 5, 138–142, doi:10.1038/Nclimate2498, 2015.
- Gregory, J. M., Ingram, W. J., Palmer, M. A., Jones, G. S., Stott, P. A., Thorpe, R. B., Lowe, J. A., Johns, T. C., and Williams, K. D.: A new method for diagnosing radiative forcing and climate sensitivity, *Geophys. Res. Lett.*, 31, L03205, doi:10.1029/2003gl018747, 2004.
- Hall, A.: The role of surface albedo feedback in climate, *J. Climate*, 17, 1550–1568, doi:10.1175/1520-0442(2004)017<1550:Trosaf>2.0.Co;2, 2004.
- Hansen, J., Sato, M., Ruedy, R., Nazarenko, L., Lacis, A., Schmidt, G. A., Russell, G., Aleinov, I., Bauer, M., Bauer, S., Bell, N., Cairns, B., Canuto, V., Chandler, M., Cheng, Y., Del Genio, A., Faluvegi, G., Fleming, E., Friend, A., Hall, T., Jackman, C., Kelley, M., Kiang, N., Koch, D., Lean, J., Lerner, J., Lo, K., Menon, S., Miller, R., Minnis, P., Novakov, T., Oinas, V., Perlwitz, J., Rind, D., Romanou, A., Shindell, D., Stone, P., Sun, S., Tausnev, N., Thresher, D., Wielicki, B., Wong, T., Yao, M., and Zhang, S.: Efficacy of climate forcings, *J. Geophys. Res.-Atmos.*, 110, D18104, doi:10.1029/2005jd005776, 2005.
- Hasselmann, K., Sausen, R., Maierreimer, E., and Voss, R.: On the cold start problem in transient simulations with coupled atmosphere-ocean models, *Clim. Dynam.*, 9, 53–61, 1993.
- Held, I. M., Winton, M., Takahashi, K., Delworth, T., Zeng, F. R., and Vallis, G. K.: Probing the fast and slow components of global warming by returning abruptly to preindustrial forcing, *J. Climate*, 23, 2418–2427, doi:10.1175/2009jcli3466.1, 2010.
- Herger, N., Sanderson, B. M., and Knutti, R.: Improved pattern scaling approaches for the use in climate impact studies, *Geophys. Res. Lett.*, 42, 3486–3494, doi:10.1002/2015GL063569, 2015.
- Hofmann, M. and Rahmstorf, S.: On the stability of the atlantic meridional overturning circulation, *P. Natl. Sci. Acad. USA*, 106, 20584–20589, doi:10.1073/pnas.0909146106, 2009.
- Huntingford, C. and Cox, P. M.: An analogue model to derive additional climate change scenarios from existing gcm simulations, *Clim. Dynam.*, 16, 575–586, 2000.
- IPCC: Summary for policymakers, in: *Climate change 2013: The physical science basis. Contribution of working group I to the fifth assessment report of the intergovernmental panel on climate change*, edited by: Stocker, T. F., Qin, D., Plattner, G. K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M., Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- Ishizaki, Y., Shiogama, H., Emori, S., Yokohata, T., Nozawa, T., Ogura, T., Abe, M., Yoshimori, M., and Takahashi, K.: Temperature scaling pattern dependence on representative concentration pathway emission scenarios, *Climatic Change*, 112, 535–546, doi:10.1007/s10584-012-0430-8, 2012.
- Jones, C., Lowe, J., Liddicoat, S., and Betts, R.: Committed terrestrial ecosystem changes due to climate change, *Nat. Geosci.*, 2, 484–487, doi:10.1038/Ngeo555, 2009.

- Jonko, A. K., Shell, K. M., Sanderson, B. M., and Danabasoglu, G.: Climate feedbacks in cesm3 under changing CO₂ forcing. Part ii: Variation of climate feedbacks and sensitivity with forcing, *J. Climate*, 26, 2784–2795, doi:10.1175/Jcli-D-12-00479.1, 2013.
- Li, S. and Jarvis, A.: Long run surface temperature dynamics of an a-ogcm: The hadcm3 4xco(2) forcing experiment revisited, *Clim. Dynam.*, 33, 817–825, doi:10.1007/s00382-009-0581-0, 2009.
- Manabe, S., Bryan, K., and Spelman, M. J.: Transient-response of a global ocean atmosphere model to a doubling of atmospheric carbon-dioxide, *J. Phys. Oceanogr.*, 20, 722–749, doi:10.1175/1520-0485(1990)020<0722:Troago>2.0.Co;2, 1990.
- Meehl, G. A., Moss, R., Taylor, K. E., Eyring, V., Stouffer, R. J., Bony, S., and Stevens, B.: Climate model intercomparisons: Preparing for the next phase, *Eos Trans. AGU*, 95, 77–78, 2014.
- Meinshausen, M., Raper, S. C. B., and Wigley, T. M. L.: Emulating coupled atmosphere-ocean and carbon cycle models with a simpler model, MAGICC6 – Part 1: Model description and calibration, *Atmos. Chem. Phys.*, 11, 1417–1456, doi:10.5194/acp-11-1417-2011, 2011.
- Meraner, K., Mauritsen, T., and Voigt, A.: Robust increase in equilibrium climate sensitivity under global warming, *Geophys. Res. Lett.*, 40, 5944–5948, doi:10.1002/2013GL058118, 2013.
- Mitchell, J. F. B., Wilson, C. A., and Cunningham, W. M.: On CO₂ climate sensitivity and model dependence of results, *Q. J. Roy. Meteor. Soc.*, 113, 293–322, 1987.
- Mitchell, T. D.: Pattern scaling – an examination of the accuracy of the technique for describing future climates, *Climatic Change*, 60, 217–242, 2003.
- Myhre, G., Highwood, E. J., Shine, K. P., and Stordal, F.: New estimates of radiative forcing due to well mixed greenhouse gases, *Geophys. Res. Lett.*, 25, 2715–2718, 1998.
- Oueslati, B., Bony, S., Risi, C., and Dufresne, J. L.: Interpreting the inter-model spread in regional precipitation projections in the tropics: role of surface evaporation and cloud radiative effects, *Clim. Dynam.*, 47, 2801–2815, doi:10.1007/s00382-016-2998-6, 2016.
- Santer, B., Wigley, T., Schlesinger, M., and Mitchell, J. F. B.: Developing climate scenarios from equilibrium gcm results, Report No. 47, Max Planck Institute for Meteorology, Hamburg, 1990.
- Schaller, N., Cermak, J., Wild, M., and Knutti, R.: The sensitivity of the modeled energy budget and hydrological cycle to CO₂ and solar forcing, *Earth Syst. Dynam.*, 4, 253–266, doi:10.5194/esd-4-253-2013, 2013.
- Schaller, N., Sedláček, N. J., and Knutti, R.: The asymmetry of the climate system's response to solar forcing changes and its implications for geoengineering scenarios, *J. Geophys. Res.-Atmos.*, 10, 5171–5184, 2014.
- Screen, J. A.: Arctic amplification decreases temperature variance in northern mid- to high-latitudes, *Nat. Clim. Change*, 4, 577–582, doi:10.1038/Nclimate2268, 2014.
- Seneviratne, S. I., Luthi, D., Litschi, M., and Schar, C.: Land-atmosphere coupling and climate change in europe, *Nature*, 443, 205–209, doi:10.1038/Nature05095, 2006.
- Seneviratne, S. I., Corti, T., Davin, E. L., Hirschi, M., Jaeger, E. B., Lehner, I., Orlowsky, B., and Teuling, A. J.: Investigating soil moisture-climate interactions in a changing climate: A review, *Earth-Sci. Rev.*, 99, 125–161, doi:10.1016/j.earscirev.2010.02.004, 2010.
- Shaffer, G., Huber, M., Rondanelli, R., and Pedersen, J. O. P.: Deep time evidence for climate sensitivity increase with warming, *Geophys. Res. Lett.*, 43, 6538–6545, doi:10.1002/2016GL069243, 2016.
- Solomon, S., Qin, D., Manning, M., Chen, Z., Marquis, M., Averyt, K. B., Tignor, M., and Miller, H. L.: Contribution of working group i to the fourth assessment report of the intergovernmental panel on climate change, Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2007.
- Stott, P., Good, P., Jones, G., Gillett, N., and Hawkins, E.: The upper end of climate model temperature projections is inconsistent with past warming, *Environ. Res. Lett.*, 8, 014024, doi:10.1088/1748-9326/8/1/014024, 2013.
- Tebaldi, C., and Arblaster, J. M.: Pattern scaling: Its strengths and limitations, and an update on the latest model simulations, *Climatic Change*, 122, 459–471, doi:10.1007/s10584-013-1032-9, 2014.
- Williams, K. D., Ingram, W. J., and Gregory, J. M.: Time variation of effective climate sensitivity in gcms, *J. Climate*, 21, 5076–5090, doi:10.1175/2008jcli2371.1, 2008.
- Zelinka, M. D., Klein, S. A., Taylor, K. E., Andrews, T., Webb, M. J., Gregory, J. M., and Forster, P. M.: Contributions of different cloud types to feedbacks and rapid adjustments in cmip5, *J. Climate*, 26, 5007–5027, doi:10.1175/Jcli-D-12-00555.1, 2013.