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Underestimation of global photosynthesis in Earth System Models due to representation of vegetation structure

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Key Points:

- Earth System Models (ESMs) may significantly underestimate global photosynthesis because they do not take vegetation structure into account.
- Introducing vegetation clumping into ESMs with multi-layered canopy schemes alleviates light limitation of photosynthesis at lower canopy levels.
- In our study, the addition of vegetation clumping into the land surface scheme of the UKESM resulted in an additional uptake of carbon by photosynthesis of 5.53 PgC yr\textsuperscript{-1} globally and 4.18 PgC yr\textsuperscript{-1} between 20°S-20°N latitude.
Abstract

The impact of vegetation structure on the absorption of shortwave radiation in Earth System Models (ESMs) is potentially important for accurate modelling of the carbon cycle and hence climate projections. A proportion of incident shortwave radiation is used by plants to photosynthesize and canopy structure has a direct impact on the fraction of this radiation which is absorbed. This paper evaluates how modelled carbon assimilation of the terrestrial biosphere is impacted when clumping derived from satellite data is incorporated. We evaluated impacts of clumping on photosynthesis using the Joint UK Land Environment Simulator, the land surface scheme of the UK Earth System Model. At the global level, Gross Primary Productivity (GPP) increased by 5.53 ± 1.02 PgC yr\(^{-1}\) with the strongest absolute increase in the tropics. This is contrary to previous studies that have shown a decrease in photosynthesis when similar clumping data sets have been used to modify light interception in models. In our study additional transmission of light through upper canopy layers leads to enhanced absorption in lower layers in which photosynthesis tends to be light limited. We show that this result is related to the complexity of canopy scheme being used.

Plain Language Summary

Plants need sunlight to photosynthesize; however, the way in which light absorption is typically described by climate models is not very realistic because it does not take into account structural differences in forest canopies. Identifying more realistic ways to represent these processes in forests would allow us to better predict photosynthesis and to have a greater understanding of the impact of future climate change. In our paper we discuss a method to include information about vegetation structure derived from satellites in climate models. Our results indicate that such models underestimate the amount of light reaching plants in the lower layers of dense forests. Consequently, global photosynthesis is likely underestimated in climate models due to a lack of consideration of plant structural variability.

1 Introduction

Understanding the global carbon cycle is critically important for understanding current and future climate change. The terrestrial biosphere sequesters around 25% of anthropogenic carbon emissions (Le Quéré et al., 2018) but there remains uncertainty around exactly what processes drive this (Ciais et al., 2019) and whether or not this sink will be maintained in the future. A reduction in sink strength due to climatic factors could be a significant positive feedback to climate change. To be able to model the future evolution of this uptake of carbon requires the ability to correctly model the underlying processes. This paper focuses specifically on photosynthesis in the terrestrial biosphere and the how we model the light interception in plants which drives this.

The uptake of carbon by terrestrial photosynthesis is the largest component flux in the global carbon cycle. Despite this its overall magnitude and global spatial distribution remains poorly understood. Estimates of Gross Primary Productivity (GPP) in the literature range from 120 PgC yr\(^{-1}\) to 175 PgC yr\(^{-1}\). The estimate of global GPP presented in the first IPCC report was set in the interval 90-120 PgC yr\(^{-1}\) (Watson et al., 1990) followed by all the other IPCC reports giving a fixed global GPP value equals to 120 PgC yr\(^{-1}\) (Melillo et al., 1995; Prentice et al., 2001; Denman et al., 2007). More recently, the last IPCC report (Ciais et al., 2013) updated the value of global GPP to 123 ± 8 PgC yr\(^{-1}\) based on model tree ensemble (MTE) and Eddy Covariance...
(EC) flux (Beer et al., 2010); however, the value of GPP strongly depends on the method used and they often disagree in long-term trends and spatial patterns (Jung et al., 2011; Anav et al., 2015; Jiang & Ryu, 2016; Knauer et al., 2017; He, Chen, Liu, et al., 2017; He, Chen, Croft, et al., 2017; MacBean et al., 2018).

For the global carbon budget 2007–2016, an imbalance of 0.6 PgC yr$^{-1}$ was estimated, indicating possible underestimated values in carbon sinks, such as global photosynthesis (Le Quéré et al., 2018). Welp et al. (2011) estimated global GPP to be somewhere in between 150-175 PgC yr$^{-1}$ based on $^{18}$O/$^{16}$O, and in a study based on $^{13}$C it was found that more GPP should be attributed to the Amazon region (Chen et al., 2017). Koffi et al. (2012) presented a study based on data assimilation with atmospheric CO$_2$ and ecosystem models with an estimated global GPP of 146 ± 19 PgC yr$^{-1}$. More recent studies using Solar Induced Florescence (SIF) predicted global GPP to be 144 PgC yr$^{-1}$; closer to most Earth System Models (ESMs) estimates than the MTE or MODIS data sets (Anav et al., 2015).

A key process required for modelling photosynthesis is the interception of light, which is typically achieved using a vegetation radiative transfer (RT) model. A commonly used vegetation RT model in many ESMs is the two-stream scheme of Sellers (1985) and a key assumption in the Sellers scheme is that leaves are randomly arranged in a plane parallel medium. This assumption is in common with many other vegetation RT schemes and is almost ubiquitous amongst those used in Climate and Earth System models. In reality, however, vegetation does not arrange itself in such a perfectly random fashion. An important question, therefore, is to ask what extent this assumption affects predictions of the photosynthetic flux of carbon into the land surface.

A simplification that results from the plane-parallel turbid medium approximation is a lack of representation of gaps in the canopies. The term ‘gaps’ is used here in the sense of ‘openness’, i.e., canopy openings, which light goes through without being intercepted. For most natural forest stands, savannah, and shrubland, various sizes of gaps exist between and within tree crowns. Neglecting these gaps has been shown to result in errors when estimating shortwave radiation interception.

Previous studies have shown that two-stream schemes can exhibit significant biases in comparison to more accurate 3D radiative transfer models and observations (Pinty et al., 2006; Ni-Meister et al., 2010; Kobayashi et al., 2012; Loew et al., 2014; Hogan et al., 2018). Despite this two-stream schemes remain attractive due to their computational efficiency. Highly detailed 3D radiative transfer models exist but they cannot be directly used in ESMs due to their computational expense (Yang et al., 2001) and the large number of parameters required (Loew et al., 2014). One approach to account for 3D canopy structure in two-stream schemes is to include simple parameterizations of 3D effects on shortwave radiation partitioning (Pinty et al., 2006). In the present study, we modify JULES, the land surface scheme of the UK Earth System Model (UKESM), to ingest a global dataset of canopy clumping derived from satellite data, in order to determine the impact of vegetation canopy structure on modelled global photosynthesis.

Although this modification results less light absorption by vegetation in the model it also drives an additional uptake of carbon by photosynthesis of 5.53 PgC yr$^{-1}$ globally and 4.18 PgC yr$^{-1}$ between 20°S–20°N latitude. The primary mechanism we attribute this to is the increased amount of light reaching lower layers of the canopy in which photosynthesis tends to be limited by available light. Conversely photosynthesis in the upper canopy layers, which absorb less light once clumping is included, are not typically light limited and so the reduction in absorbed radiation has less impact.
2 Models and Methods

2.1 Model description

The most commonly used method to account for structure in a vegetation RT model is to introduce a clumping index (Ω) (Nilson, 1971) to scale leaf area index (LAI). This can be easily implemented into the two-stream scheme; wherever LAI appears in the equations it is scaled by Ω. Hence the two stream equations become:

\[
\frac{d\mu}{d\text{LAI}} + [1 - (1 - \beta)\omega]I^\uparrow - \omega \beta I^\downarrow = \omega \mu \beta_0 \exp(-KL \cdot \Omega),
\]

\[
\frac{d\mu}{d\text{LAI}} + [1 - (1 - \beta)\omega]I^\downarrow - \omega \beta I^\uparrow = \omega \mu (1 - \beta_0) \exp(-KL \cdot \Omega)
\]

where \(I^\uparrow\) and \(I^\downarrow\) are the upward and downward diffuse radiative fluxes normalised by the incident flux at the top of the canopy, \(\mu\) is the cosine of the Sun zenith angle, or the incident beam, \(K\) is the optical depth of direct beam per unit leaf area and is equal to \(G(\mu)/\mu\), where \(G(\mu)\) is the projected area of leaf elements in the direction \(\cos^{-1}(\mu)\) (Ross, 1981), \(\mu\) is the average inverse diffuse optical depth per unit leaf area, \(\omega\) is the scattering coefficient and is given by \(\rho_{\text{leaf}} + \tau_{\text{leaf}}\), the leaf reflectance and transmittance respectively, and \(L\) is the cumulative LAI from the top of the canopy. \(\beta\) and \(\beta_0\) are upscattering parameters for the diffuse and direct beams, respectively. In this context \(\Omega\) corresponds to the structure factor described in Pinty et al. (2006) except that it is assumed not to vary with zenith angle.

The Joint UK Land Environment Simulator (JULES; Best et al., 2011; Clark et al., 2011), is the land surface scheme of the new UK Earth System Model (UKESM). It uses the Sellers RT model to calculate light interception and absorption in vegetation. The option to include a clumping index was added in version 4.6 with a default value of 1.0 (i.e., no clumping), allowing user to prescribe other values where data is available. The variable was originally implemented in JULES with a single value per plant functional type (PFT), and it was tested and evaluated over crops by Williams et al. (2017) who showed that it was necessary to include clumping (i.e. \(\Omega<1.0\)) to correctly model the productivity of maize for a field site in Nebraska, USA. For this paper we modified JULES to read in a spatially varying map of clumping for each PFT (described in Section 2.2).

We used JULES version 4.6 with the Global Land (GL) 4.0 configuration (Walters et al., 2014) with the WATCH-Forcing-Data-ERA-Interim data set (Weedon et al., 2014) at 0.5° spatial resolution and temporal resolution of 3 hours. The Harmonized World Soil Database version 1.0 data set (Nachtergaele et al., 2008) and the model of runoff production (TOPMODEL) were applied following Clark and Gedney (2008). Leaf area index was determined prognostically by the JULES phenology module (Cox, 2001) updated every 10 days. Prior to performing the global scale model simulations, the soil moisture and temperature were brought to equilibrium using a 5-year global spin-up by cycling 1 year of meteorological data. JULES GL4.0 uses 5 PFTs: broadleaf trees, needle-leaf tress, C3 grasses, C4 grasses, and shrubs.

By default, JULES computes light interception and photosynthesis in 10 vertical canopy layers. Leaf-level photosynthesis in each layer is estimated as the minimum rate of three assimilation regimes as proposed by Farquhar et al. (1980) and modified by Collatz et al. (1991, 1992): (i) the Rubisco-limited rate or carbon limiting regime; (ii) the light-limited rate, and; (iii) the carbon compound export limitation for C3 plants or PEP-carboxylase export limitation for C4 plants, referred to as the electron transport or export limiting regime. The multilayer approach
simulates the transition between the Farquhar limiting regimes at each canopy layer, resulting in increased carbon limitation towards the top of the canopy and increased light limitation towards the bottom of the canopy (Clark et al., 2011). The light-limited rate of photosynthesis in each layer is proportional to the fraction of absorbed photosynthetically active radiation (fAPAR) in that layer. Consequently, including clumping in the radiative transfer scheme directly affects the light limited rate of photosynthesis but not the Rubisco or export limited rates.

We performed two runs of JULES for the year 2008 with and without a prescribed value of clumping index. The year was chosen as it is close to the date of production of the clumping map (i.e., 2006) and an ENSO neutral year, unlike 2006-2007, which was a weak El Niño.

### 2.2 Global clumping index map

The global clumping map of He et al. (2012) was used to provide clumping index data for JULES. It has a spatial resolution of 500 m and was produced for the year of 2006. We assume that the global clumping index map in 2006 is still reliable for modelling GPP in 2008 since the inter-annual variation of clumping index is general small (He et al., 2016). The data were derived from the NASA-MODIS BRDF/albedo product (MCD43) by considering the difference in forward and backward scattering from the surface, which is primarily controlled by the structure of the vegetation. This follows the methodology of Chen et al. (2005) but with an additional correction to the magnitude of the MODIS hotspot. The method uses a 4-Scale BRDF model (Chen et al. 1997) that considers different scales of canopy clumping: tree groups, tree crowns, branches and shoots. This is equivalent to the assumptions implicit in clumping as implemented in JULES. Pinty et al. (2006) provide a detailed discussion of this type of clumping index as applied to two-stream models.

We scaled up the He et al. (2012) data to the resolution of the model run (0.5°) on a per-PFT basis by using the GLC2000 land cover data (Bartholome & Belward, 2005). The GLC2000 is also used in the production of the clumping dataset and to prescribe the distribution of the 5 PFTs used by JULES. The total clumping index map is shown in Figure 1. Values less than one indicate clumping, with smaller values indicating greater clumping. The most clumped areas are the boreal forests and areas with sparse vegetation, while the least clumped areas are in the presence of grasses, e.g., over savannahs in Africa and crops in the USA and Asia. Tropical forests show intermediate levels of clumping, which does not fit with many below canopy observations of clumping. He et al. (2012) argue that ground based measurements generally underestimate clumping in dense forests (i.e., overestimate the clumping index value) because they are overly affected by lower-level branches. Pisek et al. (2013) further confirmed that in moderate to dense forests with developed bottom layers, in situ measurements of clumping near the surface tend to considerably underestimate the overall canopy-level clumping. Olivas et al. (2013) found that the mean LAI above 1 m using litter-fall collection was 5.54 ± 0.3 at an old-growth tropical rainforest, while the effective LAI from hemispherical photographs was only 3.45 ± 0.1, implying a clumping index of 0.62.

### 2.3 Benchmarking data

We used the MTE global GPP dataset (Jung et al., 2011) as a reference. It is a monthly global data product at 0.5° resolution which uses a statistical method based on machine learning techniques referred to as model tree ensembles (MTE). The MTE global GPP was trained against flux tower GPP estimates at site level using fAPAR from satellite observations and meteorological data as explanatory variables. Site level GPP estimates from 178 FLUXNET sites
were incorporated in the production of the data following quality filtering and partitioning of net ecosystem exchange into GPP and ecosystem respiration based on Lasslop et al. (2010). The MTE product is available since 1982 but it is important to interpret it carefully since flux tower observations started a decade after that with a limited number of sites sparsely distributed and mainly across Europe and North America. Therefore, there is a large uncertainty of the MTE GPP over regions with limited flux tower sites including most parts of Africa and South America, as well as Tropical and Northern Asia (Anav et al., 2015).

3 Results

3.1 The impact of vegetation canopy structure on modelled global fAPAR

The first order impact of adding clumping to the vegetation radiative transfer scheme in JULES is to reduce fAPAR. Figure 2 shows a global map of fAPAR differences between JULES with and without clumping included. fAPAR decreases across the entire globe when clumping is added because it acts to decrease the effective leaf area available to intercept light. However, in addition to reducing the overall fAPAR, the relative distribution of absorption vertically through the canopy is also modified. Layers at the bottom canopy have more light directly incident upon them due to greater transmission through the layers above and therefore can potentially also absorb more PAR. Because clumping is applied to all layers evenly, each layer absorbs proportionally less of the PAR directly incident upon it, but the total amount of incident PAR on layers except for the top one will always be increased relative to the model without clumping. Hence the total absorption of PAR in a layer can increase even though its fAPAR decreases as long as there is sufficient additional radiation reaching it.

The average value of fAPAR for the globe in 2008 according to JULES without clumping is 0.607 ± 0.022 (95% confidence interval). Applying the clumping index shifts the average value to 0.576 ± 0.021, or the equivalent of a total average decrease of 0.032 ± 0.002 (-5.3%). Some locations of the Earth have much larger divergences in fAPAR, for instance Southwest Canada and Northwest USA, Northeast Russia, and high-altitude regions such as the Himalayas and the Andes; these are areas typically associated with needle-leaved trees.

3.2 The impact of vegetation canopy structure on global GPP

In our model experiment the addition of clumping systematically increases carbon assimilation throughout the globe, resulting in an additional 5.53 ± 1.02 PgC yr⁻¹ in GPP. Figure 3 shows the difference in GPP between JULES with clumping (JULES-Clump) and the default version of JULES (i.e., without clumping). The strongest difference between the two model setups is found in the tropics (20°S - 20°N) with additional GPP of 4.18 PgC yr⁻¹, or 75% of the total additional carbon, followed by 1.10 PgC yr⁻¹, or approximately 20% of the total extra GPP in the Northern Hemisphere (20°N - 90°N), and 0.25 PgC yr⁻¹ in the Southern Hemisphere (90°S - 20°S), which corresponds to approximately 5% of the total extra GPP.

Figure 4a shows the difference in the absolute difference between JULES-Clump and MTE-GPP, and JULES and MTE-GPP. Regions in blue indicate that including clumping moves the JULES prediction toward the MTE estimate, and red areas indicate the opposite, i.e., JULES-Clump presents larger discrepancies than JULES in comparison to the MTE-GPP product. Tropical forests, the temperate forests in North America, and most of the boreal forests generally move closer the MTE data in JULES-Clump. The red areas on Figure 4a, associated with increasing differences between the MTE and modelled GPP prediction when clumping is
included are mainly found in the African and Brazilian savannahs, and sparser areas in the presence of grasses, especially C4 grasses. C4 grasses have previously been shown to be over-productive in JULES (Harper et al., 2016) and adding clumping makes it more productive in this study. It is also likely that the MTE data set itself shows an inaccurate representation of GPP for C4 grasses, since this PFT is not well sampled in the eddy covariance data that the MTE dataset is based upon.

Figure 4c shows the total GPP in PgC yr⁻¹ for each box in Figure 4b for the MTE-GPP product, JULES and JULES-Clump, respectively. The error bars on the MTE product were calculated as the weighted sum of the averaged standard deviation of the ensemble mean of the 25 best model trees associated with the MTE-GPP product for the year 2008. JULES-Clump consistently shows a higher GPP than the default JULES for all the evaluated areas with a larger absolute impact over the boxes in the tropics, i.e., Central and South America, ΔGPP = 2.03 PgC yr⁻¹, or 36.7% of the total additional GPP generated by the addition of clumping, followed by Africa, ΔGPP = 1.10 PgC yr⁻¹, or 19.9% of the total additional GPP, and South and Southeast Asia, ΔGPP = 1.05 PgC yr⁻¹, or 19.0% of the total additional GPP. Alone, the tropics are responsible for an extra 4.18 PgC per year (75.6% of the global ΔGPP).

Globally, the 5.53 PgC yr⁻¹ caused by the inclusion of vegetation clumping is equivalent to an additional 4.8% of GPP for the year of 2008. Although for the majority of regions in Figure 4b JULES GPP are within the error bars of the MTE product, JULES-Clump is closer to the estimates of the MTE, except for Africa, where JULES is lower than the MTE GPP and JULES-Clump is higher than it. The most significant change is observed over Central and South America where the prediction of GPP without clumping is low compared to the MTE GPP.

The additional GPP resulting from including clumping is not evenly distributed vertically though the canopy. The difference in zonal mean GPP in each canopy layer between JULES with and without clumping is shown in Figure 5. In particular there is a strong enhancement of GPP in the lower canopy layers for the tropics, whereas the top 3 or 4 layers exhibit reduced photosynthesis. This is caused by the increase in PAR absorption in the lower layers described in Section 3.1. Because these layers tend to be light limited this results in a significant boost to the overall canopy photosynthesis compared to the upper layers which are generally not light limited (Jogireddy et al., 2006; Mercado et al., 2007; Alton et al., 2007; Huntingford et al., 2008). For the bottom two layers of the canopy GPP increased more than 50% throughout all latitudes. This adds further weight to the arguments of He et al. (2018) who highlight the importance of shade leaves in global photosynthesis.

3.3 Is the impact of vegetation canopy structure on global GPP impacted by diffuse radiation?

Throughout all simulations performed in this study the percentage of diffuse incident shortwave radiation was held constant and equal to 40% as a proxy average value for the whole globe (Harper et al., 2016). However, the consideration of gaps through the addition of clumping into the radiative transfer scheme in JULES can enhance the amount of shortwave radiation reaching bottom layers of the vegetation canopy, as previously discussed. This is true for both natures of incident light, i.e., either direct, collimated beams, or diffuse, isotropic shortwave radiation. However, is the impact of canopy clumping on GPP affected by the amount of diffuse radiation?

In order to verify the effect of diffuse light on the impact of clumping on GPP, a test was performed for 12 FLUXNET sites for the year of 2008 with JULES and JULES-Clump for four
different ratios of incident diffuse shortwave radiation to global incident shortwave radiation: 20%, 40% (used in all the other runs), 60%, and 80% (Figure 6). Results indicate that across all the evaluated sites, differences in monthly mean GPP fluxes between JULES and JULES-Clump are independent of the amount of diffuse shortwave radiation reaching the surface.

One observable pattern in Figure 6 is the diffuse fertilisation effect (Mercado et al., 2009) clearly noticed in sites with high LAI values, e.g., the tropics (BR-Ma2, BR-Sa1) throughout the year, temperate forests (JP-Tak, US-MMS, US-Ha1) and boreal/needle-leaved forests (FI-Hyy, FI-Kaa, DE-Tha) in Summer time, while places with smaller or no noticeable differences between JULES and JULES-Clump are noticed in sites that are not limited by light, because of small LAI values, often associated with drier/grassland sites (ES-Es1, ES-LMa, US-FPe).

4 Discussion

The only other study of which we are aware that tackles this question on a global scale is that of Chen et al. (2012). In that study the authors used a related dataset to prescribe clumping in the Boreal Productivity Simulator (BEPS; Liu et al., 1997) but found that global GPP was reduced by 12.1 PgCyr⁻¹. The critical difference between our study and that of Chen et al. (2012) is in the treatment of canopy radiative transfer: our model uses multiple canopy layers each with different proportions of sunlit and shaded leaves, whereas Chen et al. (2012) use a single layer split into sunlit and shaded leaves (a so-called ‘two-leaf’ model). As discussed in the previous sections our result is caused by the greater penetration of light into lower layers boosting photosynthesis in layers that are light-limited. The phenomena we illustrate is to some degree analogous to the so-called “diffuse light fertilisation effect” which has been shown previously to enhance global GPP after the eruption of Mount Pinatubo in 1991 (Mercado et al., 2009). Diffuse light is able to penetrate further down into the canopy than direct beam irradiance.

Single layer models cannot redistribute absorbed radiation vertically in the manner we have shown using a layered canopy model. Consequently, there is no preferential alleviation of light-limited photosynthesis at lower levels and no boost to overall canopy photosynthesis.

Other examples of models with multi-layered canopy schemes include EALCO (Wang et al., 2001), EDv2.1 (Medvigy et al., 2009), LPJ-GUESS (Smith et al., 2001; Smith et al., 2014), SDGVM (Woodward et al., 1995), and TECO (Wang et al., 2009), and so, similar results would be expected from these models assuming the model structure allows for the inclusion of clumping in the canopy radiation scheme.

There is some empirical evidence from field based studies that supports our finding that structure increases GPP (Ahl et al., 2004; Duursma & Makela, 2007; Hardiman et al., 2011; Bohn & Huth, 2017). Hardiman et al. (2011) showed departures from randomness in forest canopies boosted productivity in a transition zone between boreal forests and Northern mixed hardwood. The authors suggested that changes in canopy structure can contribute to resilience of the functioning of ecosystems trees. Atkins et al. (2018) affirms that the inclusion of canopy structural complexity metrics in canopy light absorption models could increase confidence in predictions of biogeochemical cycles and energy balance. Their study including sites from the National Ecological Observation Network and university field stations found that canopy structure was strongly coupled with fAPAR under high light environments, while under low light conditions, when diffuse light predominates, light scattering weakened the dependency of fAPAR on structure. Also, the authors found that a multivariate model including parameters of canopy structure and leaf area index explained around 89% of the inter-site variance in fAPAR. Another observational study by Fahey et al. (2016) found an important contribution of bottom
layers of a North American site to canopy productivity as whole. The authors found a connection
between sub-canopy tree growth and fAPAR, and indicated a relationship between sub-canopy
light availability and canopy structure. Although, they found that sub-canopy growth response
was not mediated by fAPAR alone.

On a global scale it will be necessary to provide observation of canopy structure from
remote sensing instruments. The method of He et al. (2012) can, potentially, be repeated for
every year of the MODIS archive and is applicable to other missions with similar characteristics
such as Sentinel-3. Arguably, however, this problem also needs addressing using observations
that are more directly related to forest structure such as space borne LiDAR from missions such
as NASA GEDI (Hancock et al., 2019), or long wavelength RADAR from JAXA’s ALOS
PALSAR the upcoming ESA Biomass mission. Terrestrial and airborne observations will also
be critically important (Longo et al., 2016; Ferraz et al., 2018; Rödig et al., 2018) and the
increased interest in terrestrial scanning LiDAR may help to answer some of these questions
(Disney et al., 2010; Mulatu et al., 2019).

Our result that tropical photosynthesis is being underestimated in JULES likely applies to
the terrestrial biosphere components of all ESMs. Multilayered models will respond in the same
way when clumping is introduced, i.e. with greater penetration of light to lower levels. We also
argue that single layer models do not represent the impact of clumping of photosynthesis
correctly. It is clear, however, that much more investigation is required to understand the correct
way to represent structure in these models. The technique used in JULES to include clumping is
relatively simple albeit based on well-established theoretical considerations (e.g., Nilson 1971).
We note that there are more sophisticated approaches available (Kucharik et al., 1999; Pinty et
al., 2006; Ni-Meisters et al., 2010) but these rely on additional parameters which must either be
calculated by the underlying land surface model, or input as ancillary data.

5 Conclusion

Our work suggests that ESMs may significantly underestimate photosynthesis in tropical
forests because they do not take vegetation structure into account. The dominant effect that
introducing clumping has in our study is to alleviate light limitation at lower canopy levels. This
tends to have the greatest impact where leaf area index is high and where photosynthesis is not
limited by light in higher canopy layers. In our study this effect accounted for an additional
uptake of carbon by photosynthesis globally of 5.53 PgCyr⁻¹ and 4.18 PgCyr⁻¹ between 20°S-
20°N latitude.

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(2012) is available for download through the following
link: https://daac.ornl.gov/VEGETATION/guides/Global_Clumping_Index.html. The global
clumping index map per PFT used in this study and a UNIX patch file for JULES version 4.6 is
available from https://github.com/braghiere/JULES-Clump. The JULES code is available from
the UK Met Office code repository: https://code.metoffice.gov.uk/. Model simulation results are
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**Figure captions**

**Figure 1.** Global map of the MODIS derived clumping index at 0.5° resolution for the year 2006 scaled up from the 500m He et al. (2012) clumping dataset.

**Figure 2.** Spatial distribution of total fAPAR difference (JULES with clumping – JULES without clumping) for the year of 2008.

**Figure 3.** Difference in GPP between JULES with clumping and JULES without clumping. Global average values are indicated at the bottom of the figures in PgC yr\(^{-1}\) with the 95% confidence interval. Grey areas represent regions with no data.

**Figure 4. a.** The difference in the absolute GPP between JULES without clumping and the MTE data, and JULES with clumping and the MTE data. Regions in blue indicate model improvement by addition of vegetation clumping; **b.** map showing the regions used in the analysis; **c.** Total (area weighted sum over box area) JULES (green), JULES-Clump (green) and observation based (MTE; black dots and error bars) GPP fluxes for the year of 2008 at regional scales. Error bars indicate the weighted sum of the averaged standard deviation of the ensemble mean of the 25 best model trees associated with the MTE-GPP product.

**Figure 5.** Zonal mean vertical profile of **a.** absolute and difference in GPP between JULES-Clump and JULES without clumping; **b.** Total GPP zonal mean of MTE, JULES-Clump (red), and JULES (blue). 1 standard deviation (±1σ) of the spatial mean for each product is represented by the filled areas.

**Figure 6.** Monthly mean fluxes of GPP for 12 FLUXNET sites from JULES (continuous line) and JULES-Clump (dashed line) for four different percentages of incident diffuse shortwave radiation: 20% (red), 40% (blue), 60% (green), and 80% (yellow).