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Accepted Version

Leung, T. Y. ORCID: <https://orcid.org/0000-0003-0056-284X>,
Leutbecher, M., Reich, S. and Shepherd, T. G. ORCID:
<https://orcid.org/0000-0002-6631-9968> (2019) Atmospheric
predictability: revisiting the inherent finite-time barrier. *Journal
of the Atmospheric Sciences*, 76 (12). pp. 3883-3892. ISSN
1520-0469 doi: <https://doi.org/10.1175/JAS-D-19-0057.1>
Available at <https://centaur.reading.ac.uk/86240/>

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To link to this article DOI: <http://dx.doi.org/10.1175/JAS-D-19-0057.1>

Publisher: American Meteorological Society

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1 **Atmospheric predictability: revisiting the inherent finite-time barrier**

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ABSTRACT

13 The accepted idea that there exists an inherent finite-time barrier in deter-
14 ministically predicting atmospheric flows originates from Edward N. Lorenz's
15 1969 work based on two-dimensional (2D) turbulence. Yet, known ana-
16 lytic results on the 2D Navier-Stokes (N-S) equations suggest that one can
17 skilfully predict the 2D N-S system indefinitely far ahead should the initial-
18 condition error become sufficiently small, thereby presenting a potential con-
19 flict with Lorenz's theory. Aided by numerical simulations, the present work
20 re-examines Lorenz's model and reviews both sides of the argument, paying
21 particular attention to the roles played by the slope of the kinetic energy spec-
22 trum. It is found that when this slope is shallower than -3 , the Lipschitz con-
23 tinuity of analytic solutions (with respect to initial conditions) breaks down
24 as the model resolution increases, unless the viscous range of the real system
25 is resolved – which remains practically impossible. This breakdown leads
26 to the inherent finite-time limit. If, on the other hand, the spectral slope is
27 steeper than -3 , then the breakdown does not occur. In this way, the apparent
28 contradiction between the analytic results and Lorenz's theory is reconciled.

29 **1. Introduction**

30 Now an accepted fact in dynamical meteorology, the existence of an inherent finite-time barrier
31 in predicting atmospheric flows was first conceptually shown by Lorenz (1969). Using a simple
32 model, he estimated the predictability limit to be slightly over two weeks – a result echoed by
33 recent studies with real-world operational models at major numerical weather prediction centres
34 (Buizza and Leutbecher 2015; Jutd 2018; Selz 2019; Zhang et al. 2019). Although advances in
35 probabilistic prediction make it possible to extract predictable signals beyond this limit, the ex-
36 tended predictability mainly results from temporal averaging of the predicted fields, together with
37 the slowly varying components of the climate system (Buizza and Leutbecher 2015). Moreover,
38 the loss of information in probabilistic prediction is reflected in the growth of deterministic error,
39 and under statistically stationary conditions, saturation of the error spectrum corresponds to the
40 predicted probability distribution matching that of the climatology, so that the loss of determin-
41 istic and probabilistic predictability are matched. Indeed, empirical evidence suggests that the
42 decay of forecast skill behaves broadly similarly across deterministic and probabilistic predictions
43 (Buizza and Leutbecher 2015). Thus the study of deterministic error growth can be used to un-
44 derstand the mechanisms limiting the range of predictability to a finite horizon and the role of
45 multi-scale interactions in error growth (Rotunno and Snyder 2008; Durran and Gingrich 2014;
46 Sun and Zhang 2016). Such an analysis requires averaging over multiple cases to ensure robust
47 results, which is a somewhat different notion of predictability compared to the fully probabilistic
48 notion used today. However, since the earlier works adopted the deterministic approach with the
49 averaging, we shall take the same approach as we revisit their works.

50 In his original work, Lorenz (1969) classified fluid systems into two categories:

- 51 • those whose error at any future time can be made arbitrarily small by suitably reducing the
52 initial error, and
- 53 • those whose error at any future time cannot be reduced below a certain limit unless the initial
54 error is zero.

55 Using an appropriate skill function, these systems can be equivalently characterised in terms of
56 *range of predictability* (or simply *predictability*): the former category has an indefinite range and
57 the latter has only an inherently finite range. The reader is referred to Appendix A for a motivation
58 of the concept and more details about the skill function. By modelling atmospheric flows by the
59 two-dimensional (2D) barotropic vorticity equation and assuming a $-\frac{5}{3}$ spectral slope along the
60 inertial range of the kinetic energy (KE) spectrum of the unperturbed flow, he argued that such
61 flows have an inherently finite range of predictability. (The inertial range is a continuous part
62 of the spectrum where a specific power-law relationship is followed so that the flow restricted to
63 such scales is self-similar. It is identified by a spectral slope which is the slope of the spectrum as
64 appearing in a log-log plot.)

65 Although the barotropic vorticity equation with large-scale forcing produces a steeper spectral
66 slope of -3 , and unbalanced dynamics are required to produce a spectral slope of $-\frac{5}{3}$ in more
67 realistic models (Sun and Zhang 2016), it has been shown that predictability is determined much
68 more by the spectral slope than by the nature of the dynamics (Rotunno and Snyder 2008). Thus, it
69 is appropriate to use the barotropic vorticity equation to study predictability with a range of spectral
70 slopes, recognising that this only addresses one aspect of what limits atmospheric predictability in
71 practice.

72 Closely related to this system are the incompressible 2D Navier-Stokes (2D N-S) equations,
73 whose well-posedness (existence of a unique solution to the initial-value problem that depends

continuously on the initial conditions) was first rigorously shown by Ladyzhenskaya also in the second half of the twentieth century (Robinson 2001). As we will see in Section 4, it is not difficult to show that well-posedness implies an indefinite range of predictability in the sense of Lorenz.

The present paper aims to reconcile the difference between the inherently finite predictability result of Lorenz and the indefinite predictability corollary of Ladyzhenskaya’s proof, in the context of incompressible 2D flows. Section 2 reviews Lorenz’s argument of its inherent finite-time behaviour. In Section 3 we reproduce Lorenz’s numerical results and discuss the predictability in the directly simulated 2D barotropic vorticity model. An account of the well-posedness and indefinite predictability of the incompressible 2D N-S equations is presented in Section 4, with which we reconcile Lorenz’s result of inherently finite predictability in Section 5. The major findings are summarised in Section 6.

2. Lorenz’s argument of inherently finite predictability

The model of Lorenz (1969) is based on the dimensionless 2D barotropic vorticity equation

$$\frac{\partial \theta}{\partial t} + J(\psi, \theta) = 0, \quad \theta = \Delta \psi \quad (1)$$

where ψ is the velocity streamfunction (related to the velocity \mathbf{u} by $\mathbf{u} = -\nabla \times (\psi \hat{\mathbf{k}})$), $\Delta = \nabla \cdot \nabla$, $\nabla = \left(\frac{\partial}{\partial x}, \frac{\partial}{\partial y} \right)$ and $J(A, B) = \frac{\partial A}{\partial x} \frac{\partial B}{\partial y} - \frac{\partial A}{\partial y} \frac{\partial B}{\partial x}$. Assuming a doubly periodic domain, Lorenz expanded the variables ψ and θ in Fourier series and re-wrote the linearised error equation of (1) in Fourier components. Then he made various assumptions to an ensemble of error fields for the linearised error equation of (1), most notably homogeneity and a slight generalisation of the quasi-normal closure. The resulting equation was then passed into the large-domain and continuous-spectrum limit.

94 The derivation is arguably more straightforward if the domain is the whole \mathbb{R}^2 space and the
 95 variables are Fourier-transformed rather than expanded in Fourier series. We have checked that
 96 this method returns the same equation as the limiting equation of Lorenz, up to a constant multi-
 97 plicative factor.

98 A further assumption of isotropy simplifies the equation, which was then discretised and numer-
 99 ically approximated. Depending on the specification of a KE spectrum for the unperturbed flow,
 100 a matrix of constant coefficients C was constructed so that the vector Z of error KE at different
 101 scales (each scale K collectively represents wavenumbers $k = 2^{K-1}$ to $k = 2^K$) evolves according
 102 to the linear model

$$\frac{d^2}{dt^2}Z = CZ, \quad \text{or equivalently} \quad \frac{d}{dt} \begin{pmatrix} Z \\ W \end{pmatrix} = \begin{pmatrix} 0 & I \\ C & 0 \end{pmatrix} \begin{pmatrix} Z \\ W \end{pmatrix}, \quad (2)$$

103 where W is the first time-derivative of Z .

104 As Rotunno and Snyder (2008) mentioned, the computation of C involves computing integrals
 105 of nearly singular functions. We have been cautious about these integrations and have made sure
 106 that our integrations for C are accurate, some details of which are provided in Appendix B.

107 To time-integrate equation (2), it is necessary that the initial conditions for Z and W are speci-
 108 fied. Lorenz did not explicitly give an initial condition for W , but as Rotunno and Snyder (2008)
 109 assumed $W(t = 0) \equiv 0$ in their predictability experiments, we shall prescribe the same for our
 110 numerical simulations (Section 3(a)). The non-linear effects were accounted for by removing the
 111 corresponding components of Z , W and C when the error KE saturated at a particular scale, where-
 112 upon an inhomogeneous forcing term was added to the right-hand-side of equation (2) to account
 113 for the saturated scale's contribution to the error growth at the unsaturated scales (details available
 114 in Appendix C). Time-integration with the resulting lower-dimensional system was carried on,
 115 until all scales became saturated. The evolution of the error KE spectrum in time is depicted in

116 Figure 1 (for illustrative purposes only; this is for a spectral slope of -3 whose simulation Lorenz
 117 excluded in his original work).

118 As Lorenz noted down the saturation times t_K of scale K , he found that the successive differences
 119 $t_K - t_{K+1}$ behaved approximately proportional to $2^{-\beta K}$ with β depending on the spectral slope.
 120 He therefore concluded that, given an initial error at an infinitesimally small scale, the range of
 121 predictability is inherently finite if and only if the telescoping series

$$t_K = \sum_{j=K}^{\infty} (t_j - t_{j+1}) = \sum_{j=K}^{\infty} 2^{-\beta j} \quad (3)$$

122 is summable, which is the case if and only if $\beta > 0$. By observing $\beta = \frac{2}{3}$ for the atmospherically
 123 relevant spectral slope of $-\frac{5}{3}$, he suggested inherently finite predictability for the atmosphere. Ad-
 124 ditionally, he found that $\beta = \frac{1}{3}$ for a hypothetical spectral slope of $-\frac{7}{3}$. Lorenz thus hypothesised
 125 by linear extrapolation that the range of predictability would be indefinite if the spectral slope were
 126 steepened to -3 .

127 This result is echoed by arguments on dimensional grounds (Vallis 1985; Lilly 1990). Assuming
 128 that t_K in equation (1) depends only on the wavenumber k and the one-dimensional KE spectral
 129 density $E(k)$ of the background flow, one has $t_K \sim (k^3 E(k))^{-0.5}$ as this is the only way the physical
 130 units of k and $E(k)$ can combine to give the dimension of time. With $k \sim 2^K$, one obtains $t_K \sim$
 131 $2^{-\beta K}$, with $\beta = \frac{2}{3}, \frac{1}{3}$ and 0 for the spectral slopes of $-\frac{5}{3}, -\frac{7}{3}$ and -3 respectively, the same result
 132 as Lorenz's.

133 3. Numerical simulations

134 We performed a series of numerical simulations, first on the Lorenz model (2) followed by a
 135 forced-dissipative version of the full 2D barotropic vorticity system (1), to see whether indefinite
 136 predictability is indeed achieved with a KE spectral slope of -3 as Lorenz hypothesised.

137 *a. Lorenz's model*

138 Rotunno and Snyder (2008) solved for the growth of the error KE spectrum for a background
139 spectral slope of $-p$ where $p = 3$. In order to assess the range of predictability in Lorenz's frame-
140 work, we extended their calculations to study the relationship between K and t_K .

141 Having computed the matrix C as in Rotunno and Snyder (2008), we solved the linear matrix
142 system (2) explicitly, that is, by writing out the general solution in terms of the eigenvalues and
143 eigenvectors of

$$\begin{pmatrix} 0 & I \\ C & 0 \end{pmatrix}$$

144 and projecting the initial condition onto such an eigenspace to determine the constants of the
145 general solution. This exact approach is a good and easy alternative to the numerical schemes
146 used by Lorenz (1969), Rotunno and Snyder (2008) and its extension by Durran and Gingrich
147 (2014). Details of the solution procedure can be found in Appendix C.

148 Figure 1 shows the evolution of the error for the -3 spectrum as in Rotunno and Snyder (2008),
149 and Figure 2 shows the saturation times t_K as a function of the scale K . Note that in Figure 2 t_K
150 is plotted instead of $t_K - t_{K+1}$ against K , but the choice makes little difference when $\beta > 0$ since
151 if $t_K - t_{K+1}$ is proportional to $2^{-\beta K}$ then so is t_K (cf. equation (3)). It is clear that the saturation
152 times t_K scale as $2^{-\beta K}$ with a small but positive β (0.05) along the inertial range, so that the sum
153 in equation (3) is still finite for $p = 3$, contrary to Lorenz's prediction. Indeed, arguing in the same
154 way as Lorenz, our result indicates inherently finite predictability for a -3 spectrum which is
155 contrary to Lorenz's hypothesis, although we acknowledge that $\beta = 0.05$ is just marginally away
156 from the critical value of zero. We did, however, recover Lorenz's result for the case of a $-\frac{5}{3}$
157 spectrum (not shown).

158 *b. Forced-dissipative 2D barotropic vorticity equation*

159 The difference between our value of β and Lorenz’s for the -3 spectrum deserves particular
 160 attention because it amounts to a qualitative contrast between inherently finite and indefinite ranges
 161 of predictability. To further investigate this, we performed direct numerical simulations (DNS) on
 162 this $p = 3$ spectrum in the form of identical-twin experiments (pairs of runs which only differ in the
 163 initial condition), and assessed the predictability following Lorenz’s methodology with necessary
 164 adaptations.

165 First of all, equation (1) had to be restricted to a doubly periodic domain and be made forced-
 166 dissipative:

$$\frac{\partial \theta}{\partial t} + J(\psi, \theta) = f + d, \quad \theta = \Delta \psi. \quad (4)$$

167 The forcing and dissipation, however small, are necessary for generating statistically stationary
 168 KE spectra in the DNS. To generate a -3 spectral slope, following standard practice (Maltrud
 169 and Vallis 1991), forcing was applied at the large scale: $f(t)$ was chosen to be an independent
 170 white-noise process for each 2D wavevector whose scalar wavenumber k falls in the narrow band
 171 ($\pm 10\%$) around $k = 20$. The dissipation d was a highly scale-selective hyperviscosity $d \sim -\Delta^6 \theta$.

172 It is worth noting that equation (4) would also be the vorticity form of the incompressible 2D
 173 N-S equations

$$\frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla) \mathbf{u} = -\nabla p + f(x, t) + \nu \Delta \mathbf{u}, \quad \nabla \cdot \mathbf{u} = 0 \quad (5)$$

174 if d were chosen to be $d = \nu \Delta \theta$, $\nu > 0$. We would have liked to run these DNS on the 2D N-S
 175 equations which will be discussed in Section 4, but the build-up of KE at the smallest scales as
 176 a numerical artefact was so strong that we had to either increase ν – which would substantially
 177 shorten the inertial range and thus reduce the reliability of our conclusions – or choose a more

178 scale-selective dissipation. We opted for the latter, as is standard practice in simulations of 2D
179 turbulence (Maltrud and Vallis 1991).

180 We performed five pairs of identical-twin experiments on equation (4) by varying the random
181 seed that generated the pre-perturbation (original) initial condition. Within each pair, notably,
182 the realisations of the large-scale stochastic forcing $f(t)$ in the control and perturbed runs were
183 identical. The model was pseudo-spectral with a truncation wavenumber of $k_t = 512$, in which the
184 $J(\psi, \theta)$ term was computed in the physical domain via a pair of Fast Fourier Transforms with the
185 spectral de-aliasing filter proposed by Hou and Li (2007). The original initial condition for each
186 of the five cases was an already-developed homogeneous and (approximately) isotropic turbulence
187 with a clean logarithmically corrected -3 spectrum in the inertial range (Figure 3), which has been
188 shown to be a more accurate description of the large-scale-forced 2D turbulent spectrum for finite
189 inertial ranges (Bowman 1996).

190 The perturbations were introduced spectrally at each of the 2D wavevectors \mathbf{k} for a specified
191 value of $k_p = |\mathbf{k}|$. A random phase shift independently drawn from a uniform distribution was
192 applied on a pre-determined part $\gamma \in [0, 1]$ of the spectral coefficients $\hat{\psi}(\mathbf{k})$ and thus $\hat{\theta}(\mathbf{k})$, where
193 the hat indicates Fourier coefficients. It can be shown that $\gamma(k_p)$ and $\mathbb{E}(E_e(k_p))$, the expected value
194 of the one-dimensional error KE spectral density at wavenumber k_p , are related by $\mathbb{E}(E_e(k_p)) =$
195 $2\gamma^2 E(k_p)$. By specifying the relative error $\frac{\mathbb{E}(E_e(k_p))}{E(k_p)}$, we could work out γ and thus generate the
196 perturbation fields, to which we added the original initial conditions to obtain the perturbed initial
197 conditions.

198 Since our truncation wavenumber $k_t = 512$ corresponds to $K = 9$ of Lorenz's paper, there would
199 only be 8 values of $t_K - t_{K+1}$, among which only 4 or 5 would be in the inertial range. It would be
200 inaccurate to determine β from such a few data points, so we have transformed Lorenz's argument

201 to incorporate information from all wavenumbers k , not only from the scale K as a collection of
202 wavenumbers.

203 To transform the argument, recall that $t_K \sim 2^{-\beta K}$ when $\beta > 0$, and let $T(k)$ be the saturation
204 time of wavenumber k . Since $k \sim 2^K$ and both t_K and T represent saturation times, we may write
205 $T \sim k^{-\beta}$ and conclude that the T should vary with k as a power-law if Lorenz's results hold.
206 The argument will break down when β becomes zero, that is, when the threshold for indefinite
207 predictability is reached.

208 In this study, the perturbations were introduced at $k_p = 256$. The saturation threshold was chosen
209 to be 1.315 times the KE spectral density of the control flow, or equivalently 0.6575 times the
210 maximum permissible error energy, in accordance with Lorenz (1969). We applied sensitivity tests
211 and found that the results are largely insensitive to the saturation threshold. Figures 4 and 5 show
212 respectively the evolution of the error KE spectrum, and the saturation times T across different
213 wavenumbers k which fit the $T \sim k^{-\beta}$ relationship for $\beta = 0.24$, averaged over the five cases.
214 (The five cases exhibited very similar qualitative behaviour, showing that our results are robust to
215 initial conditions, hence justifying the use of averaging to obtain smoother results.) Based on the
216 transformed version of Lorenz's argument, our result also suggests inherently finite predictability
217 for a (logarithmically corrected) -3 spectrum, this time with greater confidence as β is further
218 away from zero.

219 **4. Aspects from PDE theory: the incompressible 2D Navier-Stokes equations**

220 A very different approach to the problem of inherently finite or indefinite predictability is via
221 use of the analytic theory of partial differential equations (PDEs). The incompressible 2D N-S
222 equations (5), where we shall drop the word 'incompressible' for the remainder of the paper, are
223 always useful as a pedagogical first step towards understanding and modelling the motion of real

224 fluid flows in the atmosphere. As such, the analytical properties of the 2D N-S problem have been
225 extensively studied. Building on these analytic results, we now consider their implications for
226 predictability.

227 *Well-posedness and implications on predictability*

228 Unlike their three-dimensional counterpart whose regularity problem remains open, the initial-
229 value problem for the 2D N-S equations on the torus (i.e. a doubly periodic domain) has been
230 proven to be well-posed, by which we mean the existence of a unique solution that depends con-
231 tinuously on the initial conditions. Proofs of its well-posedness, for both strong and weak solutions
232 respectively, can be found in the book by Robinson (2001). In the present paper we shall use his
233 proof for weak solutions to demonstrate that the 2D N-S system is indefinitely predictable. To set
234 the context, a summary of the uniqueness proof is provided below. Interested readers may refer to
235 Robinson's book for a full proof.

236 First, the 2D N-S equations (5) are cast in the form of an ordinary differential equation in an
237 appropriate function space depending on an arbitrary, fixed positive time T . An equation for the
238 error velocity field $w = u - v$ of two solutions u and v is formulated, and its inner product with
239 w itself is taken to obtain an equation for the time-evolution of the error energy $\frac{1}{2}\|w\|^2$, where
240 $\|\cdot\|$ is the L^2 norm on the torus. This equation contains a term which can be bounded above by
241 Ladyzhenskaya's inequalities (Robinson 2001) specific to the 2D N-S equations. After some work
242 one uses Grönwall's inequality to show that

$$\|w(t)\|^2 \leq \exp\left(\int_0^t \frac{M}{\nu} \|\nabla u(s)\|^2 ds\right) \|w(0)\|^2, \quad t \in [0, T],$$

243 where M is a positive constant provided by Ladyzhenskaya's inequalities. Uniqueness follows by
 244 setting $w(0) = 0$. One can also show the continuous dependence on initial conditions, since

$$\|w(t)\| \leq \sqrt{\exp\left(\int_0^T \frac{M}{\nu} \|\nabla u(s)\|^2 ds\right)} \|w(0)\| =: L(T) \|w(0)\|, \quad t \in [0, T], \quad (6)$$

245 i.e. errors are Lipschitz in time.

246 As an immediate corollary to inequality (6), the 2D N-S system is indefinitely predictable
 247 (Palmer et al. 2014). Indeed, if a prediction is defined to lose its skill when $\|w(t)\| > \varepsilon$, then
 248 for any given time $T \in \mathbb{R}^+$, the prediction is skilful for at least up to T when the initial error
 249 $\|w(0)\|$ can be made sufficiently small, that is, smaller than $\frac{\varepsilon}{L(T)}$.

250 It is important to note that in the present Section the KE spectral slope plays no role in determin-
 251 ing the inherent finiteness or indefiniteness of predictability of the 2D N-S equations. The above
 252 argument applies to 2D N-S systems of any spectral slope.

253 5. Reconciling the contradiction with Lorenz

254 At first glance, the indefinite predictability derived in Section 4 seems to contradict Lorenz's
 255 result in Section 2 for any $p < 3$. However, we have not discussed the role of the KE spectral slope
 256 in $L(T)$ which, as we will see in the following, reconciles the conflict.

257 Central to our argument is the inequality (6) presented above. For simplicity, suppose the real
 258 system has only one inertial range of slope $-p$ (without any logarithmic correction) in its KE
 259 spectrum so that $|\hat{u}(k)|^2 \sim k^{-p}$ (note the change of notation: the hat now represents Fourier coeffi-
 260 cients in the space of one-dimensional wavenumbers k) between its large-scale cutoff wavenumber
 261 k_1 and small-scale cutoff wavenumber k_2 . Then

$$\|\nabla u_s\|^2 = \int_0^\infty k^2 |\hat{u}_s|^2 dk = \int_0^{k_1} k^2 |\hat{u}_s|^2 dk + A_0 \int_{k_1}^{k_2} k^{2-p} dk + \int_{k_2}^\infty k^2 |\hat{u}_s|^2 dk \quad (A_0 \text{ constant}), \quad (7)$$

262 where the subscript s distinguishes the system itself from a model for the system which we will
 263 denote with subscript m . The three terms on the right-hand-side of equation (7) represent contri-
 264 butions from the large scale, the inertial range and the viscous range respectively. Compared to
 265 the first two terms, the term representing the viscous range is assumed to be small. In particular,
 266 the integrand is assumed to decay rapidly enough so that $\|\nabla u_s\|^2$ remains finite (this is in fact part
 267 of the definition of the function space to which u_s belongs).

268 Now, suppose the model truncates at wavenumber $k_t \ll k_2$ and numerical dissipation kicks in at
 269 wavenumber $k_0 \in (k_1, k_t)$. For the model,

$$\|\nabla u_m\|^2 = \int_0^{k_t} k^2 |\hat{u}_m|^2 dk = \int_0^{k_1} k^2 |\hat{u}_m|^2 dk + A_0 \int_{k_1}^{k_0} k^{2-p} dk + \int_{k_0}^{k_t} k^2 |\hat{u}_m|^2 dk. \quad (8)$$

270 Again, we may neglect the contribution from the viscous range, so that

$$\|\nabla u_m\|^2 \sim \int_0^{k_1} k^2 |\hat{u}_m|^2 dk + A_0 \int_{k_1}^{k_0} k^{2-p} dk. \quad (9)$$

271 Because $k_0, k_t \ll k_2$, the second integral in relation (9) with $p < 3$ *appears* to diverge as the res-
 272 olution (k_0, k_t) increases. Combining this with inequality (6), $L(T)$ – until k_2 is reached – grows
 273 exponentially with k_0 , leading to a breakdown of the Lipschitz-continuous dependence on initial
 274 conditions in inequality (6). To keep the error $\|w(t)\|$ under control, the initial error $\|w(0)\|$ would
 275 have to decrease exponentially, but decreasing the scale of the initial error without changing its
 276 magnitude relative to the background KE spectral density (Lorenz’s thought experiment) would
 277 only give a polynomial decrease. The corollary of indefinite predictability therefore fails to hold.
 278 Hence the range of predictability is inherently finite in practice, even though the system is indefi-
 279 nitely predictable, because indefinite predictability cannot be achieved without making the model
 280 resolution so high that its effective resolution k_0 (and the scale of the initial error) falls within
 281 the viscous range of the real system, let alone the large-scale error has to be constrained to zero
 282 (Durran and Gingrich 2014).

283 This concept, known as ‘asymptotic ill-posedness’, was put forward by Palmer et al. (2014) as
 284 they argued that whether the three-dimensional Navier-Stokes system is well-posed is practically
 285 irrelevant to the well-established theory of inherently finite predictability. We have now extended
 286 the discussion to the 2D N-S system and given a mathematical basis to the concept in our context.

287 When $p > 3$, the second integral in relation (9) does not appear to diverge as $k_0 \rightarrow k_2$. This means
 288 one may indeed approximate $\|\nabla u_s\|^2$ by the $\|\nabla u_m\|^2$ in relation (9) with a sufficiently large value
 289 of k_0 . So would $L(T)$ of inequality (6) be approximated without regard to the model resolution,
 290 making it possible for $\|w(t)\| \leq \varepsilon$ by making $\|w(0)\|$ small enough in scale and thus achieving
 291 indefinite predictability.

292 So far our argument for the cases $p < 3$ and $p > 3$ are in harmony with Lorenz’s result in Section
 293 2. For the borderline case $p = 3$, our argument suggests practically inherently finite predictability,
 294 since $\|\nabla u_m\|^2 \sim \text{constant} + \int_{k_1}^{k_0} k^{-1} dk = \text{constant} + \log \frac{k_0}{k_1}$ which appears to diverge as $k_0 \rightarrow k_2$.
 295 This disagrees with Lorenz. Even with the logarithmic correction

$$|\hat{u}(k)|^2 \sim k^{-3} \left[\log \left(\frac{k}{k_r} \right) \right]^{-\frac{1}{3}} \quad (k_r > 0 \text{ constant}),$$

296 or more generally

$$|\hat{u}(k)|^2 \sim k^{-3} \left[A_1 \log \left(\frac{k}{k_r} \right) + A_2 \right]^{-\frac{1}{3}} \quad (A_1, A_2, k_r > 0 \text{ constants}),$$

297 to the -3 spectrum (Bowman 1996), an easy calculation along the previous lines still suggests that
 298 the range of predictability is practically inherently finite. As such, we are unable to explain the
 299 disagreement and we leave the problem open.

300 For models and systems with multiple inertial ranges, only the range immediately before the
 301 viscous range pertains to our argument concerning the large- k_0 behaviour. This applies to the real
 302 atmosphere where $p = \frac{5}{3}$ (Nastrom and Gage 1985). Since k_f for atmospheric models is smaller
 303 than k_2 by ‘at least seven or eight orders of magnitude’ (Palmer et al. 2014), the crucial assumption

304 to our discussion ($k_1 \ll k_2$) is satisfied and we conclude that atmospheric predictability is indeed
305 practically inherently finite.

306 6. Conclusions

307 Half a century on since Lorenz’s pioneering result of inherently finite atmospheric predictability,
308 we revisited his original argument by (i) re-running his simplified model of the 2D barotropic
309 vorticity equation, (ii) directly simulating the full model and (iii) comparing his conclusions with
310 the well-posedness of the 2D N-S equations as proven by Ladyzhenskaya.

311 Although his main result – that atmospheric predictability is inherently finite because the KE
312 spectral slope is shallower than -3 – has now become an ‘accepted part of the canon of dynamical
313 meteorology’ (Rotunno and Snyder 2008), the details behind the conclusion were re-assessed. For
314 the -3 spectrum, we saw a substantially different β in the DNS (Section 3(b)) than in the simpli-
315 fied model (Section 3(a)), which may be an indication that the model is inadequate in simulating
316 error growth. In both cases, nevertheless, the hypothesis of indefinite predictability ($\beta = 0$) for
317 $p = 3$ based on linear extrapolation (Section 2) was refuted.

318 The 2D N-S equations that closely relate to the 2D barotropic vorticity equation were used to ad-
319 dress the predictability problem from a more rigorous perspective. The forced-dissipative system
320 was shown to be indefinitely predictable regardless of the spectral slope (Section 4). However, we
321 found that $p = 3$ serves as a cutoff between *practically* inherently finite and indefinite predictabil-
322 ity by noting how quickly the initial error has to be brought down with increasing resolution in
323 order to maintain the bound for the error at future times (Section 5). This echoes Lorenz’s original
324 conclusions except for the borderline case $p = 3$ itself, in which case our result of inherently finite
325 predictability agrees with our own computations of Lorenz’s model and the DNS.

326 Until recently, KE spectra in global weather forecast models had only resolved the synoptic-
327 scale -3 range. As model resolutions start to extend into the $-\frac{5}{3}$ range, the strong constraints on
328 the range of predictability envisaged by Lorenz will become visible (Judt 2018). However, the
329 limits on predictability arising from initial errors on the large scales will also limit predictability
330 in practice (Durrán and Gingrich 2014), and the interplay between the two could be an interesting
331 area to explore.

332 By providing another approach to attacking the problem of predictability (via the analytic the-
333 ory of the 2D N-S equations), we look forward to similar results on more atmospherically relevant
334 PDEs such as the surface quasi-geostrophic equations (Held et al. 1995), and a more active contri-
335 bution from mathematicians on this topic.

336 *Acknowledgments.* Tsz Yan Leung is supported through a PhD scholarship awarded by the En-
337 gineering and Physical Sciences Research Council grant EP/L016613/1 ‘EPSRC Centre for Doc-
338 toral Training in the Mathematics of Planet Earth at Imperial College London and the University of
339 Reading’, with additional funding support from the European Research Council Advanced Grant
340 ‘Understanding the Atmospheric Circulation Response to Climate Change’ (ACRCC), Project
341 339390, under Theodore G. Shepherd as the Principal Investigator. The authors thank Richard
342 Scott for providing his code for modification for the numerical simulation in Section 3(b). The
343 authors also wish to thank the three anonymous reviewers for their helpful and valuable comments
344 on an earlier version of this manuscript.

345 APPENDIX A

346 **Motivating the concept *range of predictability***

347 Standard theory in deterministic dynamical systems dictates that, for a perfect model of the
 348 system, the dynamics of the error ε can be completely described by the time t , the initial error
 349 ε_0 and the initial state of the system x , so that $\varepsilon = \varepsilon(t, \varepsilon_0, x)$. Suppose now that the skill S of a
 350 forecast, particular to an initial state, can be quantified by a continuously decreasing function of
 351 some norm $\|\cdot\|$ (such as the kinetic energy) of the error. In such a case, we can write $S = S(\|\varepsilon\|) =$
 352 $S(\varepsilon) = S(t, \varepsilon_0, x)$. Now, averaging over the initial states on some (non-trivial) attractor D of the
 353 system, we may define the overall skill $\bar{S} = \int_{x \in D} S(t, \varepsilon_0, x) dx$ of a forecast. If we further assume
 354 that the error norm increases with t in the average, as is in the context of atmospheric predictions,
 355 we can infer that $\bar{S}(t, \varepsilon_0)$ is monotonically decreasing in time.

356 Let's say that a prediction loses its skill when $\bar{S} < \alpha$ (where α is a fixed threshold), which
 357 is typically realised in fluid flows by saturation of the error kinetic energy spectrum at specified
 358 scales. Perhaps a first question to the understanding of predictability can be formulated as follows:
 359 how long does it take for an initial error ε_0 to grow so that a prediction is no longer skilful? The
 360 answer \tilde{T} , known as the range of predictability, is the solution to $\bar{S}(t, \varepsilon_0) = \alpha$ for the specified ε_0 .
 361 The monotonicity assumption of \bar{S} guarantees the uniqueness of the solution \tilde{T} .

362 By formulating this question for different initial error fields we can regard \tilde{T} as a function of
 363 ε_0 . It is clear from the very definition of deterministic systems that $\varepsilon_0 = 0$ implies $\tilde{T}(\varepsilon_0) = \infty$.
 364 However, it is not quite obvious as to whether \tilde{T} could be made arbitrarily large by reducing $\|\varepsilon_0\|$ to
 365 anything positive below a threshold, or equivalently whether the equality $\liminf_{\|\varepsilon_0\| \rightarrow 0} \tilde{T}(\varepsilon_0) = \infty$
 366 holds, because \tilde{T} may behave irregularly at small $\|\varepsilon_0\|$ – or at least appear to.

367 To see the equivalence, we unwrap the statement $\liminf_{\|\varepsilon_0\| \rightarrow 0} \tilde{T}(\varepsilon_0) = \infty$ to get

$$\liminf_{\|\varepsilon_0\| \rightarrow 0} \tilde{T}(\varepsilon_0) = \infty$$

$$\Leftrightarrow \forall R \in \mathbb{R}, \liminf_{\|\varepsilon_0\| \rightarrow 0} \tilde{T}(\varepsilon_0) \geq R$$

$$\Leftrightarrow \forall R \in \mathbb{R}, \sup_{\varepsilon' > 0} \inf_{\|\varepsilon_0\| \in (0, \varepsilon')} \tilde{T}(\varepsilon_0) \geq R$$

$$\Leftrightarrow \forall R \in \mathbb{R}, \exists \varepsilon' > 0 \text{ such that } \inf_{\|\varepsilon_0\| \in (0, \varepsilon')} \tilde{T}(\varepsilon_0) \geq R$$

$\Leftrightarrow \tilde{T}$ could be made arbitrarily large by reducing $\|\varepsilon_0\|$ to anything positive below a threshold.

368 With this in mind, a system is said to have an *indefinite range of predictability*, or be *indefinitely*
 369 *predictable*, if the range of predictability could be made arbitrarily large by reducing the initial
 370 error to a small yet positive value. Systems that fail to satisfy such a condition are referred to as
 371 *inherently finitely predictable*.

372 APPENDIX B

373 **Some details regarding the computation of the matrix C**

374 The integrations were performed using `scipy.integrate.nquad` on Python which returned
 375 a warning message ‘IntegrationWarning: Extremely bad integrand behavior occurs
 376 at some points of the integration interval’ about the integrand’s singular behaviour,
 377 even if the integration domain were confined to the support of the integrand so that resources
 378 were not wasted in integrating zero regions. The warning disappeared by applying a change of
 379 coordinates (from logarithmic to Cartesian) in the integrand and accordingly the integration limits,
 380 which sped up the wall-clock time of the computation by a factor of about 9 as well. The entries of
 381 C computed by these two methods differ by no more than 0.0025%. Based on these observations,
 382 we are confident that our computations are accurate.

383 The numerical code for these computations is available in the Supplemental Material.

384 APPENDIX C

385 **Solution procedure of Lorenz's system**

- 386 1. Set a time-step h ; in this case, $h = 0.001$.
- 387 2. Initialise the run by setting time $t = 0$. Also initialise $t_0 = 0$. (t_0 is the time when the previous
388 saturation occurs.)
- 389 3. Project the initial condition onto the eigenspace of the block matrix

$$\begin{pmatrix} 0 & I \\ C & 0 \end{pmatrix}$$

390 to determine the constants of the general solution.

- 391 4. Compute the solution at time $t - t_0$ and check if any of the scales K saturates by time t .
- 392 5. If none of the scales saturates, reset $t = t + h$ and repeat step 4.
- 393 6. If, by time t , the error energy at some scale exceeds the background energy at the same scale,
394 then the clock (i.e. the variable t) is reset to the previous time-step $t - h$, and h is refined to
395 0.000001.
- 396 7. Repeat steps 4-5 with the new value of h until the error energy at some scale K exceeds the
397 background energy at the same scale. The saturation time of scale K is determined as if the
398 error energy increases linearly between times $t - h$ and t .
- 399 8. Reset $h = 0.001$ and set t_0 to be the current time t . Remove the row and column of the matrix
400 C corresponding to scale K and the corresponding entries of Z , W and all the F_j terms (more

401 on the F_j terms below). The reduced-size system

$$\frac{d^2}{dt^2}Z = CZ + \sum_j F_j$$

402 (where the summation is performed over all saturated scales) has a new inhomogeneous term
 403 F_K which accounts for the contribution of scale K 's saturated energy to the error growth at the
 404 remaining scales. (F_K equals the background energy of scale K multiplied by the removed
 405 column of C restricted to the rows corresponding to the remaining scales.) As the new system
 406 is equivalent to the first-order system

$$\frac{d}{dt} \begin{pmatrix} Z \\ W \end{pmatrix} = \begin{pmatrix} 0 & I \\ C & 0 \end{pmatrix} \begin{pmatrix} Z \\ W \end{pmatrix} + \begin{pmatrix} 0 \\ \sum_j F_j \end{pmatrix}$$

407 (with the size of the identity matrix I accordingly reduced) and $\sum_j F_j$ is a constant vector,
 408 its solution $(Z, W)^T$ can be expressed as the sum of a particular solution $(\Delta Z, \Delta W)^T$ and a
 409 solution of of the homogeneous system $(\bar{Z}, \bar{W})^T$. A particular solution to the differential
 410 equation can be obtained by solving

$$- \begin{pmatrix} 0 & I \\ C & 0 \end{pmatrix} \begin{pmatrix} \Delta Z \\ \Delta W \end{pmatrix} = \begin{pmatrix} 0 \\ \sum_j F_j \end{pmatrix}.$$

411 Re-calculate the eigenvalues and eigenvectors of the new

$$\begin{pmatrix} 0 & I \\ C & 0 \end{pmatrix}$$

412 and project the difference between the current solution at the unsaturated scales (treated as the
 413 initial condition in the variable $t - t_0$) and $(\Delta Z, \Delta W)^T$ onto such an eigenspace to determine
 414 the constants of $(\bar{Z}, \bar{W})^T$ and thus the full solution.

415 9. Repeat steps 4–8 until all scales saturate.

416 The numerical code for the solution procedure is available in the Supplemental Material.

417 **References**

- 418 Bowman, J. C., 1996: On inertial-range scaling laws. *J. Fluid Mech.*, **306**, 167–181, doi:10.1017/
419 S0022112096001279.
- 420 Buizza, R., and M. Leutbecher, 2015: The forecast skill horizon. *Quart. J. Roy. Meteor. Soc.*, **141**,
421 3366–3382, doi:10.1002/qj.2619.
- 422 Durran, D. R., and M. Gingrich, 2014: Atmospheric predictability: why butterflies are not of
423 practical importance. *J. Atmos. Sci.*, **71**, 2476–2488, doi:10.1175/JAS-D-14-0007.1.
- 424 Held, I. M., R. T. Pierrehumbert, S. T. Garner, and K. L. Swanson, 1995: Surface quasi-geostrophic
425 dynamics. *J. Fluid Mech.*, **282**, 1–20, doi:10.1017/S0022112095000012.
- 426 Hou, T. Y., and R. Li, 2007: Computing nearly singular solutions using pseudo-spectral methods.
427 *Journal of Computational Physics*, **226**, 379–397, doi:10.1016/j.jcp.2007.04.014.
- 428 Judt, F., 2018: Insights into atmospheric predictability through global convection-permitting
429 model simulations. *J. Atmos. Sci.*, **75**, 1477–1497, doi:10.1175/JAS-D-17-0343.1.
- 430 Lilly, D. K., 1990: Numerical prediction of thunderstorms – has its time come? *Quart. J. Roy.*
431 *Meteor. Soc.*, **116**, 779–798, doi:10.1002/qj.49711649402.
- 432 Lorenz, E. N., 1969: The predictability of a flow which possesses many scales of motion. *Tellus*,
433 **21**, 289–307, doi:10.3402/tellusa.v21i3.10086.
- 434 Maltrud, M. E., and G. K. Vallis, 1991: Energy spectra and coherent structures in forced
435 two-dimensional and beta-plane turbulence. *J. Fluid Mech.*, **228**, 321–342, doi:10.1017/
436 S0022112091002720.

437 Nastrom, G. D., and K. S. Gage, 1985: A climatology of atmospheric wavenumber spectra of
438 wind and temperature observed by commercial aircraft. *J. Atmos. Sci.*, **42**, 950–960, doi:10.
439 1175/1520-0469(1985)042<0950:ACOAWS>2.0.CO;2.

440 Palmer, T. N., A. Döring, and G. Seregin, 2014: The real butterfly effect. *Nonlinearity*, **27**, R123–
441 R141, doi:10.1088/0951-7715/27/9/R123.

442 Robinson, J. C., 2001: *Infinite-Dimensional Dynamical Systems: An Introduction to Dissipative*
443 *Parabolic PDEs and the Theory of Global Attractors*. Cambridge University Press, 234–257 pp.

444 Rotunno, R., and C. Snyder, 2008: A generalization of Lorenz’s model for the predictability of
445 flows with many scales of motion. *J. Atmos. Sci.*, **65**, 1063–1076, doi:10.1175/2007JAS2449.1.

446 Selz, T., 2019: Estimating the intrinsic limit of predictability using a stochastic convection scheme.
447 *J. Atmos. Sci.*, **76**, 757–765, doi:10.1175/JAS-D-17-0373.1.

448 Sun, Y. Q., and F. Zhang, 2016: Intrinsic versus practical limits of atmospheric predictabil-
449 ity and the significance of the butterfly effect. *J. Atmos. Sci.*, **73**, 1419–1438, doi:10.1175/
450 JAS-D-15-0142.1.

451 Vallis, G. K., 1985: Remarks on the predictability properties of two- and three-dimensional flow.
452 *Quart. J. Roy. Meteor. Soc.*, **111**, 1039–1047, doi:10.1002/qj.49711146608.

453 Zhang, F., Y. Q. Sun, L. Magnusson, R. Buizza, S.-J. Lin, J.-H. Chen, and K. Emanuel, 2019:
454 What is the predictability limit of midlatitude weather? *J. Atmos. Sci.*, **76**, 1077–1091, doi:
455 10.1175/JAS-D-18-0269.1.

456 **LIST OF FIGURES**

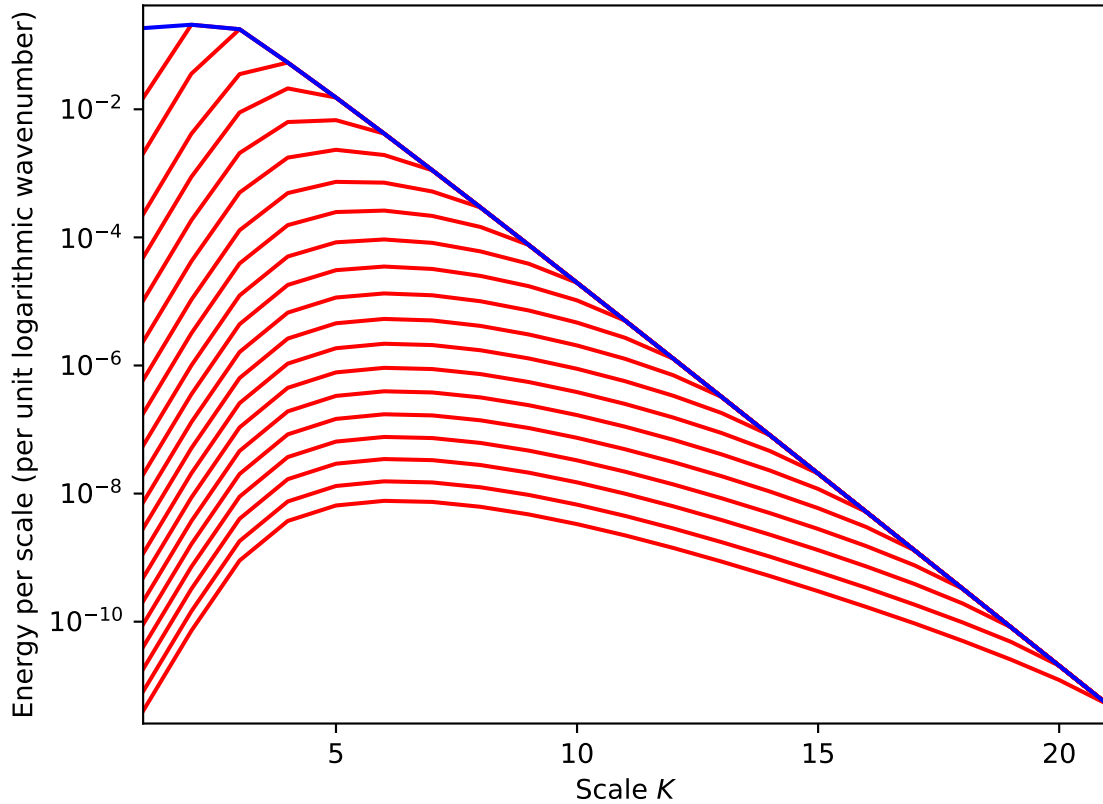
457 **Fig. 1.** The evolution of the error KE spectrum (red, from bottom to top) in the Lorenz (1969)
 458 model, with a -3 -slope background KE spectrum (appearing in this Figure as a -2 slope,
 459 as the KE per unit logarithmic wavenumber is plotted instead of the KE spectral density) as
 460 in Rotunno and Snyder (2008). The initial condition for this run is $Z(K = 20) = 2^{-40}$, $Z = 0$
 461 for all other K , and $W = 0$ for all K 25

462 **Fig. 2.** Saturation times of various scales (red) for the same model run as in Figure 1. The blue
 463 curve shows a line of fit with $\beta = 0.05$ 26

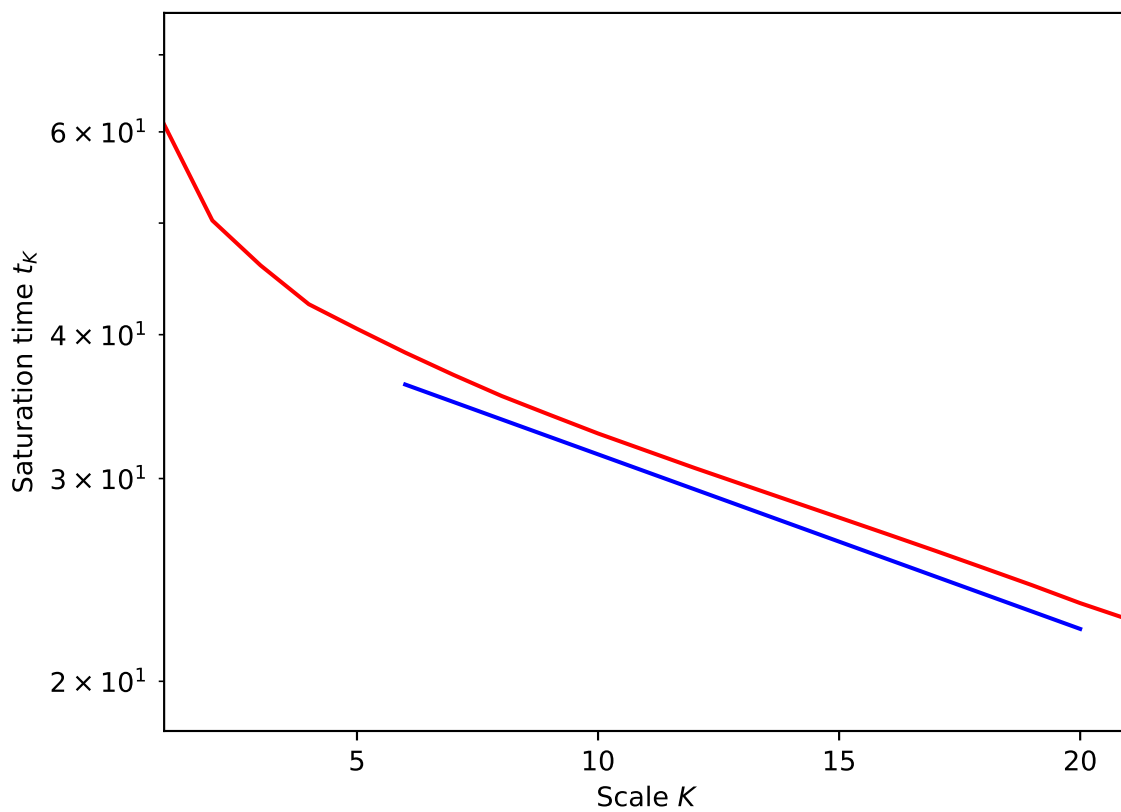
464 **Fig. 3.** KE spectrum (averaged over the five cases) of the initial condition (red), and logarithmically
 465 corrected -3 reference spectra $E(k) \sim k^{-3} \left[\log \left(\frac{k}{k_r} \right) \right]^{-\frac{1}{3}}$ ($k_r = 10$ in black, $k_r = 20$ in green),
 466 where $E(\cdot)$ is the one-dimensional KE spectral density. 27

467 **Fig. 4.** Evolution of the error KE spectrum (magenta and blue, bottom to top) for an initial pertur-
 468 bation (blue dot) at $k_p = 256$. The magenta curves are for $t = 0.3, 0.6, \dots, 2.7$ and the blue
 469 curves are for $t = 3, 6, \dots, 66$. The background KE spectra at $t = 0, 3, 6, \dots, 66$, scaled by a
 470 factor of 2, are shown in red (top to bottom), with the reference spectra in black and green
 471 as in Figure 3. The spectra are averaged over the five cases. 28

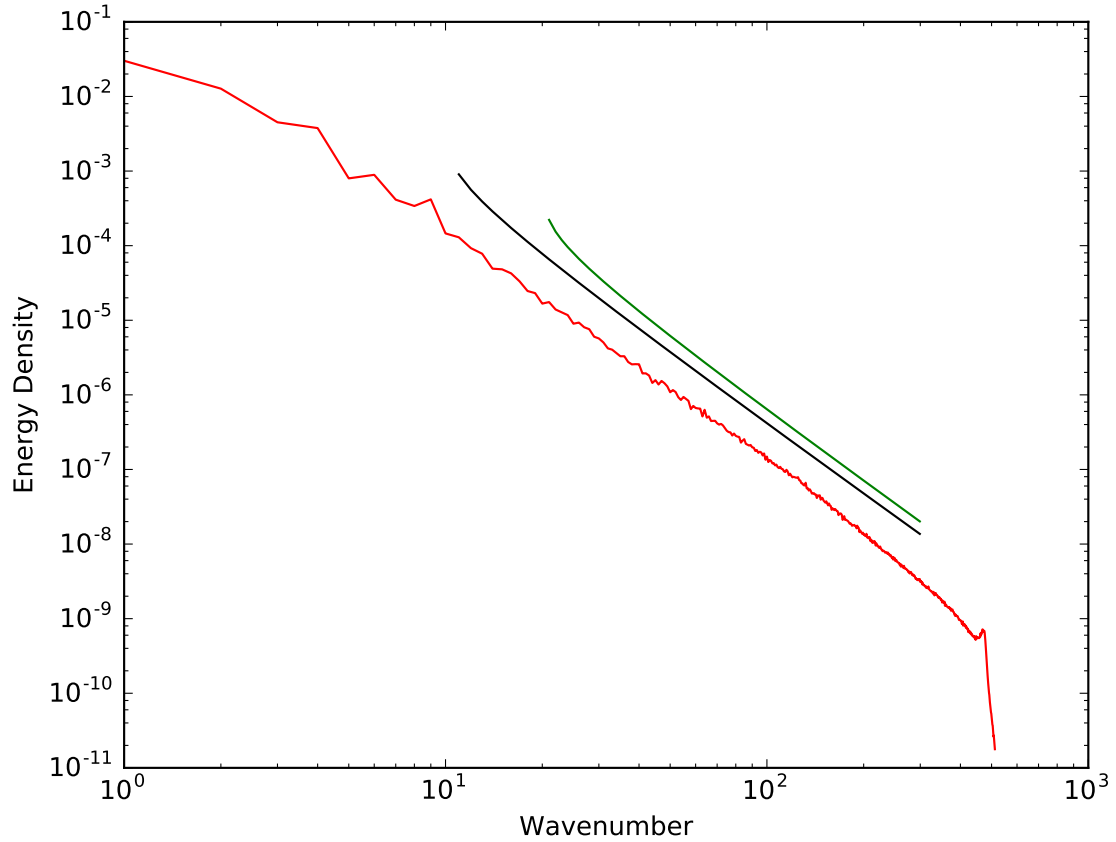
472 **Fig. 5.** Saturation times T at different wavenumbers k (red) for an initial error at wavenumber $k_0 =$
 473 256 , averaged over the five cases. The black curve shows a line of fit with $\beta = 0.24$ 29



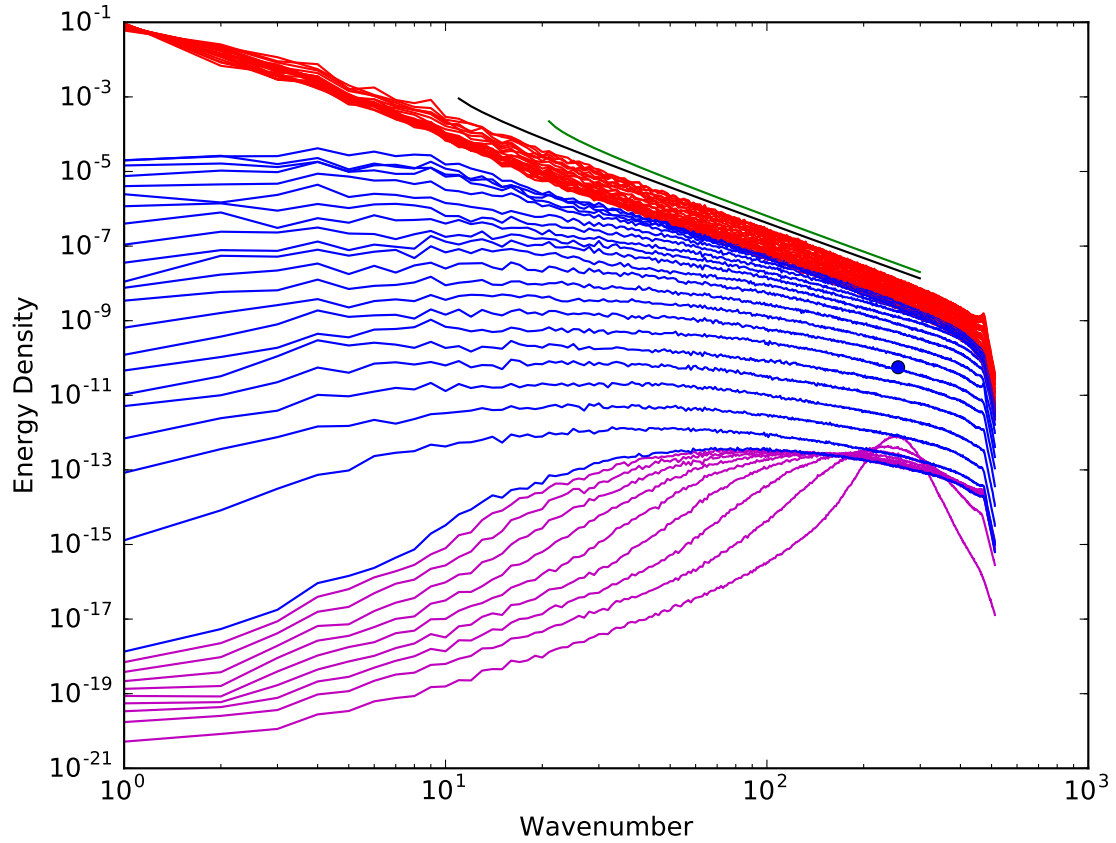
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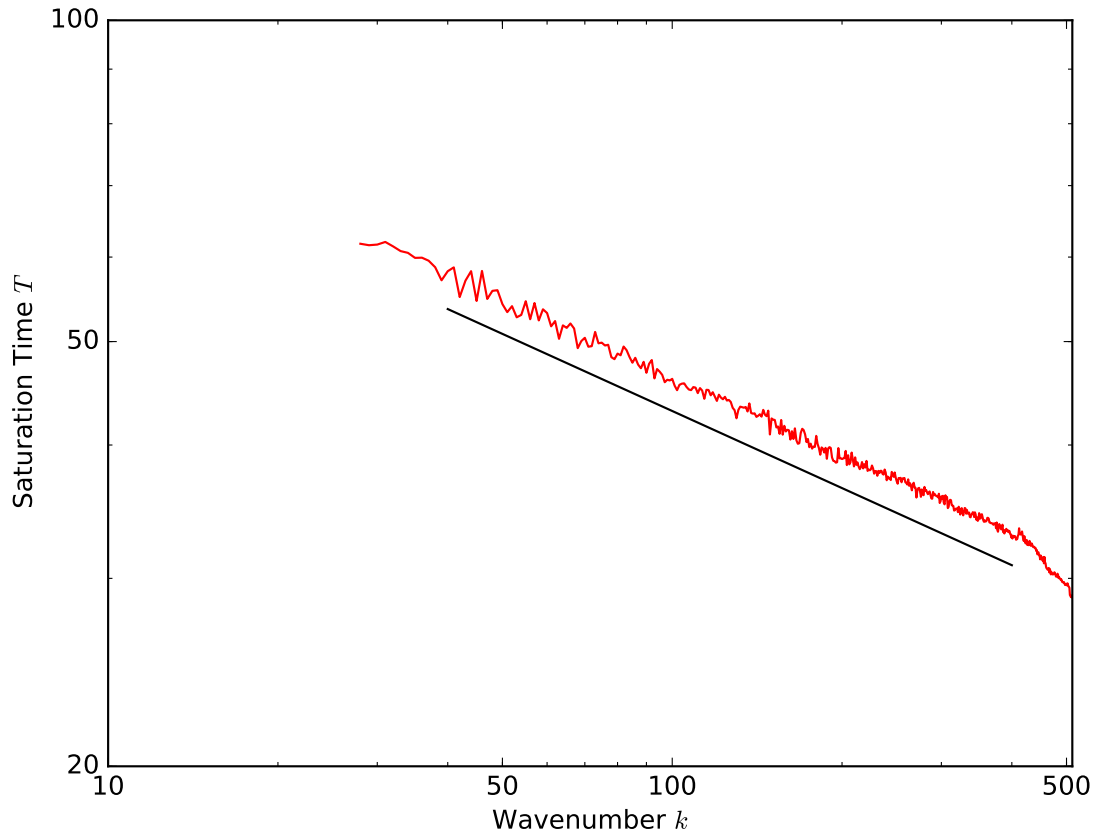
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 482 dimensional KE spectral density.



483 FIG. 4. Evolution of the error KE spectrum (magenta and blue, bottom to top) for an initial perturbation (blue
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 486 reference spectra in black and green as in Figure 3. The spectra are averaged over the five cases.



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